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Black carbon sources constrained by observations in the Russian high Arctic

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26 **ABSTRACT:** Understanding the role of short-lived climate forcers like black carbon (BC) at high northern latitudes in climate change is hampered by the scarcity of surface 27 observations in the Russian Arctic. In this study, highly time resolved Equivalent BC (EBC) 28 29 measurements during a ship campaign in the White, Barents and Kara Seas in October 2015 are presented. The measured EBC concentrations are compared with BC concentrations 30 simulated with a Lagrangian particle dispersion model coupled with a recently completed 31 global emission inventory to quantify the origin of the Arctic BC. EBC showed increased 32 values (100–400 ng m⁻³) in the Kara Strait, Kara Sea, and Kola Peninsula, and an extremely 33 high concentration (1000 ng m⁻³) in the White Sea. Assessment of BC origin throughout the 34 35 expedition showed that gas flaring emissions from the Yamal/Khanty-Mansiysk and Nenets/Komi regions contributed the most when the ship was close to the Kara Strait, north of 36 37 70°N. Near Arkhangelsk (White Sea), biomass burning in mid-latitudes, surface 38 transportation, and residential and commercial combustion from Central and Eastern Europe 39 were found to be important BC sources. The model reproduced observed EBC concentrations efficiently, building credibility in the emission inventory for BC emissions at high northern 40 latitudes. 41

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43 **1. INTRODUCTION**

Short-lived climate forcers are aerosols and gases that cause radiative forcing ^{1,2} and have 44 lifetimes of less than a few years ^{3,4}. Light-absorbing aerosols are of particular interest, since 45 they have a warming effect that is strongest over highly reflective surfaces (e.g., clouds, snow 46 and ice)⁵. Furthermore, their deposition on snow and ice decreases surface albedo, which can 47 enhance melting ^{6,7} and trigger surface warming. Most of the radiation absorption of 48 accumulation-mode aerosol is due to black carbon (BC)⁸. BC also influences cloud radiative 49 properties ^{9,10}. BC originates from incomplete combustion, e.g., of biomass or fossil fuels ^{1,11}. 50 Freshly emitted BC is hydrophobic but ageing in the atmosphere changes its properties to a 51 more hydrophilic state ¹². It is an important constituent in Arctic Haze, a phenomenon that is 52 primarily the result of long-range pollution transport from sources outside the Arctic ^{5,7,13,14}. 53 The majority of the Arctic BC originates from anthropogenic sources, especially industrial 54 applications, residential combustion, and diesel transportation activities ¹⁵, while other 55 important sources include fires in boreal forests and agricultural regions especially from 56 spring to fall $^{16-18}$. 57

Near the surface, about 50% of the BC north of 60°N originates from Russia¹⁹, where 58 emission inventory data are highly uncertain²⁰. Emissions from flaring of gas associated with 59 oil production are prone to particularly high uncertainty because both activity data and 60 61 emission factors are largely lacking. According to the Global Gas Flaring Reduction Partnership (GGFR) (http://www.worldbank.org/en/programs/gasflaringreduction), nearly 50 62 billion m³ of gas are flared in Russia annually. The Russian flaring emissions in the Yamal 63 and Khanty-Mansiysk regions are directly within the major low-altitude pathway of sub-64 Arctic air masses penetrating into the Arctic ¹⁶ and thus Stohl et al. ²¹ estimated that they 65 contribute about 42% of the annual average BC surface concentrations in the Arctic.. 66

67 However, limited measurements are available that would enable constraining this particular source of BC in the Russian Arctic. For instance, in the whole Russian territory 68 north of 50°N, continuous measurements of equivalent BC (EBC) are performed only at Tiksi 69 station (71.36°N; 128.53°E) ^{22,23}, which is far from the major industrial sources in Russia. 70 Based on isotope measurements, one recent observational study ²⁴ suggests that the 71 contribution of gas flaring emissions to BC at Tiksi is lower than estimated by Stohl et al.²¹. 72 However, new bottom-up inventories ^{25,26} contain gas flaring emissions that are even higher 73 than those used by Stohl et al.¹⁹. To clarify the role of gas flaring emissions, any EBC 74

measurements from regions closer to the oil production facilities of Russia would be extremely valuable, In these regions, BC has been measured only with low time resolution during a few ship campaigns ^{21,27–29}. However, to relate such measurements to particular source regions, measurements with high time resolution are necessary. In a comparison with the few available observations, modeled BC concentrations were found to be too low ²¹, but a comprehensive analysis was not possible because of the low time resolution of these measurements.

In the present study, we report highly time resolved EBC concentrations measured during the "Sever-2015" expedition through the White Sea, Barents Sea, and Kara Sea in October 2015. We compare the EBC measurements recorded during the cruise with predicted BC concentrations simulated with a Lagrangian particle dispersion model (LPDM). Furthermore, we investigate and quantify the origin of the BC observed during the cruise using modeling results coupled with the most recent emission inventory for BC. This is done to assess how the oil and gas industrial emissions in high northern latitudes affect Arctic BC.

89 2. METHODOLOGY

90 2.1 Expedition and Analysis of Equivalent Black Carbon. The expedition "Sever-91 2015" was carried out onboard the research vessel "Akademik Treshnikov" of the Russian Arctic and Antarctic Research Institute from 9 to 25 October 2015. The ship is the modern 92 vessel of RMRS (Russian Maritime Register of Shipping, class notation KM Arc7AUT2) and 93 it uses three propulsion WÄRTSILÄ diesel 4-stroke engines with 600 rpm. The ship track in 94 the Arctic Ocean and the research vessel are shown in Figure S 1 together with the main gas 95 96 flaring facilies. The cruise started on 10 October from the port of Arkhangelsk (64.58 N, 97 40.50°E; point A on the map), and continued through the delta of the Dvina river towards the 98 White Sea and Kanin Nos (point 1 in Figure S 1) in the Barents Sea. Then it passed the Kara 99 Strait (point 2 in Figure S 1) and the Kara Sea until it reached the archipelago Severnaya Zemlya (79.35°N, 101.83°E; point B). After a stay of two days near the research station "Ice 100 Base Cape Baranova" on the Bolshevik Island (from 15 to 17 October 2015), the ship turned 101 102 back. A storm forced the ship to moor in the Kara Sea (point 3 in Figure S 1) from 19 to 21 103 October before it could continue its return journey to Arkhangelsk, where it arrived on 25 104 October. Meteorological data (temperature, apparent wind speed and direction) during the cruise were obtained from the Vaisala maritime observation system MAWS-420. Real wind 105

direction and speed was estimated from the aforementioned data. Surface air temperature,pressure, and wind data are shown in Figure S 2, respectively.

