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MODELLING OF PHOTOCHEMICAL OXIDANTS
AND PRECURSORS IN EUROPE
28 MAY - 3 JUNE 1982

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SUMMARY

Model calculations of the long-range transport of photochemical oxidants to 14 rural sites in Europe (one in Austria, three in FRG, two in The Netherlands, one in France, four in Scandinavia and three in the UK) have been carried out for the time period 28 May - 3 June 1982. A lagrangian long-range transport model with the CBM-X chemistry has been applied on the EMEP-grid. The calculations have been compared with hourly measured ozone concentrations, and fair agreement found for the Scandinavian sites and some of the sites in the UK, FRG and The Netherlands.

Through sensitivity calculations it is shown that atmospheric temperature, mixing height, cloud cover and dry deposition velocities all have a significant influence on the calculated ozone concentration.

The choice of initial concentration of the precursors is also important for the calculated distribution of ozone. To be comparable with the NO_x-emissions estimated for the PHOXA area for the last days of May 1982, the annual average daily emissions estimated for the EMEP grid cells falling inside the PHOXA area, had to be increased by 40%. The annual average daily VOC emissions for EMEP were comparable to the values estimated in PHOXA. With the assumption that the annual average daily emissions estimated for all of the EMEP grid for 1982 was adjusted upwards by 40%, it was shown that a reduction of the NO_x-emissions uniformly over the grid by 25% or more leads to an increase in calculated O₃ at all sites, while a reduction in VOC-emissions by 25% or more is calculated to reduce ozone. A combined reduction or increase in both NO_x and VOC emissions by 25%, has little influence on ozone at the 14 receptor sites.

If the NO_x emissions were reduced to the EMEP emissions estimated for that part of 1982 (70% of the annual average), it was found that both NO_x- and VOC-emission control, separately or together, efficiently reduced ozone; VOC-emission control being somewhat more efficient than NO_x-emission control.

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MODELLING OF PHOTOCHEMICAL OXIDANTS AND PRECURSORS
IN EUROPE 28 MAY - 3 JUNE, 1982

1 INTRODUCTION

1.1 DESCRIPTION OF SUBPROJECTS WITH TNO

During 1985, NILU was awarded two subcontracts with TNO, Apeldoorn, The Netherlands, on projects that TNO was to carry out for the Commission of the European Communities (CEC). The first subcontract covered the participation as a cooperating institution in the "study on photochemical oxidants and precursors-phase IV-long range transport phenomena". The work-load in this project, including NILU's part, was described in a technical annex distributed in January 1985. In short, there were four stages of progression:

First stage should involve episode selection (decided later to be May 29 - June 2 1982);

The next stage model calculations of the selected episode;

Stage 3 validation of the model results and calculation of the effect on ozone of reduction of emission (including proposed emission reductions for large combustion installations within CEC and proposed emission reductions for motor vehicles);

Stage 4: Assessment of the findings.

The second subcontract covered the participation of NILU as a cooperating institution in the study "photochemical oxidants and precursors - phase V". The work-load in this project, including NILU's part, was described in a technical annex distributed in November 1985. The objective of the project was to carry out a sensitivity analysis with the PHOXA RTM III-model and the Norwegian trajectory model, and there were three stages in the project:

In the first stage, a list of input parameters most likely to have the largest influence on the model results should be made;

In the second stage, sensitivity calculations should be carried out by both models based on the list developed in stage 1;

Third stage, assessment of findings with emphasis on O₃ and NO₂.

1.2 OUTLINE OF REPORT

Since both subcontracts with TNO cover model calculations of the photochemical episode 29 May - 2 June 1982, the results of the calculations carried out at NILU are reported together for both subcontracts.

Model calculations using the Norwegian lagrangian long-range transport model with CMB-X chemistry (Whitten et al., 1984) for the time period 28 May 1982, 1200 GMT, to 3 June 1982, 1200 GMT, have been carried out for 14 receptor points within the PHOXA grid area, see Table 1 for a list of the sites with geographical coordinates (latitude-longitude) and EMEP-coordinates, see Figure 1 for a map of the EMEP grid. Calculations have been carried out every 6 h GMT, i.e. at arrival times 0000, 0600, 1200 and 1800 GMT at each site.

More than 20 different sensitivity calculations have been carried out, 14 of them with different emission reductions of NO_x and VOC, the rest changing different aspects of the physical and chemical description of the formation of photochemical oxidants in Europe (changing the chemical scheme, temperature, mixing height, humidity, photolysis rate coefficients, deposition velocities and initial concentrations for the calculation).

In the report is given a brief outline of the Norwegian long-range transport model for photochemical oxidants and a description is given of the input data (emissions, meteorology, ground removal).

Following the model description, the results of the calculations are discussed, starting with the model validation where calculated and measured ozone concentrations for the 14 receptor points are discussed. Afterwards the results of the sensitivity calculations including the emission reductions, are discussed.

Table 1: Receptor points in the model calculation

Site No.	Site	Country	Latitude	Longitude	EMEP-coordinates	
			^o N	(^o E > 0 ^o W < 0)	x	y
1	Illmitz	A	47.77	16.77	26.01	16.83
2	Langenbrügge	FRG	52.80	10.75	21.10	17.42
3	Schauinsland	FRG	47.92	7.90	22.55	13.61
4	Deuselbach	FRG	49.77	7.05	21.88	14.46
5	Risø	DK	55.00	11.00	20.04	18.73
6	Rørvik	S	57.42	11.93	19.07	20.32
7	Langesund	N	59.00	9.75	17.63	20.60
8	Jeløy	N	59.40	10.75	17.72	21.08
9	Sappermeer	NL	53.10	6.47	19.45	16.30
10	Waarde	NL	51.25	4.05	19.40	14.47
11	Colomiers	F	43.37	1.20	21.70	8.42
12	Bottesford	UK	52.56	-0.48	17.04	14.11
13	Sibton	UK	52.50	1.50	17.85	14.57
14	Stodday	UK	54.10	3.20	17.80	16.02

2 MODEL DESCRIPTION

2.1 PREVIOUS APPLICATIONS

The Norwegian lagrangian long-range transport model with atmospheric boundary layer chemistry was described by Eliassen et al., (1982a).

In the first application of the model, ozone formation and transport to southern Norway and south-eastern Sweden during the time period 6-14 April 1979 was studied (Eliassen et al., 1982a). The air flow was primarily over eastern Europe towards southern Scandinavia.

In the second application of the model, the formation of oxidants during transport to southern Scandinavia during the time period 26 August to 14 September 1980 was studied (19 days). Ozone and PAN measurements at a number of sites in Scandinavia were used in the model validation. The flow direction was predominantly from the south-west and south during the time period (Hov et al., 1985).

2.2 METEOROLOGICAL MODEL

The model has been described in some detail previously (Eliassen et al., 1982a, Eliassen et al., 1982b, Hov et al., 1984). The pollutants are assumed to be completely vertically mixed throughout the boundary layer which has a variable depth along the 96 h long 850 mb trajectories. No mass transport takes place through the top of the well-mixed layer. Lateral diffusion is not treated explicitly, but the emission data are given in a 150 km grid where finer details than 150 km in the concentration fields are smoothed out.

In episode studies with short sampling times (like one hour), the rate of horizontal spread of instantaneous releases of pollutants may be an important parameter which should be considered (Eliassen, 1984). For a sampling time of many hours, like 6 h or more, the instantaneous diffusion of pollutant releases is dominated by the diffusion due to sampling time ("synoptic swinging of the trajectories", Smith, 1979). A sampling time of 24 h is used for sulphur species in EMEP, in which case the synoptic swinging of trajectories is the dominating factor for plume spread (Eliassen, 1984).

During transport, pollutants are emitted into the air parcel according to the emission maps for NO_x and VOC. Instantaneous concentrations are predicted upon arrival of a trajectory. The horizontal resolution of the concentration fields is determined by the choice of emission grid and density of trajectory arrival points. The combined effects of vertical wind shear and diffusion due to heat exchange is difficult to handle in lagrangian models. Trajectory models are simple numerically, however, since the integration is reduced to an ordinary time-integration along certain selected trajectories.

Trajectory positions are calculated every 2 h, as described in Petterssen (1956), based on wind observations at the 850 mb level at 0000, 0600, 1200 and 1800 GMT. The observed wind data are analysed objectively in the EMEP grid, cpr. Figure 1. In regions where wind observations are scarce, such as over sea, the wind analysis is heavily influenced by the quasi-geostrophic balanced wind predicted by the Norwegian Meteorological Institute as part of its weather prediction routine.

Alternative trajectories for transport at the 925 mb level rather than 850 mb, can be calculated by backing the analysed 850 mb wind by e.g. 10^0 and reducing it to 90% or so. Radiosonde observations close to the trajectory can give an indication as to the turning and change in speed of the wind with height. The mixing height used represents a material surface below which both old and new pollutants are mixed. The 1200 GMT mixing height is chosen. The basic data for the mixing height analysis are taken from radiosonde data (about 120 radiosonde reports are available within the grid). The estimated mixing heights are objectively analysed to produce grid values at 1200 GMT every day. At intermediate times it is assumed that each trajectory conserves its mixing height.

Objective analysis of temperature, relative humidity and absolute humidity are carried out at 0000 and 1200 GMT in the 150 km grid, as vertical averages between the surface and the 850 mb level. The temperature is used to evaluate temperature-dependent reaction rate coefficients. The relative humidity is used as a rough indication of cloud cover, which influences the photodissociation rates (see Table 2).

Table 2: Parameterization of cloud cover using the relative humidity.

Relative humidity	Cloud cover	"Effective" albedo
> 85%	1.0	0.6
75-85%	0.5	0.3
< 75%	0.0	0.0

When the relative humidity exceeds 90%, precipitation is assumed, and a wet deposition rate coefficient of $1 \times 10^{-4} \text{ s}^{-1}$ is applied to the concentrations of H_2SO_4 , HNO_3 , H_2O_2 and $\text{CH}_3\text{O}_2\text{H}$. For lower relative humidities than 90%, a first order wet deposition rate coefficient of $5 \times 10^{-6} \text{ s}^{-1}$ is applied. The individual trajectories are assigned mean values of temperature and absolute humidity at 0000 and 1200 GMT. The temperature is estimated by linear interpolation and the absolute humidity is conserved at intermediate positions.

2.3 DRY DEPOSITION

Dry deposition velocities appropriate for 1 m height are given in Table 3.

Table 3: Dry deposition velocities appropriate for 1 m above the ground.

Component	Deposition velocity (cm/s)	Comments
O ₃	0.5	daytime over land surfaces
O ₃	0.05	nighttime over land
O ₃	0.0	sea surfaces
NO ₂	0.5	daytime over land
HNO ₃	1.0	determined by aerodynamic resistance

To arrive at a model where average boundary layer concentrations are calculated rather than the concentration at 1 m, the deposition velocities given in Table 3 for O₃, NO₂ and PAN were simply reduced by 50%. Detailed calculations for June 1985 using meteorological data from the Numerical Weather Prediction Model at The Norwegian Meteorological Institute for surface pressure, surface stress, sensible heat flux density and temperature at 2 m height together with data for the surface roughness length and Businger's equations which relate the deposition velocity at the top of the surface layer (50 m height) to the deposition velocity at 1 m above the ground, show that the deposition velocity for SO₂ at 50 m typically was 50-75% of the value at 1 m (Hov et al., 1987). A reduction by a factor of 2 therefore means that the efficiency of the ground removal processes in the model perhaps are underestimated.

2.4 PHOTOCHEMICAL MODEL

The carbon bond mechanism denoted CBM-X was used in the model. It is described in detail by Whitten et al. (1984). This scheme describes the formation and decomposition of 63 chemical species through 146 chemical reactions. There are 10 classes of hydrocarbons (UNR, ETH, OLE, PAR, TOL, XYL, FORM, ALD2, KET and ACET). The photolysis rate coefficients for NO_2 , $\text{O}_3 \rightarrow \text{O}(^1\text{D})$, HCHO, ALD2(= CH_3CHO) are calculated explicitly through the integration over the absorption spectrum of the product of the wavelength dependent quantum yield, the absorption cross section and the solar flux at a given time and location. The photolysis rate coefficients of the other species which are photolysed, are calculated as a fixed fraction of the photolysis rate coefficient for NO_2 . The CBM-X mechanism is documented in Appendix 1.

An updated version of the surrogate mechanism used in the calculations with the Norwegian long-range transport model (Eliassen et al., 1982a; Hov et al., 1984) was also used for comparison with the results using the CBM-X mechanism. Of the hydrocarbon emissions, 30% by volume (on a compound basis) are represented as C_2H_6 , 20% as $n\text{C}_4\text{H}_{10}$, 20% as C_2H_4 , 10% as C_3H_6 and 20% as m-xylene. The chemical scheme and the representation of the hydrocarbon emissions are discussed in more detail by Eliassen et al. (1982a) and Hov et al. (1984). It consists of about 100 chemical reactions including photochemical reactions, and 40 different species.

Dissociation rate coefficients are calculated for every 5° latitude and every 15 min of the day. The total vertically integrated atmospheric ozone column is adjusted to correspond to the season and latitude. Points along a given trajectory are allocated dissociation rate coefficients through interpolation in time and space to the appropriate latitude and local time.

In the RTM III-model calculations, a condensed version of CBM-X is used (CBM-IV). CBM-IV, has 70 chemical reactions, 24 chemical species (excluding SO_2 and sulphate), and 9 hydrocarbon classes. It is described by Whitten and Gery (1985). CBM-IV has only fairly recently

become available for use within the PHOXA-project outside of the RTM III-model. A comparison of the results of 5 days of calculations in a continental photochemical box model using the surrogate mechanism in the Norwegian long-range transport model, CBM-X and other chemical schemes, was reported by Hov et al. (1986). It turned out that the CBM-X chemistry gave somewhat lower ozone concentrations than the surrogate mechanism, a result which is confirmed in the present report.

2.5 INITIAL CONDITIONS

The initial concentrations assigned at the starting point of the 96 h long trajectories can be important for the development along the trajectory. Ground removal is the ultimate removal mechanism for ozone, and in cases with low deposition, the lifetime of ozone is much longer than four days.

In such situations four days' trajectories may not be sufficient to trace the history of an air mass. If the weather is fair at the starting point, the air masses arriving there may have accumulated photochemically active pollution for a number of days.

The integration is started with a set of concentrations corresponding to a slightly polluted atmosphere, with the removal processes in equilibrium with NO_x and NMHC emissions at 10% of the average emissions for Western Europe. The initial concentrations of the most important species are listed in Table 4.

Table 4: Initial concentrations (ppbv).

Specie	Concentration	Specie	Concentration
NO	0.001	O ₃	32
NO ₂	0.3	HNO ₃	1.0
NMHC (C)	2.2	PAN	0.04

2.6 EMISSIONS

Much work is being done to improve European emission inventories for SO_2 , NO_x , VOC and also NH_3 . Within the PHOXA-project, annual and episode specific emissions for SO_2 , NO_x , CO and 10 hydrocarbon classes according to the requirement of CBM-X, have been established. The PHOXA-grid and the extension to be used in the RTM III-calculation for OECD are shown in Figure 2. Episode specific emissions with hourly values of NO, NO_2 , SO_2 , CO and 10 hydrocarbon classes for 31 May and 1-2 June 1982 for the PHOXA-grid, were made available to NILU by TNO.

It is apparent from Figure 2 that it is not straightforward to transfer PHOXA-grid emissions to the EMEP-grid. This not only applies to the EMEP grid squares fully outside of the PHOXA-grid, but also to the EMEP-grid squares only partially covered by PHOXA-grid elements.

For the EMEP-grid, unofficial inventories of NO_x and VOC-emissions exist, starting with an inventory which was thought to be representative for about 1980 (Eliassen et al., 1982a). As information from the EMEP-countries on national NO_x -emissions is coming in to EMEP MSC-W at The Norwegian Meteorological Institute and to EMEP CCC (at NILU), improved and updated NO_x -inventories for the EMEP-grid are established.

During the last 1-2 years activity has been organized also through OECD to collect emission data for the OECD countries in Europe in a grid which coincides with the EMEP-grid, but with grid elements of size $50 \times 50 \text{ km}^2$ rather than $150 \times 150 \text{ km}^2$ at 60°N latitude. The collection of emission data within OECD and PHOXA has improved the data availability in western Europe, but the very important assessment of emissions in East Europe and the western part of USSR, has to be done through EMEP.

For the part of the PHOXA-grid which fully covers EMEP-grid squares (i.e. excluding the parts of the PHOXA-grid which only partially cover EMEP grid cells), the contents of the PHOXA episode specific emission data file for 31 May - 2 June 1982, are summarized in Table 5.

Table 5: Total emissions in the PHOXA episode specific emission file for the grid cells which fully covers EMEP grid cells, translated into annual figures.

Date	NOx (kt(NO ₂)/y)	VOC (kt/y)
31 May 1982	14326	9580
1 June 1982	14173	10783
2 June 1982	14170	11747

For the emission inventory published by Eliassen et al. (1982a), valid for about 1980, the corresponding figures to Table 5 for NOx-emissions was 10113 kt(NO₂)/y and 10793 kt/y for VOC. It can be seen that the average VOC-emissions were similar, while the PHOXA-specific NOx-emissions were about 40% higher than the estimate for the EMEP-grid. However, the VOC-emissions for the PHOXA-grid contain a very significant fraction (about 2/3) thought to be natural hydrocarbons (terpenes and isoprene) for this episode where the temperatures were high over large parts of Europe. The VOC-emissions estimated for the EMEP-grid were anthropogenic only, as an annual average. In CBM-X, natural VOC's are treated as a mixture of paraffinic and olefinic bond type molecules.

To arrive at an emission inventory for the EMEP-grid for May-June 1982, it was decided to keep the grid distribution for NOx and VOC as estimated for 1980, but increasing the NOx-values for all of the EMEP grid by 40% to get agreement with the PHOXA-area estimate. Furthermore, the average distribution on hydrocarbon classes as found from the PHOXA episode specific emission file, was applied, see Table 6. The original split into hydrocarbons in the surrogate mechanism was retained, however (see section 2.4).

In this way some of the main features from the PHOXA episode specific emission file are retained in the calculation reported here.

Table 6: Distribution of VOC-emissions on hydrocarbon classes, on a mass basis.

Hydrocarbon class	Percent of total VOC emission on a mass basis
OLE	11
PAR	66
TOL	3
XYL	4
FORM	1
ALD2	7
KET	1
ACET	1
ETH	2
UNR	4
CO	350

2.7 MATHEMATICAL FORMULATION

The mass conservation equation determining the concentration c_i of species i can be written as

$$\frac{Dc_i}{dt} = -\left(\frac{v_d}{h} + k_w\right) c_i + \frac{E_i}{h} + S_i$$

The notation is

D/dt Lagrangian (total) time derivative along a trajectory

$v_d(x,y,t)$ dry deposition

$h(x,y,t)$ mixing height

$k_w(x,y,t)$ wet deposition rate

$E_i(x,y)$ direct emission of pollutant

S_i chemical sources or sinks.

In the integration procedure the appropriate back trajectories are first calculated from the analysed wind fields. Then the quantities v_d , h , k_w , etc. originally given as Eulerian fields, are converted into lagrangian information, i.e. as a function of transport time along the trajectories. These operations transform the mass conservation equation into an ordinary differential equation in time. Lastly, this equation is integrated to obtain calculated instantaneous concentrations at the receptor points.

The integration of the mass conservation equation is done with a quasi-steady-state approximation method (QSSA). This method is explicit and applies a fixed time step. The upper limit for the computational error is estimated to be 5%.

3 MODEL CALCULATIONS 28 MAY - 3 JUNE 1982

3.1 EPISODE DESCRIPTION

There was a high pressure system located over north Europe with its center over Denmark on 30 May 1982, moving eastward and with its center over East Europe on 2 June. The wind speeds were low, and the maximum hourly ozone concentration recorded was about 160 ppb, in the Netherlands on 1 June.

In Figures 3a-g is shown the 96 h, 860 mb, back trajectories at 1200 GMT to the 14 receptor points described in Table 1, for each of the days 28 May - 3 June 1982. On 28 May, the transport direction was from the southwest for the sites in Scandinavia, the northern part of FRG (Deuselbach and Langenbrügge) and the Netherlands (Sappermeer), southerly for Waarde, Colomiers and Schauinsland, and easterly for Illmitz. On 29 May the transport direction was from the west-southwest at most sites, on the 30 May the transport direction was more from the south for the Scandinavian sites than for the rest, except Colomiers, while for 31 May - 3 June the winds were low and the transport direction variable and the transport distance over 96 h fairly modest.

The mixing height field for 1200 GMT each day 28 May - 3 June 1982 is shown in Figures 4a-g. The field is calculated by an objective analysis of the mixing height as measured by about 120 radiosondes within the grid. In particular over oceans the results is quite much influenced by the initial guess (1000 m) and in areas with very few soundings, one sounding heavily influences the mixing height field over several grid lengths in all directions. When looking at Figures 4a-4g together with the map in Figure 1, it can be seen that in general the mixing height has a maximum over continental Europe and the Soviet Union exceeding 2000 m in the beginning of the period and exceeding

about 1500 m later in the period in western Europe while still higher than 2000 m over USSR. Over the Atlantic and in the Mediterranean the mixing height in general is lower than over the continents, perhaps by as much as 500-1000 m. There are many exceptions and in some cases one or a few radiosonde observations with high mixing heights give rise to values over 2000 m also over oceanic and coastal areas.

In Figure 5 an example is given of the relative humidity field. The values are low over continental Europe in the high pressure system, and higher over oceanic regions. The relative humidity in this case is derived from the radiosonde data, where temperature and dew point temperature at constant pressure surface and at significant levels, are recorded. From these data relative humidity is derived and gridded values obtained through objective analysis.

3.2 CALCULATIONS AND MEASUREMENTS

Hourly measurements of ozone were available for 13 of the 14 sites in Table 1 for the period 28 May - 3 June 1982. The measurements are taken very near the ground surface, usually only one or a few metres above the ground. This means that the measured concentrations usually are significantly reduced at night through ground removal below the nocturnal inversion and by local emissions of NO_x becoming trapped in the shallow nocturnal mixed layer. On the other hand, in the model a concentration representative of a layer with height comparable to the noon mixing height the day before, is calculated at night. This concentration is only weakly influenced by ground removal at night, and therefore the calculated diurnal variation of O₃ is usually smaller than the measured. It should be kept in mind that for measured and calculated ozone concentrations, only the day time values when the atmospheric boundary layer is well mixed, are really comparable.

In Figures 6a-6n the measured and calculated O₃-concentrations for the 14 sites are shown. In the case of Langenbrügge, measurements of ozone were not available.

In Figure 6a is shown the results for Illmitz, which is located in a rural area in eastern Austria at 117 m.a.s.l. Illmitz has a record of

measuring high ozone concentrations, as can be seen for April-September 1985 in Table 7 where ozone measurements from 24 European rural sites are presented (Grennfelt et al., 1987). It can be seen that at Illmitz, more than half of the hourly values exceeded 60 ppb, while at the station with the second highest number of measurements exceeding 60 ppb, Schauinsland, about 1/4 of the values exceeded 60 ppb. Schauinsland is a high elevation site 1205 m.a.s.l, and is more representative of the atmosphere above the boundary layer. The difference between Illmitz and all the other sites is perhaps even more striking for the measurements where very high ozone concentrations were found. In 1985 46 hourly values exceeded 140 ppb at Illmitz, the only other site measuring above 140 ppb was Langenbrügge, and only once. The highest value measured at Illmitz in 1985 was 223 ppb, almost twice as high as the station with the second highest ozone maximum (Jeløya in Norway; 133 ppb).

In Figure 6b-6d the results for the sites in FRG are shown. The measurements at Schauinsland reflect that it is a high altitude site with very slight diurnal variation, indicating that the air which is brought over the station has not recently been near the ground or polluted with NO_x-emissions. The calculated values are much lower than the measured values towards the end of the period (1-3 June).

Deuselbach is located 480 m.a.s.l. on rural land which is partly cultivated and partly pasture. The calculations are lower than measured in the beginning, higher than measured 30-31 May, and fit well towards the end.

The measurements and calculations at the four Scandinavian sites Risø, Rørvik, Langesund and Jeløya are shown in Figures 6e-6h. All these sites are coastal or near coastal. The calculations are quite satisfactory for all four sites. When measured values are fairly low, e.g. 40-50 ppb, and the calculations do not match, it is likely that the initial concentrations for the calculation may not be well chosen and have a strong influence on the calculated receptor point concentrations.

Table 7: Number of hours (h) and days (d) with hourly ozone concentrations exceeding 120, 160, 200, 240 and 280 $\mu\text{g}/\text{m}^3$, and maximum hourly and daily ozone concentration ($\mu\text{g}/\text{m}^3$), April-September 1985. Concentrations in ppb are obtained by division with 2 (Grennfelt et al., 1987).