108 Aerosol EBC concentrations were determined continuously using an aethalometer purposely designed by the Moscow State University (MSU) and Central Aerological 109 Observatory (CAO) for ship campaigns. In this instrument, light attenuation caused by the 110 111 particles depositing on a quartz fiber filter is measured at three wavelengths (450, 550, and 650 nm). The light attenuation coefficient of the collected aerosol was calculated with the 112 method of Hansen and Rosen³⁰. EBC concentrations were determined continuously by 113 114 converting the time-resolved light attenuation to the EBC mass corresponding to the same 115 attenuation and characterized by a specific mean mass attenuation coefficient. This calibration 116 parameter was derived during parallel long-term measurements against an AE33 aethalometer (Magee Scientific) that operates at the same three wavelengths (450, 550, and 650 nm). 117

118 Attenuation coefficient batn is defined as:

119
$$batn = A(m^2) \cdot \delta ATN/V(m^3) \quad (1)$$

where *A* is the filter exposed area, and *V* is the volume of air sampled and δATN is the light attenuation defined as follows:

122
$$\delta ATN = \ln(I_o/I) \quad (2)$$

where I_o and I is the light intensity transmitted through unexposed and exposed parts of the filter, respectively. Good linear correlation between the aethalometer's attenuation coefficient *batn* and the EBC concentrations calculated with the AE33 aethalometer (at 660 nm) was achieved ($R^2 = 0.92$, see Figure S 3). This allowed estimation of EBC mass concentrations using the regression slope and intercept between b_{atn} at 650 nm and EBC of the AE33 aethalometer at 660 nm:

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$$EBC(ng m^{-3}) = 3.3 \times 10^5 \cdot A(m^2) \cdot \delta ATN/V(m^3)$$
 (3)

where 3.3×10^5 is the correction factor that includes the specific mass absorption coefficient for the MSU aethalometer calibrated against the AE33 aethalometer assuming the Mass Absorption Cross-section (MAC) adopted by AE33 equal to 9.89 m² g^{-1 31}. The uncertainty of EBC measurements from both aethalometers depends on the accuracy of the MAC value used for the conversion of the light absorption coefficient to mass concentration. The constant MAC value adopted here is an approximation, assuming a uniform state of mixing for BC in atmospheric aerosol. This can be considered a valid assumption in the case of background aerosol measurements performed in this study. Absolute uncertainties of the reported MAC values remain as high as 30-70% due to the lack of appropriate reference methods and calibration materials ³².

The level of uncertainty (1-sigma) of EBC measurements was 30 ng m⁻³ for six minutes integration time. Aethalometer filters were changed manually at the latest when ATN values approached 70 but at most times filters were changed at lower values. During rough and wet weather conditions, water droplets or sea spray affected the measurements adding higher noise to the recorded ATN signal. These short data periods were either excluded from the dataset or, where possible, treated manually by establishing an adjusted baseline for the reference ATN values.

To identify the cleanest location on the vessel (i.e., the spot least influenced by the ship 147 exhaust), particulate mass (PM) concentration was measured on all decks of the vessel using a 148 149 TSI DustTrak 8530 monitor. The best site for ambient aerosol monitoring was identified to be at the foredeck, where the aethalometer was placed, while the spot most affected by the 150 151 exhaust pipe was found at about 10 m on the upper bridge (Figure S 1). A second aethalometer of exactly the same type was therefore installed at this location to record 152 potential impact from ship pollution. EBC concentrations from the two aethalometers were 153 154 compared and the absence of contamination on the foredeck, where the aethalometer was placed (clean air site), was assured. When the apparent wind was blowing from the back of 155 156 the vessel towards the clean air site on the foredeck, all aethalometer data were removed from further analysis. For instance, such contamination might have occurred when the ship moored 157 near point 3 (Figure S 1) during the storm event and therefore these measurements were 158 removed from the dataset. 159

2.2 Emissions and Modeling of Black Carbon. The concentrations of BC were 160 simulated with version 10 of the LPDM FLEXPART (FLEXible PARTicle dispersion model) 161 ^{33,34}. The model was driven with operational meteorological analyses every three hours from 162 the European Centre for Medium-Range Weather Forecasts (ECMWF). The ECMWF data 163 had 137 vertical levels and a horizontal resolution of 1°×1°. Computational particles released 164 from the measurement locations were tracked back in time in FLEXPART's "retroplume" 165 mode ³⁵. Simulations extended over 30 d back in time, sufficient to include most aerosol 166 167 emissions arriving at the station, given a typical BC lifetime (1 week). This enabled

identifying where the measured BC came from and allowed quantification of BC source 168 169 contributions. The source contributions can also be displayed as a function of the time elapsed since the emission has occurred (i.e., "age"), which can be shown as "age spectrum" 170 171 consisting of stacked bars, where a bar's color indicates the contribution of a certain age bin 172 (0-1 days, 1-2 days,..., 29-30 days) (see Figure 1b). FLEXPART simulations were performed every hour during the cruise, with particles released from small boxes covering the latitude 173 174 and longitude ranges of the ship track during the hour. The FLEXPART retroplumes consist 175 of an emission sensitivity (often also called source-receptor relationship), which yields a 176 simulated concentration in the receptor box when multiplied with gridded emissions from an 177 inventory.