STATION	Number of hours and days												Maximum ozone concentrations	
	Total		>120		>160		>200		>240		>280		$(\mu\text{g}/\text{m}^3)$	
	h	d	h	d	h	d	h	d	h	d	h	d	h	d
Illmitz	4044	172	2226	168	994	123	405	65	152	32	46	13	446	197
Gent St. Kruiswinkel	3695	161	77	15	38	7	17	4	1	1			253	120
Risø	2875	122	107	22	19	4	3	1					210	146
Brotjacklriegel	3609	152	314	41	8	3							174	109
Deuselbach	4003	170	264	38	31	6							196	109
Langenbrügge-Waldhof	4198	183	473	67	149	23	76	15	15	7	1	1	286	138
Schauinsland	4281	183	1170	91	99	20	4	2					202	127
Westerland	4329	183	101	19	4	2							166	105
Utö	1415	58	26	7	1	1							198	120
Eibergen	3841	165	81	17	18	7							181	99
Witteveen	2582	87	88	16	17	6	3	1					217	112
Birkenes	1401	61											115	70
Jeløya	4306	181	53	12	22	5	14	5	1	1			266	118
Langesund	2648	110	29	5									133	113
Aspvreten	2928	121	272	30	62	6							198	173
Norra Kvill	1830	77	13	2	3	1							194	100
Ringamåla	3300	138	207	28	17	5	1	1					202	130
Rörvik	3954	172	233	39	35	8	5	1					214	127
Vavihill	3447	144	255	32	44	7	2	1					212	141
Vindeln	1808	78											120	65
Bottesford	4375	183	84	12	22	2	9	1					220	133
Harwell	3805	160	121	23	16	4	1	1					206	136
Wray	4223	181	51	11	4	1							176	113
Sibton	2228	99	50	11	5	1							192	125

The results for the Dutch sites Sappermeer and Waarde are shown in Figures 6i-6j. There is some underestimation of the maximum values measured at Sappermeer on 1-2 June, and even much more so in the case of the southern station located in Waarde.

For the French site Colomiers in the Pyrenees, the agreement between measurements and calculations is rather poor (Figure 6k), while for the British sites (Figures 6l-6n) the agreement is good for Bottesford, fair for Sibton and a failure to pick up the high concentrations measured at Stodday.

It should be remembered, however, that the distance between these sites is so small that they are found in the same or neighbouring grid elements. It is not to be expected that the data input to the calculations is sufficiently accurate and resolved in time and space to be able to pick up a peak measured value of about 140 ppb at Stodday and 60-70 ppb at Bottesford and Sibton on 2 June 1982.

To conclude, the agreement between the calculated and measured ozone concentration is fair for the sites in Scandinavia, FRG, UK and the Netherlands, and poor for the French and Austrian sites.

3.3 SENSITIVITY STUDIES OF PHYSICAL AND CHEMICAL PARAMETERS

The calculations with the choice of physical and chemical parameters thought to be most realistic ("reference model"), were compared with the results of calculations where some of the most important parameters were altered. In Table 8 a list is given of the sensitivity calculations which have been carried out.

The frequency distribution of ozone for the reference run and the 7 sensitivity runs is shown in Figure 7. In brief these histograms show that increasing the temperature gives rise to more cases with high calculated ozone concentrations, reducing the mixing height suppresses ozone through the higher concentrations of NO_x which are then calculated in the boundary layer, reducing photolysis rates by 50% strongly reduces the calculated ozone concentrations, assuming clear sky has the opposite effect, reducing ground removal increases ozone and reducing the initial precursor concentrations also results in lower ozone concentrations at the receptor sites.

Table 8: Specification of sensitivity studies.

Run No.	Parameter	Description
0	Reference	
1	Temperature	Increase of Kelvin-temperature by 2% everywhere (i.e. about 5 ⁰ C)
2	Mixing height	Reduction by 50%
3	Solar radiation; photolysis rates	Reduction by 50%
4	Photolysis rates	Assuming clear sky everywhere
5	Ground removal	Reduction of deposition velocities by 50%
6	Initial concentrations	Reduction of precursor concentrations (NO _x , VOC) initially by 90%
7	Chemistry	Surrogate chemistry (Eliassen et al., 1982a).

In the last histogram in Figure 7 and in Figure 8 is shown some results when applying the surrogate chemistry previously used in the Norwegian long-range transport model (Eliassen et al., 1982a). Figure 8 should be compared with Figures 6g and 6l. The surrogate chemistry generally gives rise to somewhat higher ozone concentrations than CBM-X, a result also reported earlier (Hov et al., 1986). The reactivity split may partly explain the differences. In the lumped chemistry the split explained in section 2.4 was used, and this results probably in a more reactive mixture of VOC-emissions than in the case of the split used for CBM-X (Table 6).

In Table 9 the results shown in Figures 7 and 8 are presented for each station and added up over all stations, while in Table 10 is shown the number of cases where O₃ > 60 ppb, PAN > 1 ppb, OH > 6 x 10⁶ molecules/cm³ or NO₂ > 12 ppb. For each site the concentration field was calculated for 25 arrival times (every 6 h from 28 May 1200 GMT to 3 June 1200 GMT 1982), added up 25 x 14 = 350 concentration fields were calculated.

CLEAR SKY																
Site No.	1	2	3	4	5	6	7	8	9	10	11	12	13	14	Sum	%
> 50	11	16	20	19	9	10	11	10	11	12	0	9	11	5	154	44.0
> 60	5	9	14	17	7	7	8	7	9	12	0	4	7	1	107	30.6
> 70	2	4	7	8	4	6	3	3	7	7	0	3	5	0	59	16.9
> 80	0	2	4	3	1	2	1	0	5	4	0	1	3	0	26	7.4
> 90	0	1	0	1	0	0	0	0	2	1	0	0	0	0	5	1.4

DEPOSITION VELOCITIES REDUCED BY 50%																
Site No.	1	2	3	4	5	6	7	8	9	10	11	12	13	14	Sum	%
> 50	14	16	19	18	7	9	10	9	9	11	0	7	10	5	144	41.1
> 60	5	7	14	14	6	7	8	4	9	9	0	4	6	1	94	26.9
> 70	2	3	5	5	4	4	1	2	6	4	0	2	3	0	41	11.7
> 80	0	2	4	2	0	1	0	0	3	2	0	1	2	0	17	4.9
> 90	0	1	0	0	0	0	0	0	1	0	0	0	0	0	2	.6

INITIAL PRECURSOR EMISSIONS REDUCED BY 90%																
Site No.	1	2	3	4	5	6	7	8	9	10	11	12	13	14	Sum	%
> 50	4	2	5	5	0	2	3	2	3	3	0	2	2	0	33	9.4
> 60	1	1	4	2	0	1	1	0	1	2	0	1	1	0	15	4.3
> 70	0	0	0	0	0	0	1	0	0	0	0	0	0	0	1	.3
> 80	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	.0
> 90	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	.0

SURROGATE CHEMISTRY MECHANISM																
Site No.	1	2	3	4	5	6	7	8	9	10	11	12	13	14	Sum	%
> 50	20	16	18	18	9	9	11	10	8	13	5	10	9	7	163	46.6
> 60	9	9	17	12	6	3	8	8	7	8	0	6	6	2	101	28.9
> 70	4	3	11	6	3	2	5	3	5	5	0	3	4	0	54	15.4
> 80	3	2	4	3	0	1	1	1	1	2	0	3	2	0	23	6.6
> 90	0	1	0	0	0	0	0	0	0	0	0	1	0	0	2	.6

Table 10 shows that this conclusion is valid also for PAN, OH and NO₂. It is seen that an increase in temperature caused a drop in the number of high PAN-concentrations, to be expected from the temperature dependence of the thermal decomposition rate of PAN. A lower mixing height gave rise to a larger fraction of high PAN and NO₂-concentrations than in the reference case. Lower deposition velocities had less influence on PAN and NO₂ than on O₃, indicating that ground removal is a more important loss process for O₃ than for PAN or NO₂. The surrogate chemistry scheme gave higher O₃, PAN and NO₂, but lower OH.

Table 10: Summary of influence on O₃, PAN, NO₂ and OH at all 14 receptor sites during the period 28 May² - 3 June 1982 (1200 GMT) of changes in physical and chemical parameters (cfr. Table 8).

Run	% of total number of trajectories (350)			
	O ₃ >60 ppb	PAN >1 ppb	OH >6x10 ⁶ $\frac{\text{molecules}}{\text{cm}^3}$	NO ₂ >12 ppb
0 ref	17.4	28.6	13.4	7.4
1 temperature	28.3	18.0	19.1	4.9
2 mixing height	6.9	25.4	7.1	32.0
3 photolysis	0.0	2.9	0.3	17.4
4 photolysis	30.6	36.9	20.9	5.1
5 v _g	26.9	33.7	14.3	8.0
6 C _o	4.3	14.0	12.6	10.3
7 chemistry	28.9	32.0	11.1	12.6

3.4 EMISSION REDUCTIONS AND CHANGES IN O₃, PAN AND NO₂

Calculations have been carried out to see how the concentrations of O₃, PAN and NO₂ at the 14 receptor sites specified in Table 6, change during the 28 May - 3 June 1982 period with changes in the emissions of NOx and VOC.

Uniform emission changes were carried out throughout the EMEP grid, and the scenarios are outlined in Table 11. The runs h-n are similar to the runs a-g with the NOx-emissions reduced by a factor 2. This means that the "reference case" for the runs h-n is run d, where NOx is reduced by 50% compared to run 0. Since run 0 corresponds to a run where the NOx-emissions are 1.4 times the EMEP annual average emissions on a daily basis, run d corresponds to a case where the NOx-emissions are 70% of the annual average. This is what normally is assumed for May/June-emissions in the EMEP calculations.

Table 13: Percentage of trajectories with more than 60 ppb O_3 at the arrival point for the sites in each of 4 geographical areas and for the sum of those sites. Time period 28 May 1200 GMT-3 June 1982 1200 GMT (25 trajectories per site).

Run	Description		FRG sites (2+3+4)	Scandinavian sites (5+6+7+8)	NL sites (9+10)	UK sites (12+13+14)	Sum except 1 and 11
	$\Delta NO_x(\%)$	$\Delta HC(\%)$					
0	0	0	29.3	17.0	22.0	9.3	19.0
a	0	-25	9.3	2.0	2.0	4.0	4.3
b	0	-50	0.0	0.0	0.0	0.0	0.0
c	-25	0	46.7	28.0	40.0	16.0	31.7
d	-50	0	49.3	30.0	44.0	28.0	36.7
e	-25	-25	21.3	17.0	22.0	8.0	16.7
f	+25	+25	33.3	13.0	20.0	9.3	18.3
g	-75	-75	0.0	0.0	2.0	0.0	0.3
h	-50	-25	33.3	24.0	34.0	13.3	25.3
i	-50	-50	12.0	14.0	14.0	4.0	11.0
j	-62.5	0	41.3	26.0	46.0	25.3	33.0
k	-75	0	24.0	22.0	32.0	17.3	23.0
l	-62.5	-25	32.0	24.0	36.0	18.7	26.7
m	-37.5	25	60.0	35.0	50.0	33.3	43.3
n	-87.5	-75	0.0	0.0	0.0	0.0	0.0

In Table 12 is shown for each station and added over the station, the number of trajectories with O_3 exceeding some specified value, for the reference run and for the seven first scenarios described in Table 11.

In Table 13 is shown the percentage of trajectories with more than 60 ppb of ozone at the arrival point for the sites in each of the geographical areas FRG (sites 2-4), Scandinavia (5-8), The Netherlands (9,10) and UK (12-14) and the sum of these 12 sites.

The following conclusions can be drawn:

- Reduction in VOC-emissions by 25%, gives rise to a significant ozone reduction at all sites, a 50% reduction is calculated to keep all O_3 -values below 60 ppb.
- Reduction in NO_x -emission by 25% is calculated to increase O_3 at all sites, except Illmitz and Colomiers. At the FRG sites, the number of trajectories with more than 60 ppb upon arrival increases from 29.3 to 46.7%, the corresponding numbers for the Scandinavian

sites are 17.0 to 28.0%, for the Netherlands 22.0% to 40.0% and for the UK 9.3 to 16.0%.

- A reduction in NO_x-emissions by 50% is calculated to further increase O₃ at most sites, in particular for the UK sites where the percentage of trajectories with O₃ > 60 ppb is 9.3% in the reference case, 16.0% in the ΔNO_x = -25% case and 28.0% in the ΔNO_x = -50% case.
- A reduction in NO_x and VOC emissions by 25% each is calculated to decrease O₃ slightly, mainly for the FRG-sites (from 29.3 to 21.3% of the trajectories with O₃ > 60 ppb).
- An increase in NO_x and VOC emissions by 25% each is calculated to increase the number of O₃-values above 70 ppb from 22 to 25 added over all sites, above 80 ppb from 4 to 9 cases. Otherwise the influence on ozone is slight except in Scandinavia where ozone drops.
- Reduction in NO_x and VOC emissions by 75% each is calculated to reduce O₃ significantly everywhere.