Emission fluxes were taken from the ECLIPSE (Evaluating the CLimate and Air Quality ImPacts of ShortlivEd Pollutants) version 5 emission dataset ³⁶, which is available from the website of the International Institute for Applied Systems Analysis (IIASA) (http://www.iiasa.ac.at/web/home/research/researchPrograms/air/Global emissions.html).

This inventory is appropriate for use in our study, as it accounts for BC emissions from gas 182 flaring from the main emitting facilities located west of Yamal Peninsula (Komi and Nenets 183 distinct) and in Khanty-Mansiysk (south of Yamal Peninsula)²¹. Biomass burning (BB) 184 sources, namely forest, peat, savanna, woodland fires, and from deforestation were adopted 185 from the Global Fire Emissions Database, Version 3 (GFEDv3.1)³⁷. As regards to 186 anthropogenic sources, it includes industrial combustion and processes sector (IND) 187 emissions from combustion happening in industrial boilers as well as emissions from 188 189 industrial production processes. Residential and commercial sector (DOM) includes emissions 190 from combustion in heating and cooking stoves and boilers in households and public and 191 commercial buildings like malls, hospitals and schools. Waste treatment and disposal sector 192 (WST) includes emissions from waste incineration and the treatment process. Transport sector (TRA) includes emissions from all land based transport of goods, animals and persons on 193 194 road networks as well as off-road activities e.g. on railroads, agricultural and forest lands, construction sites. Shipping in in-land waters and domestic aviation are also included in this 195 sector, but international shipping and aviation are treated as separate sectors. Finally, energy 196 197 production and distribution sector (ENE) includes emissions from combustion processes in 198 power plants and generators, emission related to distribution of energy to consumers, as well 199 as emissions from gas flaring in oil facilities.

For our simulations, we assumed that BC has a density of 2000 kg m^{-3} and follows a 200 201 logarithmic size distribution with an aerodynamic mean diameter of 0.25 µm and a logarithmic standard deviation of 0.3. Each computational particle released in FLEXPART 202 203 represents an aerosol population with a lognormal size distribution (see Stohl et al., 2005). This treatment of aerosol size distribution allows simulating several different types of 204 205 particles, each with its own size distribution. Removal processes acting differently for the different particle sizes will then affect specific particle sizes. Assumed aerodynamic mean 206 diameter and logarithmic standard deviation are used by FLEXPART's dry deposition 207 scheme, which is based on the resistance analogy 38 , and they are consistent with those used in 208 other transport models ^{18,39}. Below-cloud scavenging was determined based on the 209 precipitation rate taken from ECMWF. The in-cloud scavenging was based on cloud liquid 210 water and ice content, precipitation rate and cloud depth from ECMWF⁴⁰. The FLEXPART 211 user manual (available from http://www.flexpart.eu) provides more information on 212 FLEXPART's removal parameterizations. All FLEXPART results for the cruise can be 213 214 viewed interactively at the URL http://niflheim.nilu.no/NikolaosPY/RusArctExp 2015.py.

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3. RESULTS AND DISCUSSION

3.1 **Onboard EBC Measurements.** The EBC concentrations measured during the cruise 216 217 are shown in Figure 1a. At the beginning of the expedition (10 October 2015) when the ship was in or near the port of Arkhangelsk (White Sea), high values of EBC were measured 218 (hourly values up to 700 ng m⁻³) probably due to local pollution. Only after the ship passed 219 the industrial area of the Dvina river delta (10 October 2015 at 20:30), EBC dropped to below 220 100 ng m⁻³. In the open White Sea, EBC was 40 ng m⁻³, on average, but a small peak (~163 221 ng m⁻³) was observed near the Kola Peninsula in the morning of 11 October (06:30). In the 222 basin of the Barents Sea absorption was below the detection limit of the aethalometer, and 223 224 only in the Pechora Sea (West of Kara Strait) on 12 October (06:30) EBC concentrations rose above the minimum detection levels again, gradually increasing up to 153 ng m⁻³. In the Kara 225 Strait EBC was strongly enhanced (~220 ng m⁻³); concentrations kept increasing in the Kara 226 Sea up to a maximum of 360 ng m⁻³ (Figure 1a), in an area north of strong gas flaring 227 emissions (see Figure 1 of Stohl et al.²¹). Notice that at remote Arctic stations, measured EBC 228 concentrations are much lower, typically only around 10 ng m⁻³ at this time of the year ⁴¹, 229 which can be considered the typical Arctic background ⁴². Hence, EBC values observed in the 230 Barents Sea were relatively close to the background concentrations observed in other parts of 231 232 the Arctic, whereas in the Kara Sea EBC concentrations were strongly enhanced compared to

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this level. It is worth to note that the measured EBC concentrations are comparable to those
reported by Stohl et al.²¹ of about 200-400 ng m⁻³ during a ship cruise in the Kara Sea in
September 2011.

In the morning of 13 October (07:30), when the ship was in the Eastern Kara Sea, EBC dropped to 100 ng m⁻³, then varied between 50 and 220 ng m⁻³ until midnight of 14 October, before decreasing towards minimum detectable limits until archipelago Severnaya Zemlya. On 15 October (02:50) the ship moored in the Shokalsky's passage near station "Ice Base Cape Baranova" on the Bolshevik Island (Figure S 1) until 18 October, when the voyage back to Arkhangelsk started.

242 On the way back to Arkhangelsk, in the morning of 18 October we observed EBC concentrations reaching around 60 ng m⁻³ (Figure 1a). While these concentrations were lower 243 than those observed on the way to the Bolshevik Island, they are still much higher than the 244 Arctic background. From 19 October at 10:00 to 21 October at 22:00 the ship maneuvered in 245 the central part of the Kara Sea searching for mooring stations. At that time BC varied to 246 about 200 ng m⁻³. However, due to frequent changes of the ship's course, the ship's exhaust 247 248 might have been transported to the clean air site (see Figure S1) via complicated pathways. Therefore, enhanced EBC measurements during this period were excluded from further 249 analysis. On 21 October, when the vessel continued its voyage to Arkhangelsk, relatively high 250 251 EBC concentrations were measured, while on 22 October at 18:00 no absorption could be 252 measured. On 23 October, the ship passed through the Kara Strait recording EBC concentrations of up to 250 ng m⁻³. Measured EBC concentrations declined substantially in 253 the Barents Sea, until the ship reached the Kola Peninsula where a small peak was recorded 254 on 24 October at around 6:00. Then EBC rapidly increased along the Dvina River in the 255 White Sea with a maximum of about 1100 ng m⁻³ on 24 and 25 October 2015. When the ship 256 arrived at the port of Arkhangelsk, EBC concentrations of 1500 ng m⁻³ were measured. 257 258 Although we initially considered these high EBC concentrations close to the port of 259 Arkhangelsk as local pollution, in the next section we show that this was actually not the case.

260 **3.2** Analysis of BC sources observed during the expedition.

Figure 1b shows the modeled concentrations color-coded according to their age since emission in contrast to the measurements, while in Figure 1c the modeled concentrations are separated according to the different emission categories. It was already mentioned that the

ECLIPSE inventory includes anthropogenic and biomass burning emission sources adopted 264 from GFEDv3.1 ^{36,37}. Flaring emissions dominate the emissions from the energy (ENE), 265 sector south of the Barents and Kara Seas. Generally, the model captured periods with 266 267 enhanced concentrations (e.g., in the Kara Sea during both the outward and return trip) and 268 such with very low concentrations (e.g., in the Barents Sea) quite well. One exception is the first few hours of the cruise, when FLEXPART retroplumes showed that clean air masses 269 from the Arctic reached the vessel in the port of Arkhangelsk. It is, however, very likely that 270 the high measured EBC concentrations were caused exclusively by local pollution within the 271 272 port, which cannot be captured by FLEXPART.

273 In the morning of 12 October, FLEXPART strongly overestimated the measured BC 274 concentrations (shortly before the ship passed through the Kara Strait), then underestimated 275 them by about 50%, and finally captured them almost exactly in the Kara Sea (13 October). 276 As shown in Figure 1c, the modeled concentrations during this period had a large flaring 277 contribution (ENE in Figure 1c). The measurements during this period thus enable us 278 constraining the rather uncertain gas flaring emissions. Before the highest modeled BC peak 279 on 12 October, retroplumes arrived straight from the east, with very little influence from the 280 continent. At the time of the model peak, however, the retroplume encountered the northern 281 parts of a strong cyclone centered over the Urals during the previous days. As a consequence, 282 the retroplume turned direction over the Nenets and Komi regions and almost exactly where the ECLIPSE inventory places very high gas flaring emissions, resulting in very high values 283 284 of the footprint emission sensitivity (Figure 2a) and source contributions (Figure 1c). This 285 complex situation prevailed only for about 3 hours. After that, the retroplume circled the 286 whole cyclone and this situation prevailed constantly for more than a day and during the 287 entire passage of the Kara Sea (see Figure 2c and 3d). Based on the above analysis, it is likely 288 that the modeled BC peak on 12 October is a result of the model not capturing the complex meteorological situation accurately enough. Even a small shift in the location of where the 289 290 retroplume turned (Figure 2a) would have produced much smaller simulated BC 291 concentrations. When the meteorological situation was more stable, the model captured the 292 measured EBC concentrations rather well, especially on 13 October, when gas flaring emissions from the Yamal and Khanty-Mansiysk region contributed strongly. This suggests 293 294 that gas flaring emissions for this region in the ECLIPSE inventory are in the right order of 295 magnitude, perhaps with a slight tendency towards overestimation in the Nenets and Komi 296 regions.

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The very small EBC values in the Severnaya Zemlya archipelago were also well 297 298 captured by FLEXPART (Figure 1b). During this time, the retroplumes showed transport from the Arctic Ocean, with very little influence from land sources. Figure 3a and b depict 299 300 FLEXPART daily average emission sensitivities calculated when the vessel arrived to 301 Severnaya Zemlya (14 October 2015) and when it departed (18 October 2015). Winds shifted 302 on 18 October, with retroplumes arriving again first from southerly directions and thus 303 increasing the potential for BC uptake over the land. Indeed, both measured and modeled BC 304 concentrations increased again on 18 October.

305 On the way back, measured EBC concentrations in the Kara Sea were again captured 306 quite accurately by FLEXPART. On 19 to 20 October, BC originated mainly from the 307 Russian gas flaring sites of Yamal and Khanty-Mansiysk, confirming that these emissions 308 appear to be well captured by the ECLIPSE inventory. From the afternoon on 20 October, air 309 arrived straight from the west and was not influenced anymore by sources on the continent. This was also the case on 21 October when air came from the north. Measurements also 310 showed decreasing EBC concentrations from 20 to 21 October (from 135 ng m⁻³ on 20 311 October at 10:00 to near the detection limit on 21 October at 2:00). 312

On 22 October, as the ship approached the Kara Strait, air arrived from the southwest and gas flaring emissions from the Nenets and Komi regions were sampled again, similar to 12 October. This time, the model overestimated the measured EBC concentrations only slightly. Nevertheless, together with the results from the outward journey, this may suggest that flaring emissions in the Nenets and Komi regions are somewhat overestimated in the ECLIPSE inventory.

On 24 to 25 October, measured EBC values in the White Sea reached more than 1000 319 ng m⁻³ and FLEXPART simulated similarly high BC values. The retroplumes at this time 320 arrived from the southwest and brought polluted air masses mainly from Eastern Europe 321 322 (Figure 4a). An example of the source contributions for 25 October at 00:00 is shown in Figure 4b and 5c. At that time, the modeled concentration of BC was 1310.5 ng m⁻³, which is 323 324 close to the observed values in the range from 696 to 1501 ng m⁻³. About 10% (130.8 ng m⁻³) originated from fires over Ukraine (Figure 4c), whereas about 90% originated from 325 326 anthropogenic sources mainly in Central and Eastern Europe (Figure 4b). Excluding biomass burning, surface transportation contributed about 38%, residential and commercial 327

combustion sources up to 41%, gas flaring contributed about 8% and emissions from industrial combustion and processing between 1-2%.