The effect of reducing the NO_x-emissions and VOC-emissions, separately or together, was also calculated starting with ΔNO_x = -50%, ΔVOC = 0 as the reference case (run d). This corresponds to what is assumed as NO_x-emissions for May-June 1982 in the EMEP-inventory given that the annual NO_x-emissions are comparable in 1980 and 1982. The effect of emission reductions starting from the new reference case, is quite different from runs a-g.

- A separate reduction in VOC- or NO_x- emissions is calculated to reduce ozone almost always at all sites.
- A combined reduction in both VOC- and NO_x- emissions is often more efficient than reducing only one of them by the same percentage.
- Increase in VOC- and NO_x- emissions gives rise to an increase in O₃.

In Table 14 it can be seen that the calculated impact of emission changes on the fraction of trajectories with more than 60 ppb of O₃ upon arrival, is much the same also for the trajectories with more than 1 ppb of PAN. For NO₂ the result is different, reducing HC-emissions is calculated to increase NO₂, reducing NO_x-emissions, reduce NO₂.

Table 14: Summary of the influence of emission changes on the concentration of O₃, PAN and NO₂ at all 14 receptor sites during the period 28 MAY - 3 June 1982 (1200 GMT), cfr. Table 11.

Run	Description		% of total number of trajectories (350)		
	ΔNO _x	ΔVOC	O ₃ > 60 ppb	PAN > 1 ppb	NO ₂ > 12 ppb
0	0	0	17.4	28.6	7.4
a	0	-25	4.3	10.0	10.9
b	0	-50	0.0	0.9	12.0
c	-25	0	28.3	40.0	0.6
d	-50	0	32.3	41.7	0.0
e	-25	-25	14.9	23.1	1.4
f	+25	+25	16.9	34.3	12.6
g	-75	-75	0.6	2.6	0.0
h	-50	-25	22.3	31.7	0.0
i	-50	-50	9.4	13.4	0.0
j	-62.5	0	28.9	39.4	0.0
k	-75	0	20.0	35.1	0.0
l	-62.5	-25	23.1	32.9	0.0
m	-37.5	25	38.6	51.7	0.0
n	-87.5	-75	0.0	2.0	0.0

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Figure Captions:

Figure 1: Map of Europe with the grid system used in the EMEP model.

Figure 2: EMEP grid system with the PHOXA (RTM III) grid and the extension of the PHOXA-grid used in calculations for OECD.

Figure 3: 96 h, 850 mb backtrajectories to 14 sites (Illmitz (1), Langenbrügge (2), Schauinsland (3), Deuselbach (4), Risø (5), Rørvik (6), Langesund (7), Jeløy (8), Sappermeer (9), Waarde (10), Colomiers (11), Bottesford (12), Sibton (13), Stodday (14)) for 1200 GMT each of the days 28 May - 3 June 1982.

Figure 4: Contours of constant mixing height in m at 1200 GMT each of the days 28 May - 3 June 1982.

Figure 5: Contours of constant relative humidity (in %) at 1200 GMT, 31 May 1982.

Figure 6: Hourly measured ozone concentrations vs. calculated ozone concentrations every 6 h, as a function of time starting at 1200 GMT 28 May 1982 (Day number 1) ending at 1200 GMT 3 June 1982 (Day number 7).

Figure 7: Frequency distribution of ozone (number of trajectories with ozone concentration in the interval 1: 0-5 ppb, 2: 5-10 ppb, 3: 10-15 ppb,, 10: 45-50 ppb,, 15: 70-75 ppb,, 22: 105-110 ppb, the sum of trajectories is 350), for the reference case (7,0) and for the 7 sensitivity runs specified in Table 8 (1: 2% increase in absolute temperature, 2: H_{mix} reduced by 50%, 3: photolysis rates reduced by 50%, 4: clear sky, 5: v_g reduced by 50%, 6: c_o of precursors reduced by 90%, 7: surrogate chemistry (Eliassen et al., 1982a)).

Figure 8: Measured vs. calculated ozone for Langesund and Bottesford using surrogate chemistry.

APPENDIX

Reproduced from Whitten et al. (1984).

The expanded Carbon-Bond Mechanism (CBM-X).

REACTIONS			RATE CONSTANTS ^b	ACTIVAT. ENERGY (K) ^c
1	NO ₂	= NO + O	vary	
2	O + O ₂ + M	= O ₃	2.22-5	-690.
3	O ₃ + NO	= NO ₂	26.6	1430.
4	O + NO ₂	= NO	1.38 + 4	
5	O + NO ₂ + M	= NO ₃	2.32-3	-600.
6	O + NO + M	= NO ₂	0.00312	-411.
7	NO ₂ + O ₃	= NO ₃	4.74-2	2450.
8	O ₃	= O1D	vary	
9	O ₃	= O	4.2 - 2 x k ₁	
10	O1D + M	= O	4.29 + 4	-100.
11	O1D + H ₂ O	= OH + OH	3.26 + 5	
12	O ₃ + OH	= HO ₂	100.0	940.
13	O ₃ + HO ₂	= OH	3.0	580.
14	NO ₃ + NO	= NO ₂ + NO ₂	2.81 + 4	-250.
15	NO ₃ + NO ₂	= NO + NO ₂	0.59	1230.
16	NO ₃ + NO ₂ + M	= N ₂ O ₅	1776.0	-226.
17	N ₂ O ₅ + H ₂ O	= HNO ₃ + HNO ₃	1.9-6	
18	N ₂ O ₅ + M	= NO ₃ + NO ₂	3.12-6	10650.
19	NO + NO ₂ + H ₂ O	= HNO ₂ + HNO ₂	1.6-11	
20	HNO ₂ + HNO ₂	= NO + NO ₂ + H ₂ O	1.5-5	
21	HNO ₂	= NO + OH	0.18 x k ₁	
22	NO ₂ + OH + M	= HNO ₃	1.63-2	-560.
23	NO + OH + M	= HNO ₂	9.77-3	-427.
24	HO ₂ + NO	= OH + NO ₂	1.23 + 4	-240.
25	NO + NO + O ₂	= NO ₂ + NO ₂	7.28-10	-530.
26	H ₂ O ₂	= OH + OH	1.4-3	
27	OH + H ₂ O ₂	= HO ₂	2520.0	187.
28	OH + HNO ₂	= NO ₂	9770.0	
29	OH + HNO ₃	= NO ₃	192.0	-778.
30	NO ₃	= NO ₂ + O	26.0 x k ₁	
31	NO ₃	= NO	4.6 x k ₁	
32	HO ₂ + HO ₂	= H ₂ O ₂	4144.0	-1150.
33	HO ₂ + HO ₂ + H ₂ O	= H ₂ O ₂ + H ₂ O	0.218	-5800.
34	HO ₂ + NO ₂ + M	= PNA	1.63-3	-617.
35	PNA + M	= HO ₂ + NO ₂	5.12-6	10420.

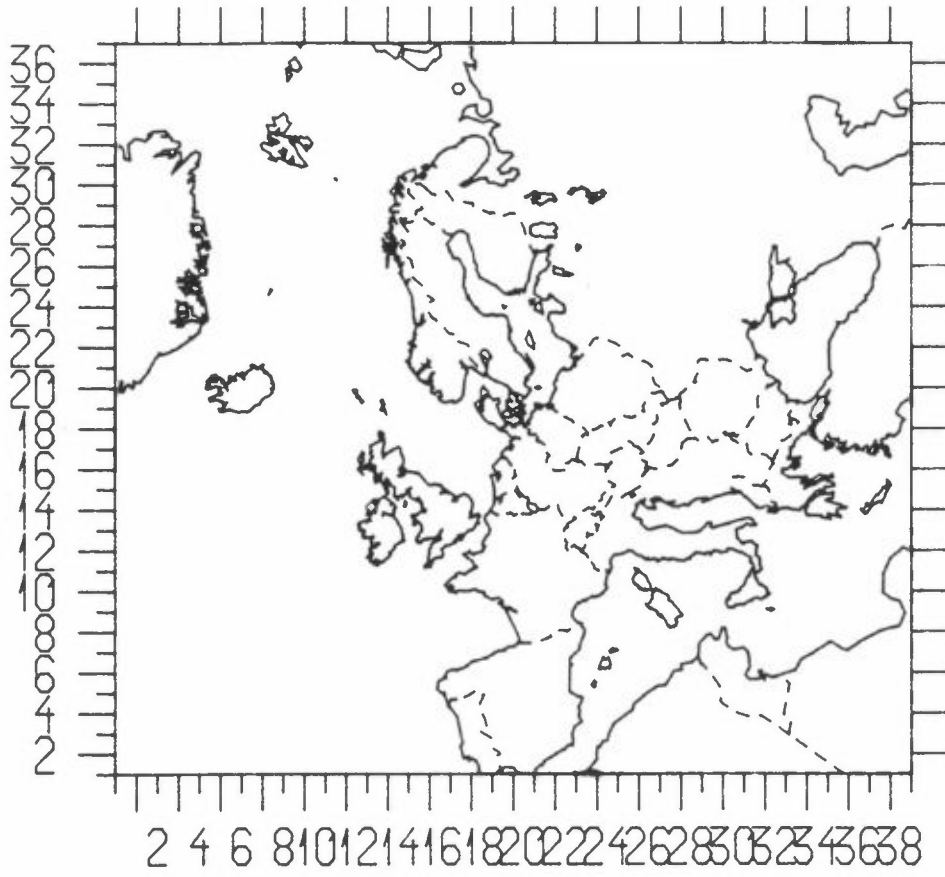
36	OH + CO + M	= HO ₂	4.0-4	
37	FORM + HO ₂	= FROX	14.8	
38	FORM + OH	= HO ₂ + CO	1.5 + 4	
39	FORM	= HO ₂ + HO ₂ + CO	vary	
40	FORM	= CO	vary	
41	FORM + O	= OH + HO ₂ + CO	237.0	1550.
42	FORM + NO ₃	= HNO ₃ + HO ₂ + CO	0.93	
43	FROX + NO	= FACD + NO ₂ + HO ₂	10400.0	
44	FROX	= HO ₂ + FORM	90.0	
45	ALD2 + O	= C ₂ O ₃ + OH	636.0	986.
46	ALD2 + OH	= C ₂ O ₃	24000.0	-250.
47	ALD2 + NO ₃	= C ₂ O ₃ + HNO ₃	3.7	
48	ALD2	= MEO ₂ + HO ₂ + CO	vary	
49	ALD2 + HO ₂	= MEO ₂ + FORM	5.0	
50	C ₂ O ₃ + NO	= MEO ₂ + NO ₂	16500.0	-250.
51	C ₂ O ₃ + NO ₂	= PAN	9000.0	-250.
52	PAN	= C ₂ O ₃ + NO ₂	0.0222	14000.
53	MEO ₂ + NO ₂ + M	= MPNA	6.00-3	-735.
54	MPNA + M	= MEO ₂ + NO ₂	9.2-5	10400.
55	MEO ₂ + NO	= MEO + NO ₂	11000.0	-180.
56	MEO + NO	= MNIT	44400.0	-200.
57	MEO + NO	= FORM + HO ₂ + NO	1920.0	
58	MEO + NO ₂	= MEN3	22200.0	
59	MEO + O ₂	= FORM + HO ₂	1.88	1313.
60	MEN3 + OH	= FORM + NO ₂	2220.0	360.
61	MNIT + OH	= FORM + NO	2370.0	340.
62	MNIT	= MEO + NO	0.3 x k ₁	
63	MEO ₂ + MEO ₂	= MEO + MEO	0 75.0	-220.
64	MEO ₂ + MEO ₂	= FORM + MEOH	328.0	-220.
65	MEO ₂ + C ₂ O ₃	= MEO + MEO ₂	4400.0	
66	C ₂ O ₃ + C ₂ O ₃	= MEO ₂ + MEO ₂	3700.0	
67	MEO ₂ + O ₂	= FORM	8900.0	-1300.
68	C ₂ O ₃ + HO ₂	= PROX	9600.0	
69	AONE	= MEO ₂ + C ₂ O ₃	0.0002 x k ₁	
70	AONE + OH	= ANO ₂	580.0	
71	ANO ₂ + NO	= NO ₂ + FORM + C ₂ O ₃		12000.0
72	GLY	= CO + CO	0.0052 x k ₁	
73	GLY	= CO + FORM	0.0016 x k ₁	
74	GLY	= HO ₂ + HO ₂ + CO + CO		0.0007