330 Figure 5 depicts calculated normalized bias for the daily average measured EBC and modeled BC concentrations along the ship track in the White, Barents and Kara Seas. This 331 statistic expresses the difference (model-observed) over the observed values. It is a useful 332 333 indicator for assessing the models' performance because it avoids over-inflating the observed range of values, especially at low concentrations and it is used here to show the locations 334 335 where modeled concentrations over- or underestimated the observations. The model is least biased when the gas flaring sources contribute the most to surface concentrations of BC, 336 337 namely in the Pechora Sea (west of Kara Sea), in the Kara Strait and in the Kara Sea on the 338 way to the Bolshevik Island, as well as in the middle of the Kara Sea (point 3 in Figure S 1), 339 and close to the port of Arkhangelsk on the way back to Arkhangelsk. The extremely low concentrations calculated by the model in the beginning of the cruise in contrast to the high 340 EBC concentrations (\approx 700 ng m⁻³) led to negative biases near the port of Arkhangelsk and in 341 the industrial area of Dvina river delta. On the contrary, the lack of absorption in the 342 343 aethalometer near the Bolshevik Island from 15 to 17 October resulted in significant 344 overestimated predicted BC concentrations by the model and high positive biases (Figure 5).

The very good agreement ($R^2 = 0.76$) between modeled and measured concentrations 345 was confirmed by the root mean square error (RMSE). Whereas R^2 is a relative measure of 346 fit, RMSE is an absolute measure of fit. It can be interpreted as the standard deviation of the 347 unexplained variance; hence it is in the same units as the response variable. Lower values of 348 RMSE indicate better fit. RMSE is a good measure of how accurately the model predicts the 349 350 response, and is the most important criterion for fit if the main purpose of the model is prediction. The RMSE when including all data was estimated to be 230 ng m⁻³. This high 351 352 value is more or less expected here considering that the RMSE calculates the square error, hence it is very sensitive to larger errors. In the present case, if the points from the initial 353 354 period of the cruise (Arkhangelsk and Dvina river industrial area) that were subject to local pollution are excluded, the RMSE falls to 85 ng m⁻³, which is very low compared to the range 355 of values observed during the cruise $(0-1500 \text{ ng m}^{-3})$. 356

Overall, we found that the model had no systematic bias compared to the observations, which supports the validity of the ECLIPSE emission inventory for northern Russia. The good agreement especially in the region where flaring emissions are important suggests that flaring emissions are also captured quite well in this inventory. This is particularly true for the Yamal and Khanty-Mansiysk regions, whereas there may be some overestimation of flaring emissions in the Nenets and Komi regions. Local pollution cannot be captured neither by our model due to poor temporal and spatial resolution of the available operational wind fields, nor by the emission inventory used (available in 0.5° resolution). When local pollution was insignificant (e.g., in regions far from urban and industrial areas), emissions from residential and commercial combustion, as well as surface transportation were also captured well.

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368 ASSOCIATED CONTENT

369 Supporting Information

Figure S 1 shows the ship track of the research vessel "Akademik Treshnikov" in the Arctic Ocean and the main flaring facilities located in high latitudes. Figure S 2 depicts the measured meteorological conditions during cruise namely surface air temperature and pressure, and wind velocity and direction. Finally, Figure S 3 shows the quality of the EBC measurements (QA/QC) in terms of comparison of attenuation coefficients of the aethalometers used onboard (MSU) against EBC concentrations obtained with the AE33 aethalometer. This material is available free of charge via the Internet at <u>http://pubs.acs.org</u>.

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388 **REFERENCES**

389 390 391 392	(1)	Bond, T. C.; Doherty, S. J.; Fahey, D. W.; Forster, P. M.; Berntsen, T.; Deangelo, B. J.; Flanner, M. G.; Ghan, S.; Kärcher, B.; Koch, D.; et al. Bounding the role of black carbon in the climate system: A scientific assessment. <i>J. Geophys. Res. Atmos.</i> 2013 , <i>118</i> (11), 5380–5552.
393 394 395 396	(2)	IPCC. Climate Change 2013: The Physical Science Basis. Contribution to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change.; Stocker, T. F., Qin, D., Plattner, GK., Tignor, M. M. B., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., Midgley, P. M., Eds.; Cambridge University Press, 2013.
397 398 399 400	(3)	Shindell, D. T.; Chin, M.; Dentener, F.; Doherty, R. M.; Faluvegi, G.; Fiore, A. M.; Hess, P.; Koch, D. M.; MacKenzie, I. A.; Sanderson, M. G.; et al. A multi-model assessment of pollution transport to the Arctic. <i>Atmopsheric Chem. Phys.</i> 2008 , <i>8</i> , 5353–5372.
401 402 403	(4)	Bowerman, N. H. A.; Frame, D. J.; Huntingford, C.; Lowe, J. A.; Smith, S. M.; Allen, M. R. The role of short-lived climate pollutants in meeting temperature goals. <i>Nat. Clim. Chang.</i> 2013 , <i>3</i> (12), 1021–1024.
404 405 406 407	(5)	 Quinn, P. K.; Bates, T. S.; Baum, E.; Doubleday, N.; Fiore, A. M.; Flanner, M.; Fridlind, A.; Garrett, T. J.; Koch, D.; Menon, S.; et al. Short-lived pollutants in the Arctic: their climate impact and possible mitigation strategies. <i>Atmos. Chem. Phys.</i> 2008, <i>8</i>, 1723–1735.
408 409	(6)	Warren, S. G.; Wiscombe, W. J. A Model for the Spectral Albedo of Snow. II: Snow Containing Atmospheric Aerosols. <i>J. Atmos. Sci.</i> 1980 , <i>37</i> , 2734–2745.
410 411 412	(7)	Flanner, M. G.; Zender, C. S.; Randerson, J. T.; Rasch, P. J. Present-day climate forcing and response from black carbon in snow. <i>J. Geophys. Res. Atmos.</i> 2007 , <i>112</i> (11), 1–17.
413 414 415	(8)	Petzold, A.; Ogren, J. A.; Fiebig, M.; Laj, P.; Li, S. M.; Baltensperger, U.; Holzer- Popp, T.; Kinne, S.; Pappalardo, G.; Sugimoto, N.; et al. Recommendations for reporting black carbon measurements. <i>Atmos. Chem. Phys.</i> 2013 , <i>13</i> (16), 8365–8379.
416 417 418	(9)	Popovicheva, O. B. Combustion-derived carbonaceous aerosols (soot) in the atmosphere: Water interaction and climate effects. In <i>Aerosol Science and Technology</i> ; Agranovski, I., Ed.; Wiley - VCH Verlag GmbH & Co. KGaA, 2010; pp 127–157.