75	GLY + OH	= HO ₂ + CO + CO	15000.0	
76	MGLY	= C ₂ O ₃ + HO ₂ + CO		0.02 x k ₁
77	MGLY + OH	= MGPX	26000.0	
78	MGPX + NO	= C ₂ O ₃ + NO ₂	12000.0	
79	OH + M	= MEO ₂	2.1-5	
80	PAR + OH	= RO ₂	150.0	
81	PAR + OH	= RO ₂ R	1000.0	
82	RO ₂ + NO	= NO ₂ + HO ₂ + ALD2 + X	12000.0	
83	RO ₂ R + NO	= NO ₂ + ROR	12000.0	
84	RO ₂ R + NO	= NTR	1000.0	
85	ROR + NO ₂	= NTR	22000.0	
86	ROR	= KET + HO ₂	150000.0	
87	ROR	= ALD2 + D	90000.0	7000.
88	ROR	= AONE + D + X + X	240000.0	
89	X + PAR		10000.0	
90	D + PAR	= RO ₂	3000.0	
91	D + PAR	= AO ₂ + X + X	7000.0	
92	D + KET	= C ₂ O ₃	10000.0	
93	AO ₂ + NO	= NO + HO ₂ + AONE	12000.0	
94	KET	= C ₂ O ₃ + RO ₂ + X + X		0.0003 x k ₁
95	O + OLE	= MEO ₂ + C ₂ O ₃ + X	300.0	324.
96	O + OLE	= RO ₂ + HO ₂ + CO + X		880.
97	O + OLE	= ALD2	4740.0	324.
98	OH + OLE	= MEO ₂ + ALD2 + X	42000.0	-537.
99	O ₃ + OLE	= ALD2 + CRIG	0.0054	1897.
101	O ₃ + OLE	= ALD2 + HOTA	0.0036	1897.
102	O ₃ + OLE	= FORM + HTMA	0.0036	1897.
103	NO ₃ + OLE	= PNO ₂	11.4	
104	O + ETH	= MEO ₂ + HO ₂ + CO	1080.0	800.
105	OH + ETH	= MEO ₂ + FORM	1.2E + 4	-382.
106	O ₃ + ETH	= FORM + CRIG	0.001	2840.
107	O ₃ + ETH	= FORM + HOTA	0.0017	2840.
108	HOTA		25.0	
109	HOTA	= CO	58.0	
110	HOTA	= HO ₂ + HO ₂	10.0	
111	HOTA	= FACD	7.0	
112	HTMA		200.0	
113	HTMA	= MEO ₂ + CO + OH	400.0	

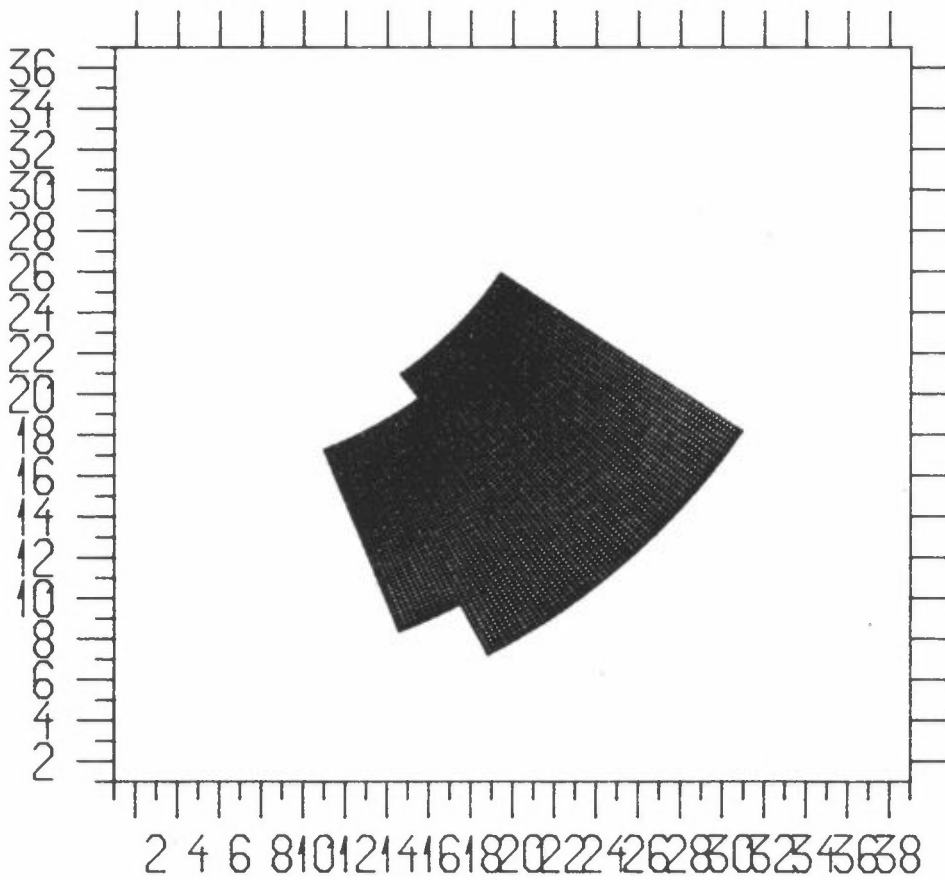
114	HTMA	= MEO ₂ + HO ₂	320.0	
115	HTMA	= HO ₂ + CO + FORM + HO ₂	80.0	
116	CRIG + NO	= FORM + NO ₂	10000.0	
117	CRIG + NO ₂	= FORM + NO ₃	1000.0	
118	CRIG + H ₂ O	= FACD	0.006	
119	CRIG + FORM	= OZD	30.0	
120	CRIG + ALD2	= OZD	30.0	
121	MCRG + NO	= ALD2 + NO ₂	10000.0	
122	MCRG + NO ₂	= ALD2 + NO ₃	1000.0	
123	MCRG + H ₂ O	= ACAC	0.006	
124	MCRG + FORM	= OZD	30.0	
125	MCRG + ALD2	= OZD	30.0	
126	PNO ₂ + NO	= DNIT	1000.0	
127	PNO ₂ + NO	= NTR + HO ₂ + NO ₂	10000.0	
128	TOL + OH	= OPEN + GLY	5500.0	
129	TOL + OH	= PHEN + HO ₂	3500.0	
130	TOL + OH	= BO ₂	750.0	
131	OPEN + NO	= NO ₂ + HO ₂ + MGLY + GLY	10000.0	
132	PHEN + NO ₃	= PHO + HNO ₃	14000.0	
133	BZA + OH	= BZO ₂	20000.0	
134	BZO ₂ + O ₂	= PBZN	2500.0	
135	BZO ₂ + NO	= NO ₂ + PHO ₂ + CO	3700.0	
136	BO ₂ + NO	= NO ₂ + BZA + HO ₂	12000.0	
137	PBZN	= BZO ₂ + O ₂	0.022	14000.
138	BZA		0.004 x k ₁	
139	PHO + NO ₂	= NPHN	20000.0	
140	PHO ₂ + NO	= NO ₂ + PHO	12000.0	
141	XYL + OH	= XYLO	2000.0	
142	XYLO + NO	= NO ₂ + HO ₂ + TLA	12000.0	
143	TLA + OH	= TLO ₂	20000.0	
144	TLO ₂ + NO	= NO ₂ + PHO	4000.0	
145	XYL + OH	= HO ₂ + PHEN	10000.0	
146	XYL + OH	= OPEN + MGLY	24000.0	

b All rate constants are in appropriate ppm and min units at 298K and one atmosphere of air. That is, all unimolecular and photolysis reactions (+ hv) are in units of min⁻¹, all bimolecular reactions are in units of ppm⁻¹ min⁻¹, and all ter-molecular reactions are in units of ppm⁻² min⁻¹. Powers of ten are indicated by + or - n.

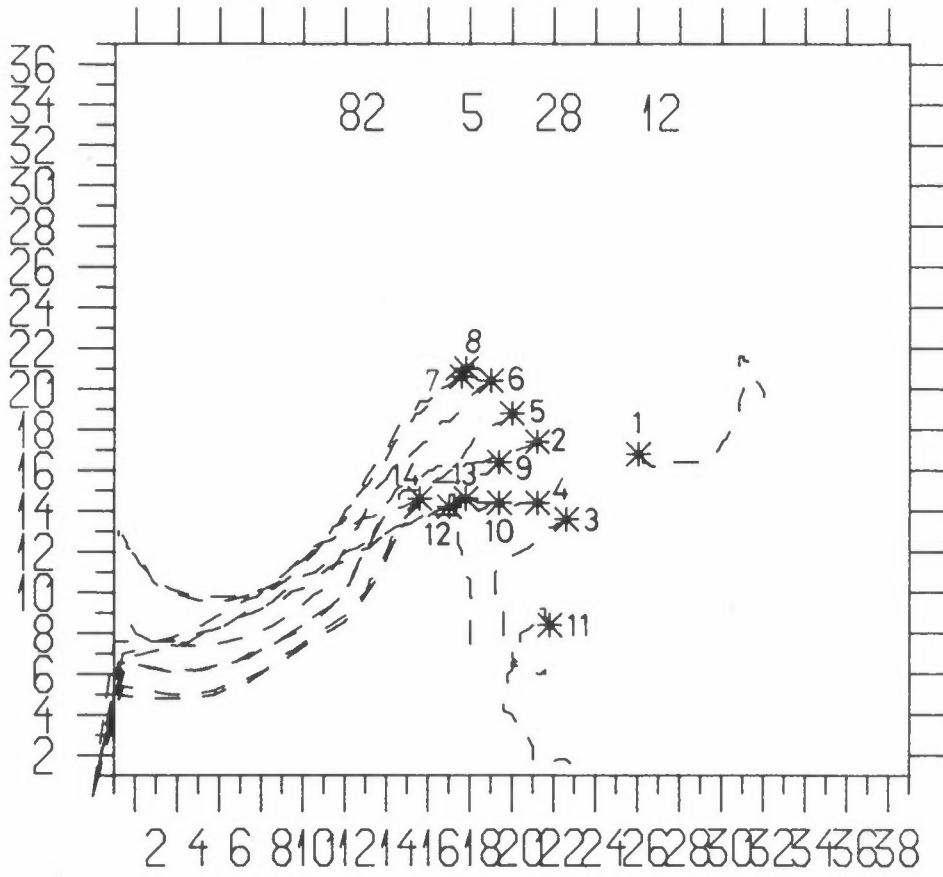
c The numbers listed are Arrhenious temperature coefficients in degrees Kelvin for B in the expression $k = A \exp(-B/T)$.