419 420 421	(10)	 Yun, Y.; Penner, J. E.; Popovicheva, O. The effects of hygroscopicity on ice nucleation of fossil fuel combustion aerosols in mixed-phase clouds. <i>Atmos. Chem. Phys.</i> 2013, <i>13</i> (8), 4339–4348.
422 423 424 425	(11)	Popovicheva, O.; Kistler, M.; Kireeva, E.; Persiantseva, N.; Timofeev, M.; Kopeikin, V.; Kasper-Giebl, A. Physicochemical characterization of smoke aerosol during large- scale wildfires: Extreme event of August 2010 in Moscow. <i>Atmos. Environ.</i> 2014 , <i>96</i> (August 2011), 405–414.
426 427 428 429	(12)	Diapouli, E.; Popovicheva, O.; Kistler, M.; Vratolis, S.; Persiantseva, N.; Timofeev, M.; Kasper-Giebl, A.; Eleftheriadis, K. Physicochemical characterization of aged biomass burning aerosol after long-range transport to Greece from large scale wildfires in Russia and surrounding regions, Summer 2010. <i>Atmos. Environ.</i> 2014 , <i>96</i> , 393–404.
430 431 432	(13)	Eleftheriadis, K.; Vratolis, S.; Nyeki, S. Aerosol black carbon in the European Arctic: Measurements at Zeppelin station, Ny-??lesund, Svalbard from 1998-2007. <i>Geophys.</i> <i>Res. Lett.</i> 2009 , <i>36</i> (2), 1–5.
433 434	(14)	AMAP. AMAP assessment 2015: Black carbon and ozone as Arctic climate forcers; Arctic Monitoring and Assessment Programme (AMAP), Oslo, Norway, 2015.
435 436	(15)	Wang, Q.; Chen, X. Nuclear accident like Fukushima unlikely in the rest of the world? <i>Environ. Sci. Technol.</i> 2011 , <i>45</i> (23), 9831–9832.
437 438 439 440	(16)	Stohl, A.; Berg, T.; Burkhart, J. F.; Fjæraa, a. M.; Forster, C.; Herber, A.; Hov, Ø.; Lunder, C.; McMillan, W. W.; Oltmans, S.; et al. Arctic smoke – record high air pollution levels in the European Arctic due to agricultural fires in Eastern Europe. <i>Atmos. Chem. Phys. Discuss.</i> 2006 , <i>6</i> (5), 9655–9722.
441 442 443 444	(17)	Stock, M.; Ritter, C.; Herber, A.; von Hoyningen-Huene, W.; Baibakov, K.; Gräser, J.; Orgis, T.; Treffeisen, R.; Zinoviev, N.; Makshtas, A.; et al. Springtime Arctic aerosol: Smoke versus haze, a case study for March 2008. <i>Atmos. Environ.</i> 2012 , <i>52</i> (March 2008), 48–55.
445 446 447 448	(18)	Evangeliou, N.; Balkanski, Y.; Hao, W. M.; Petkov, A.; Silverstein, R. P.; Corley, R.; Nordgren, B. L.; Urbanski, S. P.; Eckhardt, S.; Stohl, A.; et al. Wildfires in northern Eurasia affect the budget of black carbon in the Arctic-a 12-year retrospective synopsis (2002-2013). <i>Atmos. Chem. Phys.</i> 2016 , <i>16</i> (12), 7587–7604.

449 450 451 452	(19)	 AMAP. <i>The Impact of Black Carbon on Arctic Climate (2011)</i>; Berntsen, T., Burkhart, J. F., Christensen, J., Flanner, M., Kupiainen, K., Lihavainen, H., Shepherd, M., Shevchenko, V., Skov, H., Vestreng, V., Eds.; Arctic Monitoring and Assessment Programme (AMAP), Oslo, Norway, 2011; Vol. 4.
453 454 455	(20)	Cofala, J.; Amann, M.; Klimont, Z.; Kupiainen, K.; Höglund-Isaksson, L. Scenarios of global anthropogenic emissions of air pollutants and methane until 2030. <i>Atmos. Environ.</i> 2007 , <i>41</i> (38), 8486–8499.
456 457 458 459	(21)	Stohl, A.; Klimont, Z.; Eckhardt, S.; Kupiainen, K.; Shevchenko, V. P.; Kopeikin, V. M.; Novigatsky, A. N. Black carbon in the Arctic: The underestimated role of gas flaring and residential combustion emissions. <i>Atmos. Chem. Phys.</i> 2013 , <i>13</i> (17), 8833–8855.
460 461	(22)	Cheng, M. D. Geolocating Russian sources for Arctic black carbon. <i>Atmos. Environ.</i> 2014 , <i>92</i> , 398–410.
462 463 464	(23)	Uttal, T.; Makshtas, A.; Laurila, T. The Tiksi International Hydrometeorological Observatory - An Arctic Members Partnership. <i>Bull. World Meteorol. Organ.</i> 2013 , <i>62</i> (1), 22–26.
465 466 467	(24)	Winiger, P.; Andersson, A.; Eckhardt, S.; Stohl, A.; Semiletov, I. P.; Dudarev, O. V.; Charkin, A.; Shakhova, N.; Klimont, Z.; Heyes, C.; et al. Siberian Arctic black carbon sources constrained by model and observation. <i>Proc. Natl. Acad. Sci.</i> 2017 , 1–8.
468 469 470 471	(25)	Huang, K.; Fu, J. S.; Prikhodko, V. Y.; Storey, J. M.; Romanov, A.; Hodson, E. L.; Cresko, J.; Morozova, I.; Ignatieva, Y.; Cabaniss, J. Russian anthropogenic black carbon: Emission reconstruction and Arctic black carbon simulation. <i>J. Geophys. Res.</i> <i>Atmos.</i> 2015 , <i>120</i> (21), 11306–11333.
472 473	(26)	Huang, K.; Fu, J. S. Data Descriptor : A global gas fl aring black carbon emission rate dataset from 1994 to 2012. <i>Nature</i> 2016 , 1–11.
474 475 476 477 478	(27)	Panchenko, M. V.; Kozlov, V. S.; Pol'kin, V. V.; Golobokova, L. P.; Pogodaeva, T. V.; Khodzher, T. V.; Lisitzin, A. P.; Shevchenko, V. P. Investigations of microphysical and chemical composition of aerosol in near-water layer of the atmosphere over the White Sea. <i>Proc. SPIE 6522, Thirteen. Jt. Int. Symp. Atmos. Ocean Opt. Atmos. Phys.</i> 2006 , <i>66221A</i> .

479 480 481	(28)	Kopeikin, V. M.; Repina, I. A.; Grechko, E. I.; Ogorodnikov, B. I. Measurements of soot aerosol content in the near-water atmospheric layer in the southern and northern hemispheres. <i>Atmos. Ocean. Opt.</i> 2010 , <i>23</i> (6), 500–507.
482 483 484 485 486	(29)	Sakerin, S. M.; Bobrikov, A. A.; Bukin, O. A.; Golobokova, L. P.; Pol'Kin, V. V.; Pol'Kin, V. V.; Shmirko, K. A.; Kabanov, D. M.; Khodzher, T. V.; Onischuk, N. A.; et al. On measurements of aerosol-gas composition of the atmosphere during two expeditions in 2013 along the Northern Sea Route. <i>Atmos. Chem. Phys.</i> 2015 , <i>15</i> (21), 12413–12443.
487 488	(30)	Hansen, A. D. A.; Rosen, H. Horizontal inhomogeneities in the particulate carbon component of the Arctic haze. <i>Atmos. Environ.</i> 1985 , <i>19</i> (12), 2175–2180.
489 490 491 492	(31)	Drinovec, L.; Močnik, G.; Zotter, P.; Prévôt, A. S. H.; Ruckstuhl, C.; Coz, E.; Rupakheti, M.; Sciare, J.; Müller, T.; Wiedensohler, A.; et al. The "dual-spot" Aethalometer: An improved measurement of aerosol black carbon with real-time loading compensation. <i>Atmos. Meas. Tech.</i> 2015 , <i>8</i> (5), 1965–1979.
493 494 495 496	(32)	Zanatta, M.; Gysel, M.; Bukowiecki, N.; M??ller, T.; Weingartner, E.; Areskoug, H.; Fiebig, M.; Yttri, K. E.; Mihalopoulos, N.; Kouvarakis, G.; et al. A European aerosol phenomenology-5: Climatology of black carbon optical properties at 9 regional background sites across Europe. <i>Atmos. Environ.</i> 2016 , <i>145</i> , 346–364.
497 498 499	(33)	Stohl, A.; Hittenberger, M.; Wotawa, G. Validation of the lagrangian particle dispersion model FLEXPART against large-scale tracer experiment data. <i>Atmos. Environ.</i> 1998 , <i>32</i> (24), 4245–4264.
500 501 502	(34)	Stohl, a.; Forster, C.; Frank, A.; Seibert, P.; Wotawa, G. Technical note: The Lagrangian particle dispersion model FLEXPART version 6.2. <i>Atmos. Chem. Phys. Discuss.</i> 2005 , <i>5</i> (4), 4739–4799.
503 504 505 506	(35)	Stohl, A.; Forster, C.; Eckhardt, S.; Spichtinger, N.; Huntrieser, H.; Heland, J.; Schlager, H.; Wilhelm, S.; Arnold, F.; Cooper, O. A backward modeling study of intercontinental pollution transport using aircraft measurements. <i>J. Geophys. Res.</i> <i>Atmos.</i> 2003 , <i>108</i> (D12), 4370.
507 508	(36)	Stohl, A.; Aamaas, B.; Amann, M.; Baker, L. H.; Bellouin, N.; Berntsen, T. K.; Boucher, O.; Cherian, R.; Collins, W.; Daskalakis, N.; et al. Evaluating the climate and

509 510		air quality impacts of short-lived pollutants. Atmos. Chem. Phys. 2015, 15 (18), 10529–10566.
511 512 513	(37)	van der Werf, G. R.; Randerson, J. T.; Giglio, L.; Collatz, G. J.; Kasibhatla, P. S.; Arellano, A. F., J. Interannual variability in global biomass burning emissions from 1997 to 2004. <i>Atmos. Chem. Phys.</i> 2006 , <i>6</i> (11), 3423–3441.
514 515	(38)	Slinn, W. G. N. Predictions for particle deposition to vegetative canopies. <i>Atmos. Environ.</i> 1982 , <i>16</i> , 1785–1794.
516 517 518	(39)	Shiraiwa, M.; Kondo, Y.; Moteki, N.; Takegawa, N.; Sahu, L. K.; Takami, A.; Hatakeyama, S.; Yonemura, S.; Blake, D. R. Radiative impact of mixing state of black carbon aerosol in Asian outflow. <i>J. Geophys. Res. Atmos.</i> 2008 , <i>113</i> (24), 1–13.
519 520 521	(40)	Grythe, H.; Kristiansen, N. I.; Groot Zwaaftink, C. D.; Eckhardt, S.; Ström, J.; Tunved, P.; Krejci, R.; Stohl, A. A new aerosol wet removal scheme for the Lagrangian particle model FLEXPART. <i>Geosci. Model Dev. Discuss.</i> 2016 , No. October, 1–34.
522 523 524 525 526	(41)	Eckhardt, S.; Quennehen, B.; Olivié, D. J. L.; Berntsen, T. K.; Cherian, R.; Christensen, J. H.; Collins, W.; Crepinsek, S.; Daskalakis, N.; Flanner, M.; et al. Current model capabilities for simulating black carbon and sulfate concentrations in the Arctic atmosphere: A multi-model evaluation using a comprehensive measurement data set. <i>Atmos. Chem. Phys.</i> 2015 , <i>15</i> (16), 9413–9433.
527 528 529 530	(42)	Stone, R. S.; Sharma, S.; Herber, A.; Eleftheriadis, K.; Nelson, D. W. A characterization of Arctic aerosols on the basis of aerosol optical depth and black carbon measurements. <i>Elem. Sci. Anthr.</i> 2014 , <i>2</i> , 1–22.