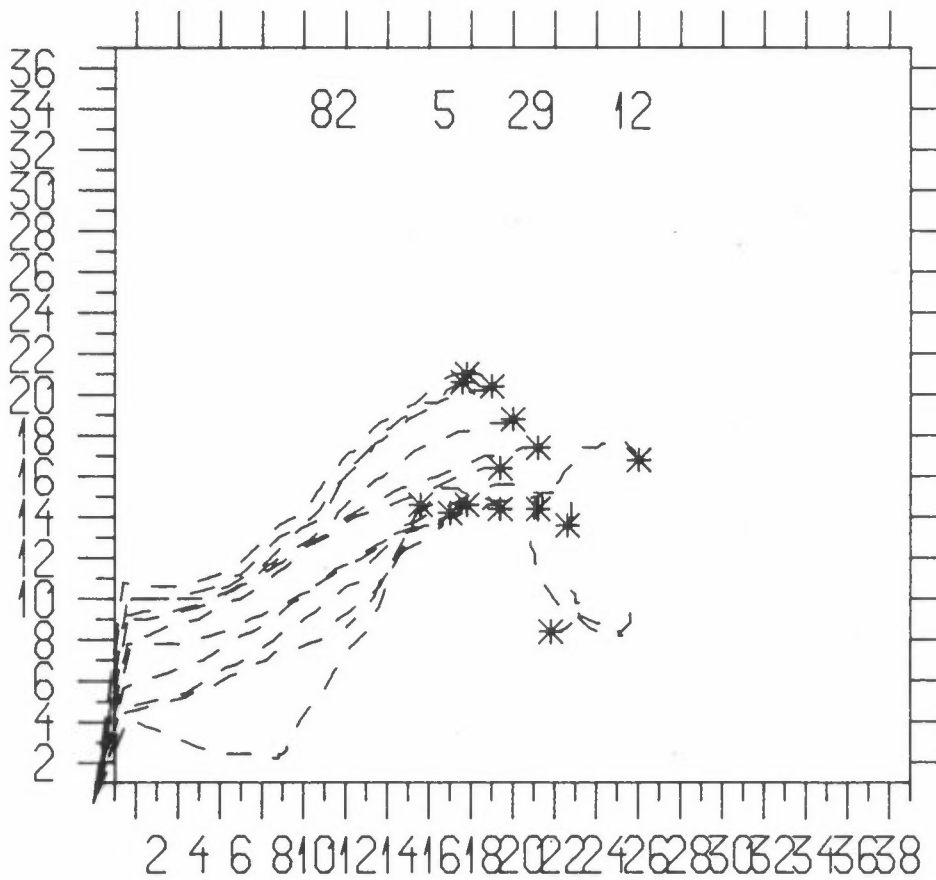
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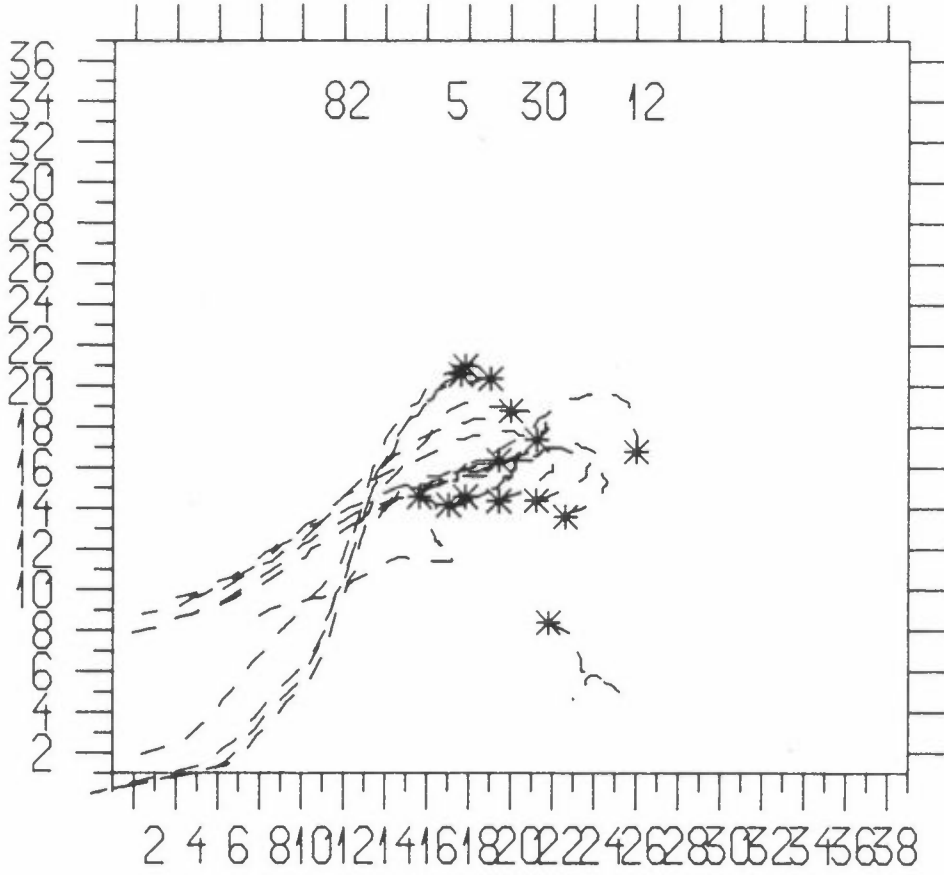
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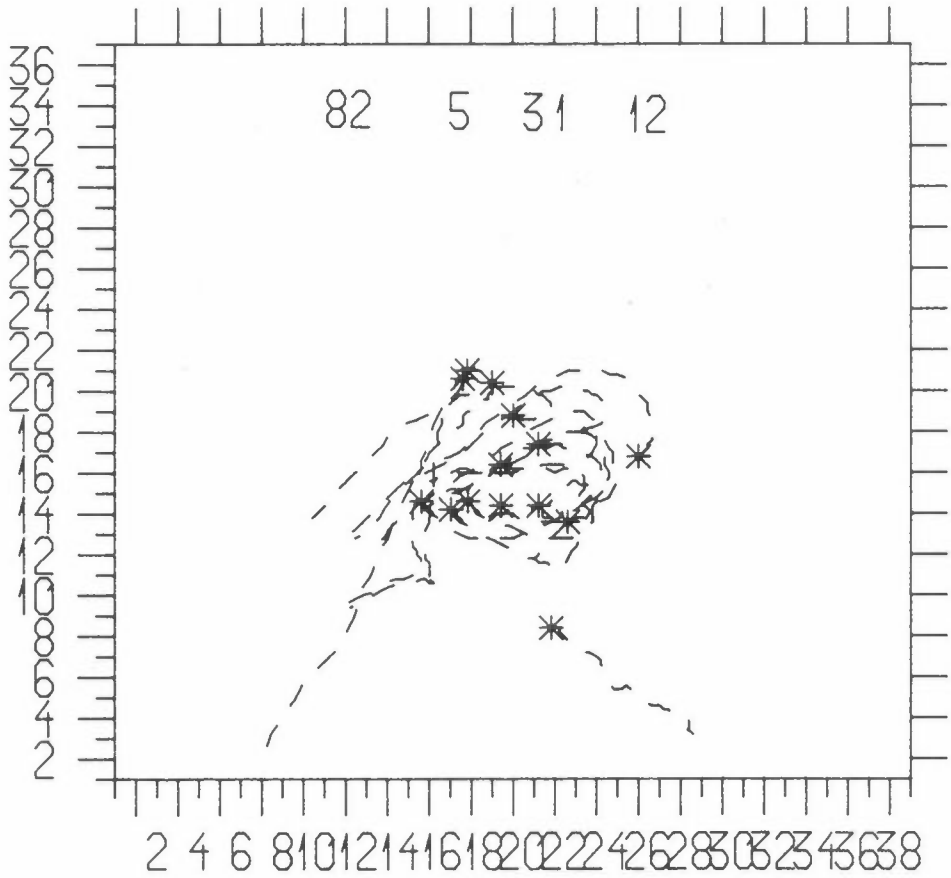
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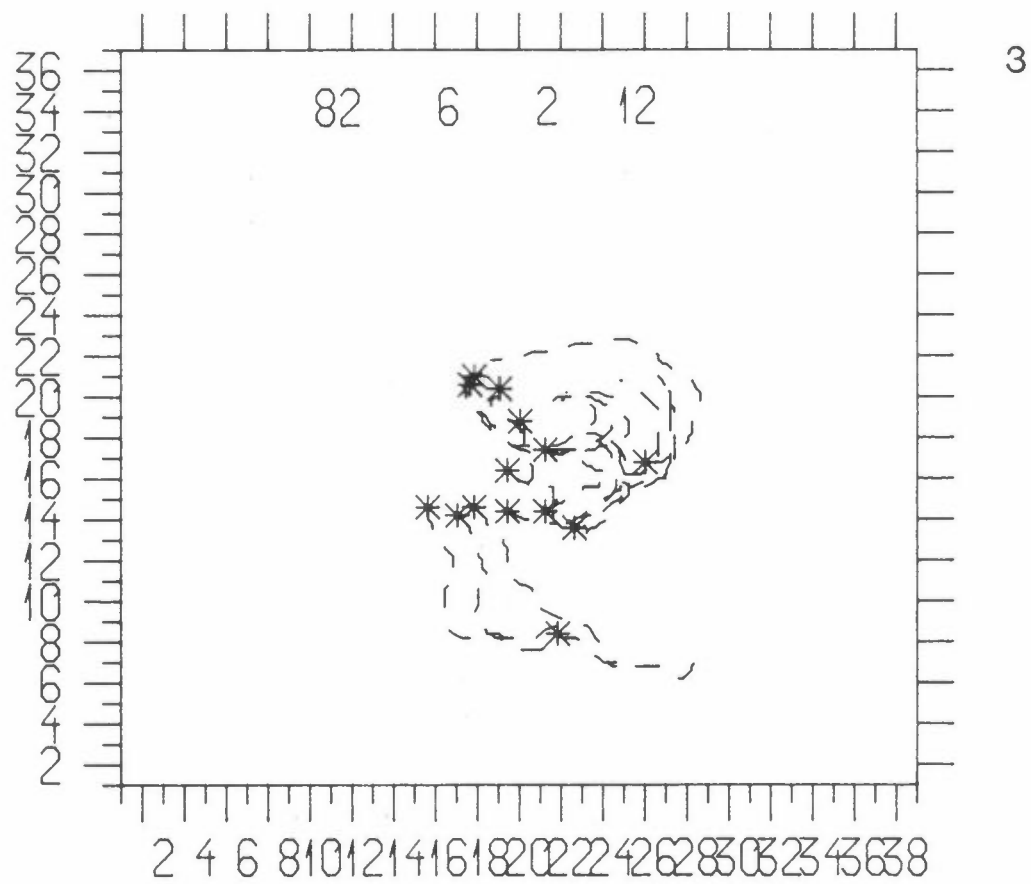
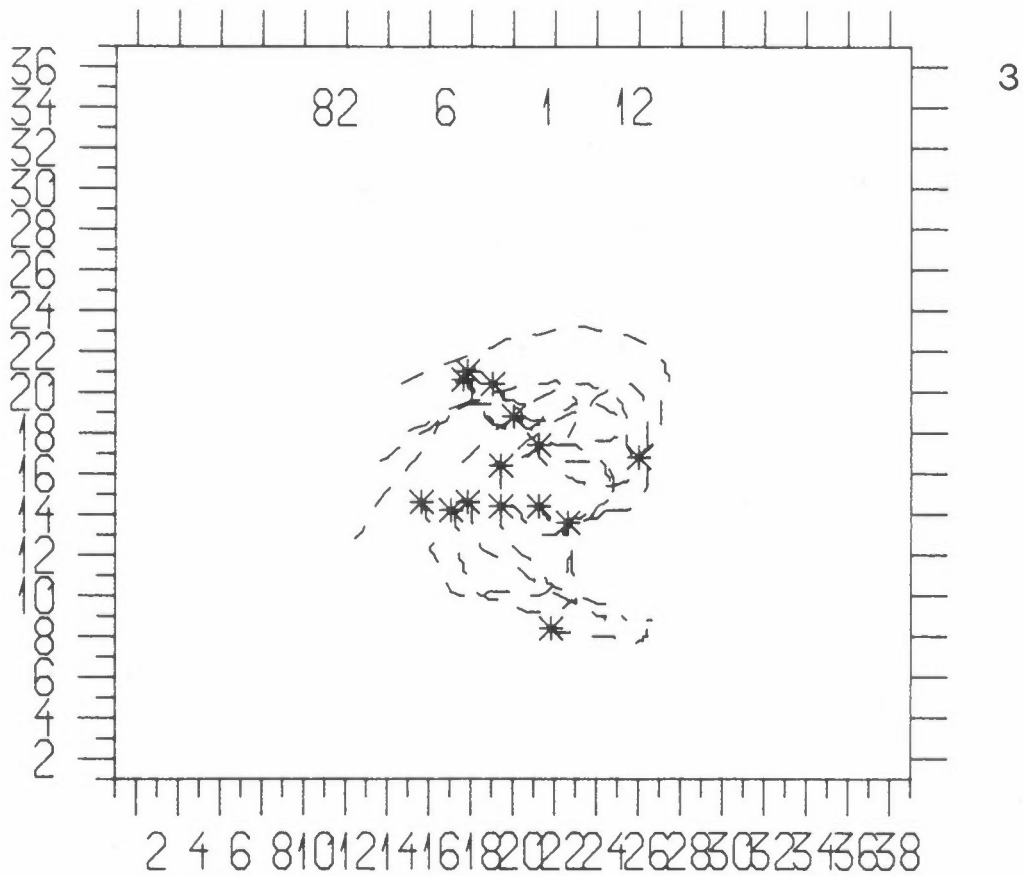
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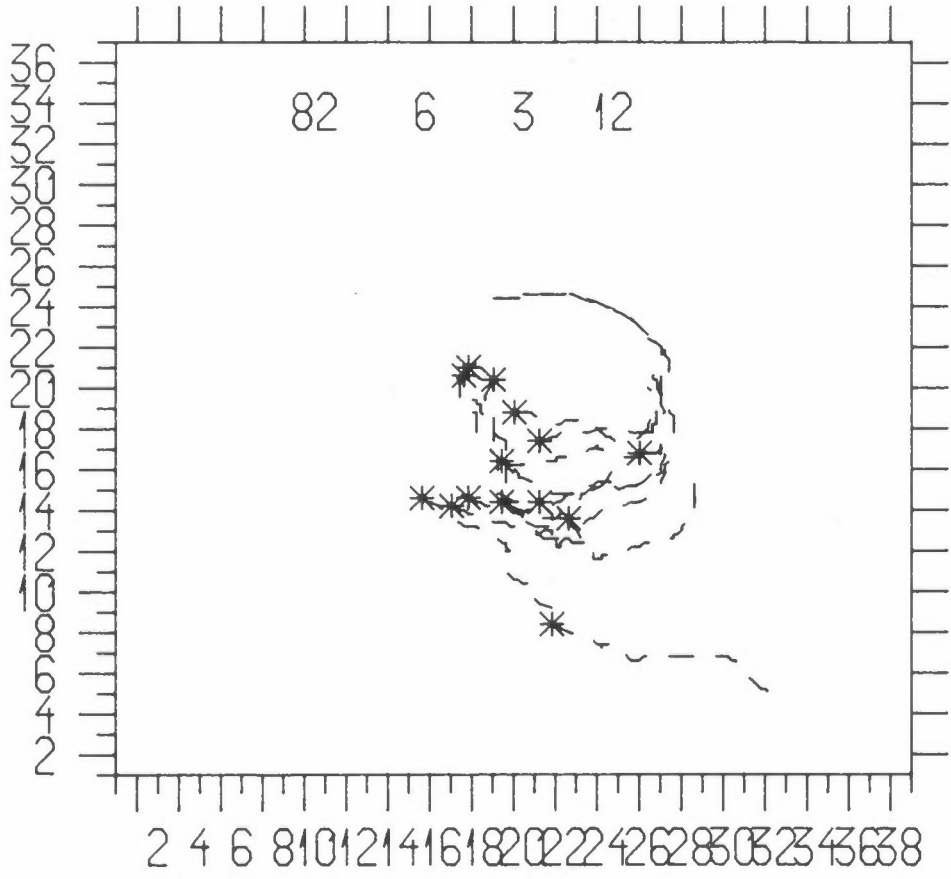