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531 FIGURE CAPTIONS FOR MANUSCRIPT

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533 Figure 1. (a) Time series of equivalent black carbon (EBC) mass concentrations during the expedition cruise. Numbers and letters in red brackets show geographical names during the 534 535 cruise according to Figure S 1. (b) Age spectra of modeled BC (colors) from all possible 536 sources showing the contribution of emissions each day back in time to the surface 537 concentration of BC. Hourly means of measured BC concentrations are shown as a black line. (c) Contribution from different emission source types to the BC surface concentrations. The 538 539 emission sources of biomass burning (BB), waste burning (WST), industrial combustion and 540 processing (IND), surface transportation (TRA), power plants, energy conversion, and 541 extraction (ENE), and residential and commercial (DOM) have been adopted from GFEDv3.1 and ECLIPSE inventories ^{36,37}. Notice the different scale used in all three panels from 24 to 542 543 25 October, when measured and modeled concentrations were much higher than for the rest 544 of the cruise. Flaring emissions are included in the energy sector (ENE).

545 Figure 2. (a) Footprint emission sensitivity and (b) contribution from anthropogenic sources 546 to surface BC concentrations on 12 October 2015 at 05:00. (c) Footprint emission sensitivity 547 and (d) contribution from anthropogenic sources to surface BC concentrations on 13 October 2015 at 11:44. Values written in black report the simulated concentration of BC at the 548 549 receptor (ship) for the same time period from all anthropogenic sources, while colored ones denote the continental contribution from anthropogenic sources. Magenta shows contribution 550 from South America, orange from Europe, yellow from Australia, green from North America, 551 552 cyan from Africa and blue from Asia.

Figure 3. Daily average footprint emission sensitivities when the vessel (a) arrived (14
October 2015) and (b) departed (18 October 2015) from the "Ice Base Cape Baranova"
station.

Figure 4. (a) Footprint emission sensitivity when the ship had passed the Dvina River and before arrival to the port of Arkhangelsk (25 October 2015 00:00). (b) Contribution from anthropogenic sources and (c) biomass burning to the simulated surface concentration of BC at the same date and time. Black values show the concentration of BC at the receptor (ship) for the time period from all anthropogenic and biomass burning sources. Colored values denote continental contribution from anthropogenic sources; magenta show contribution from

- 562 South America, orange from Europe, yellow from Australia, green from North America, cyan
- 563 from Africa and blue from Asia.
- 564 Figure 5. Distribution of normalized bias, i.e., (model-observed)/observed, for the measured
- 565 EBC and the BC concentrations predicted by FLEXPART. The biases were calculated for the
- 566 daily average concentrations and for the ship location at midnight of each day (00:00).



(a) Time series of equivalent black carbon (EBC) mass concentrations during the expedition cruise. Numbers and letters in red brackets show geographical names during the cruise according to Figure S 1. (b) Age spectra of modeled BC (colors) from all possible sources showing the contribution of emissions each day back in time to the surface concentration of BC. Hourly means of measured BC concentrations are shown as a black line. (c) Contribution from different emission source types to the BC surface concentrations. The emission sources of biomass burning (BB), waste burning (WST), industrial combustion and processing (IND), surface transportation (TRA), power plants, energy conversion, and extraction (ENE), and residential and commercial (DOM) have been adopted from GFEDv3.1 and ECLIPSE inventories 36,37. Notice the different scale used in all three panels from 24 to 25 October, when measured and modeled concentrations were much higher than for the rest of the cruise. Flaring emissions are included in the energy sector (ENE). Figure 1

261x182mm (300 x 300 DPI)



(a) Footprint emission sensitivity and (b) contribution from anthropogenic sources to surface BC concentrations on 12 October 2015 at 05:00.
 (c) Footprint emission sensitivity and (d) contribution from anthropogenic sources to surface BC concentrations on 13 October 2015 at 11:44. Values written in black report the simulated concentration of BC at the receptor (ship) for the same time period from all anthropogenic sources, while colored ones denote the continental contribution from anthropogenic sources. Magenta shows contribution from South America, orange from Europe, yellow from Australia, green from North America, cyan from Africa and blue from Asia.

Figure 2 275x214mm (300 x 300 DPI)



Daily average footprint emission sensitivities when the vessel (a) arrived (14 October 2015) and (b) departed (18 October 2015) from the "Ice Base Cape Baranova" station. Figure 3 320x500mm (300 x 300 DPI)



Distribution of normalized bias, i.e., (model-observed)/observed, for the measured EBC and the BC concentrations predicted by FLEXPART. The biases were calculated for the daily average concentrations and for the ship location at midnight of each day (00:00). Figure 5

155x117mm (300 x 300 DPI)





(a) Footprint emission sensitivity when the ship had passed the Dvina River and before arrival to the port of Arkhangelsk (25 October 2015 00:00). (b) Contribution from anthropogenic sources and (c) biomass burning to the simulated surface concentration of BC at the same date and time. Black values show the concentration of BC at the receptor (ship) for the time period from all anthropogenic and biomass burning sources. Colored values denote continental contribution from anthropogenic sources; magenta show contribution from South America, orange from Europe, yellow from Australia, green from North America, cyan from Africa and blue from Asia.

> Figure 4 248x180mm (300 x 300 DPI)



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