3



3



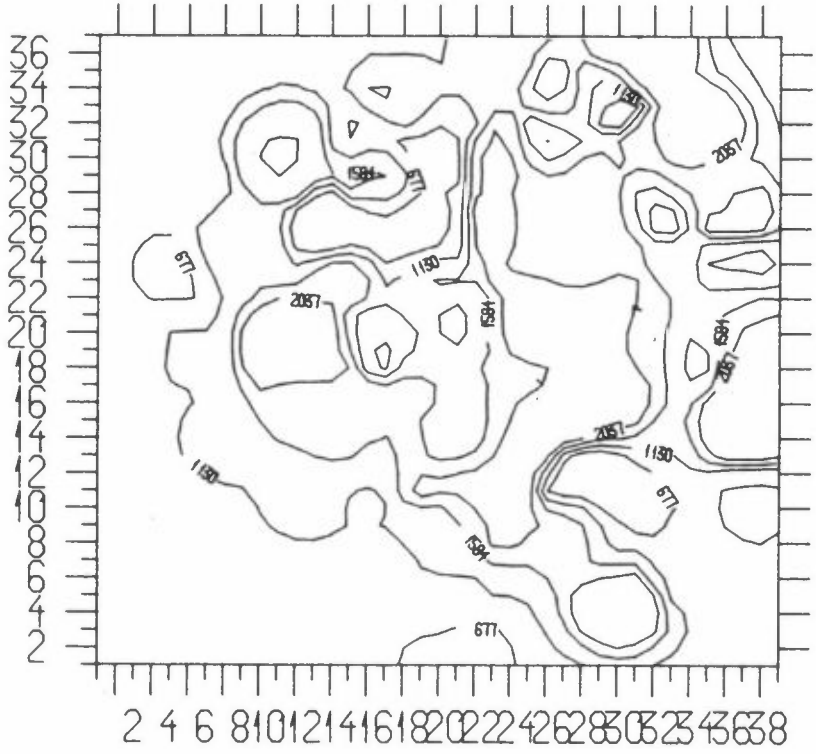


3

82 5 28 12

MIXING HEIGHT

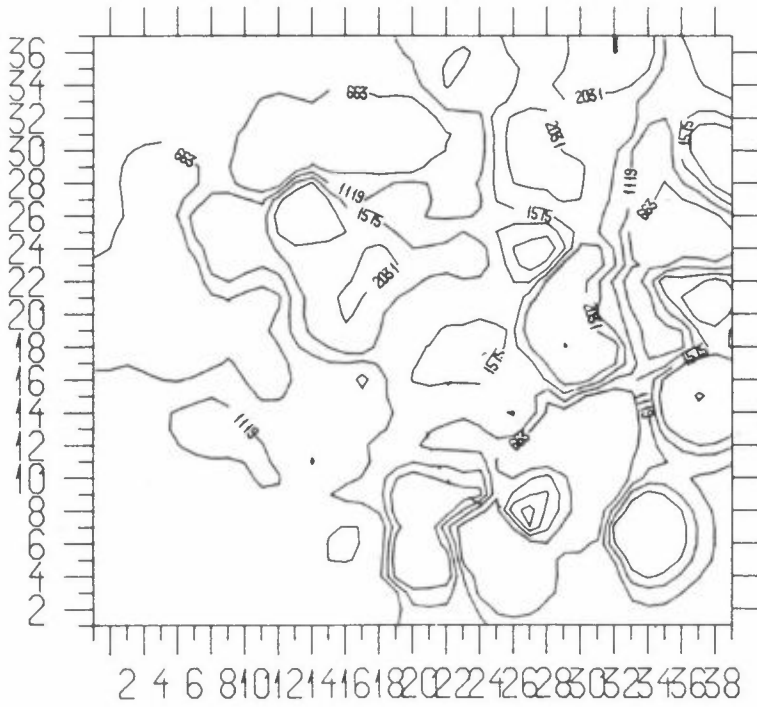
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CONTOURS FROM 224. TO 2491., INTERVAL 453.

82 5 29 12

MIXING HEIGHT

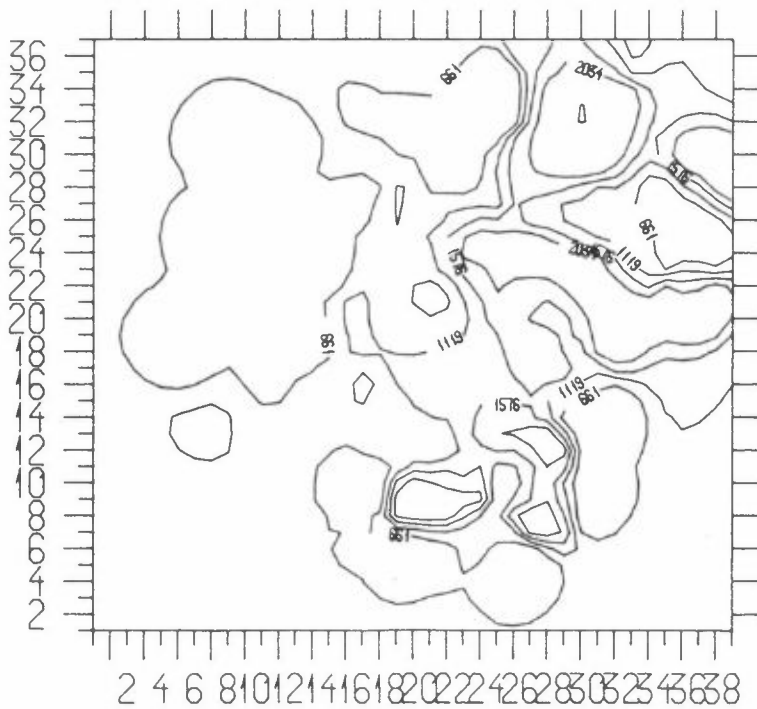


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CONTOURS FROM 208. TO 2487., INTERVAL 456.

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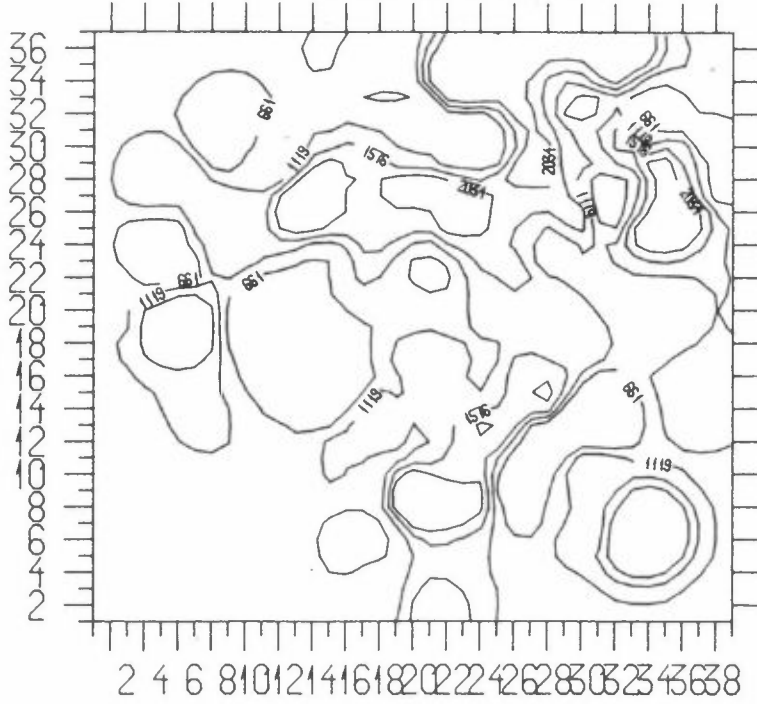


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CONTOURS FROM 204. TO 2492., INTERVAL 458.

82 5 31 12

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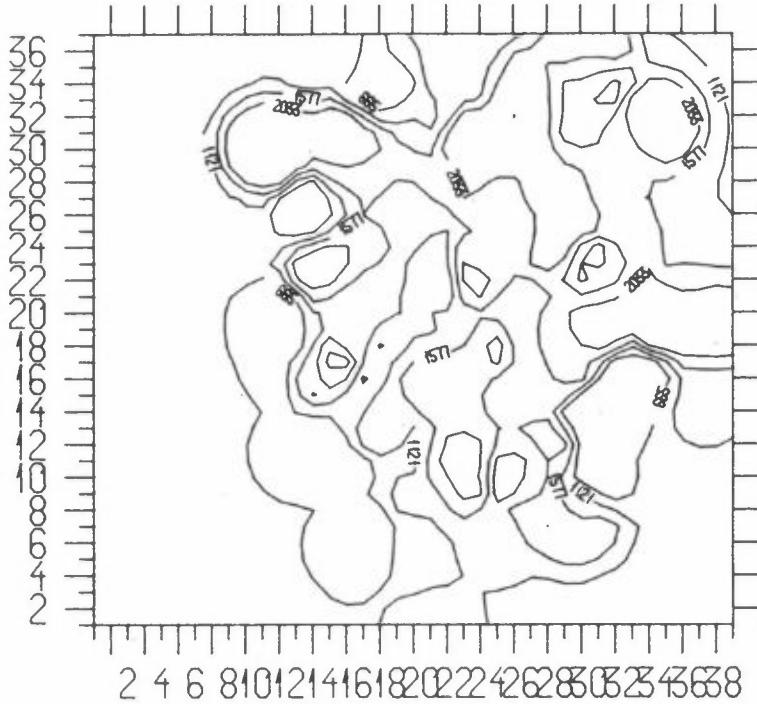


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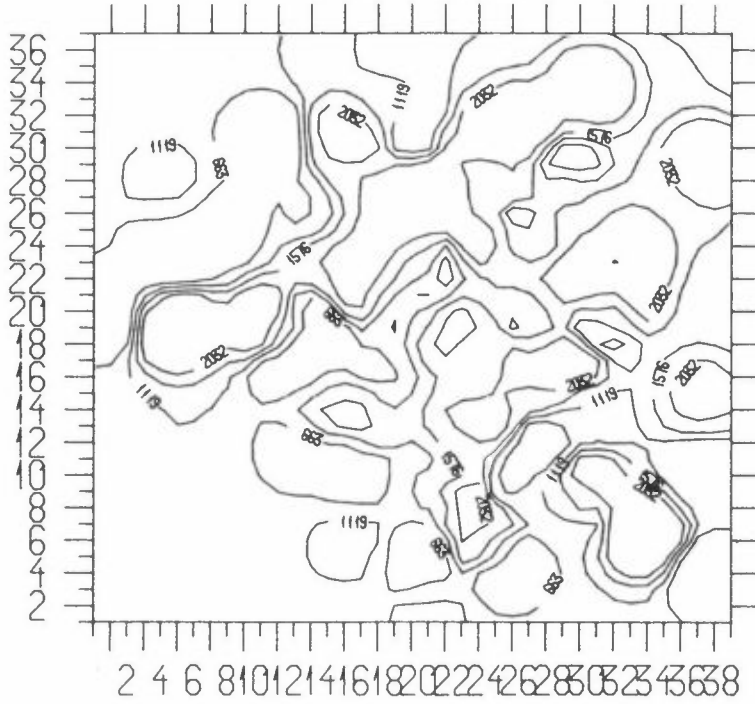


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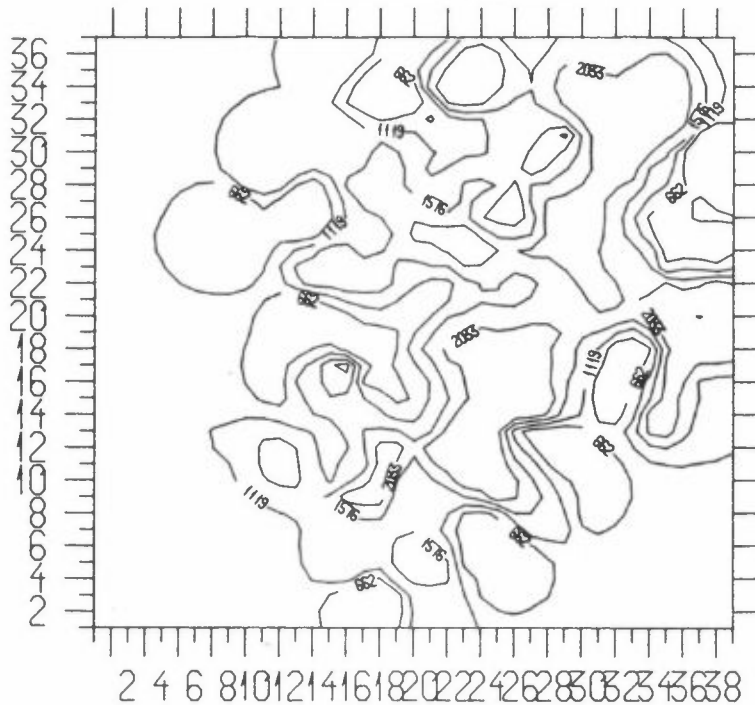


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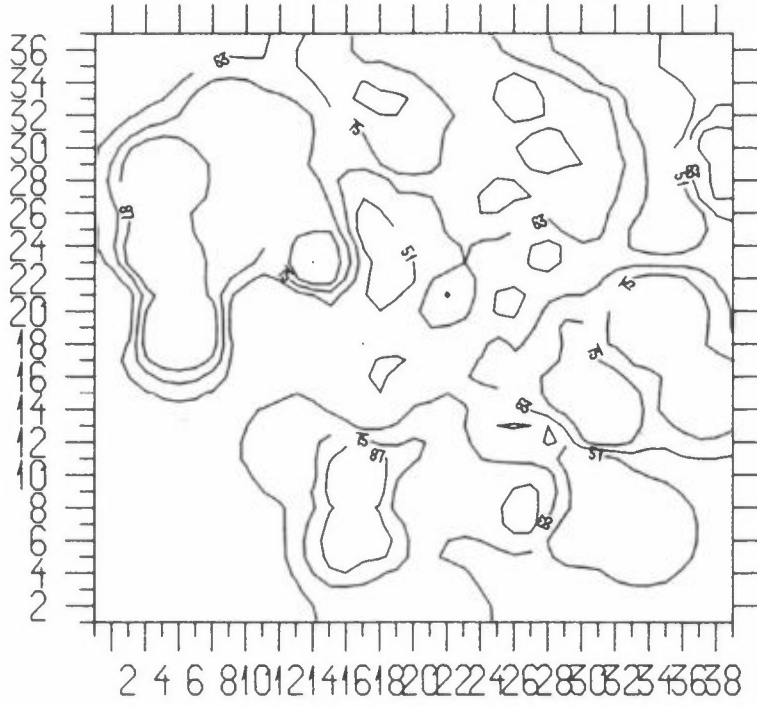


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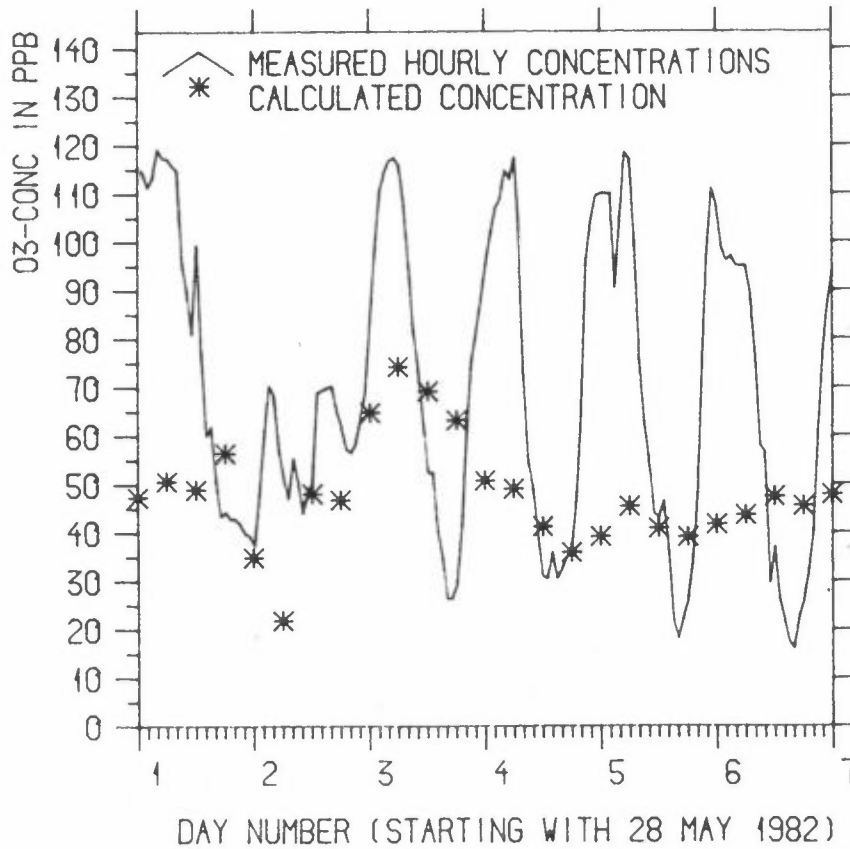
RELATIVE HUMIDITY



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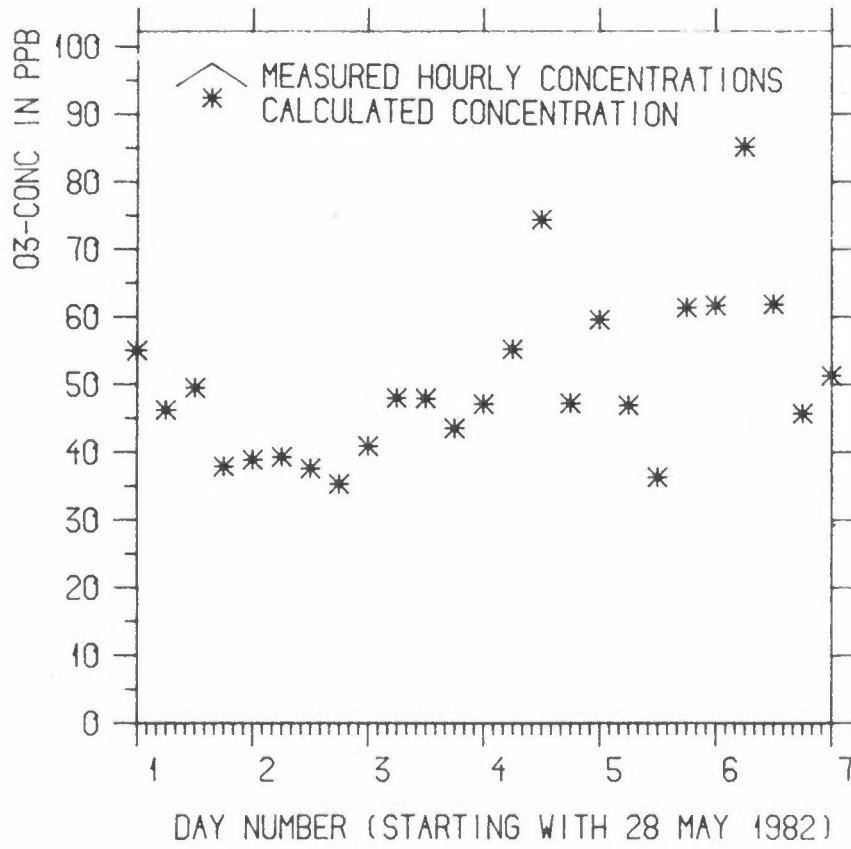
CONTOURS FROM 39. TO 99., INTERVAL 12.

ILLMITZ



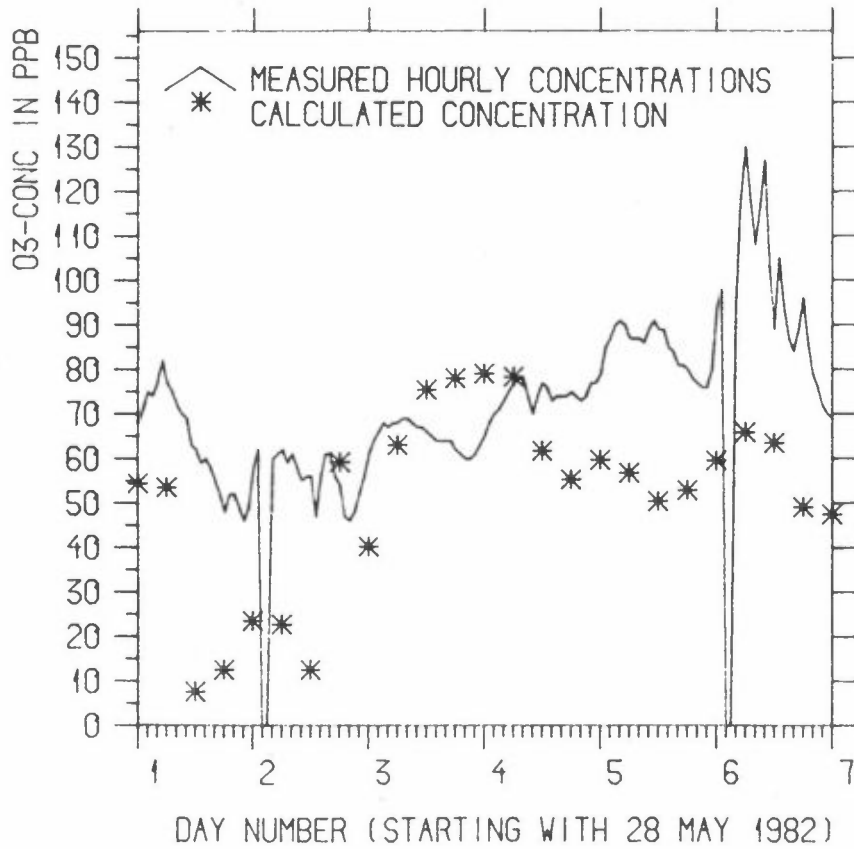
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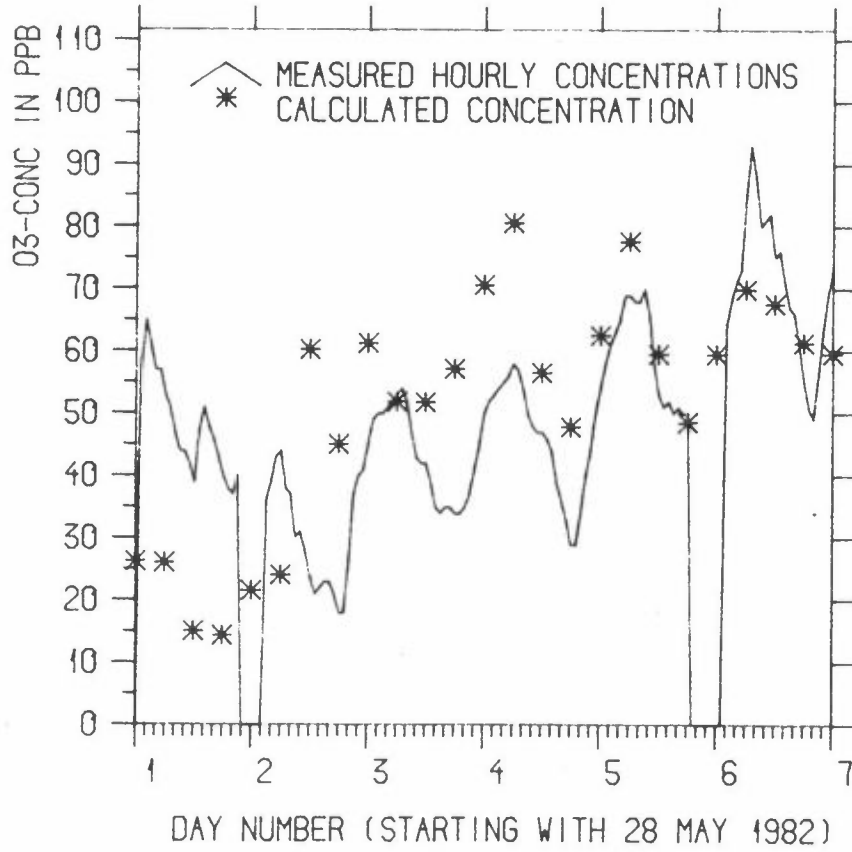
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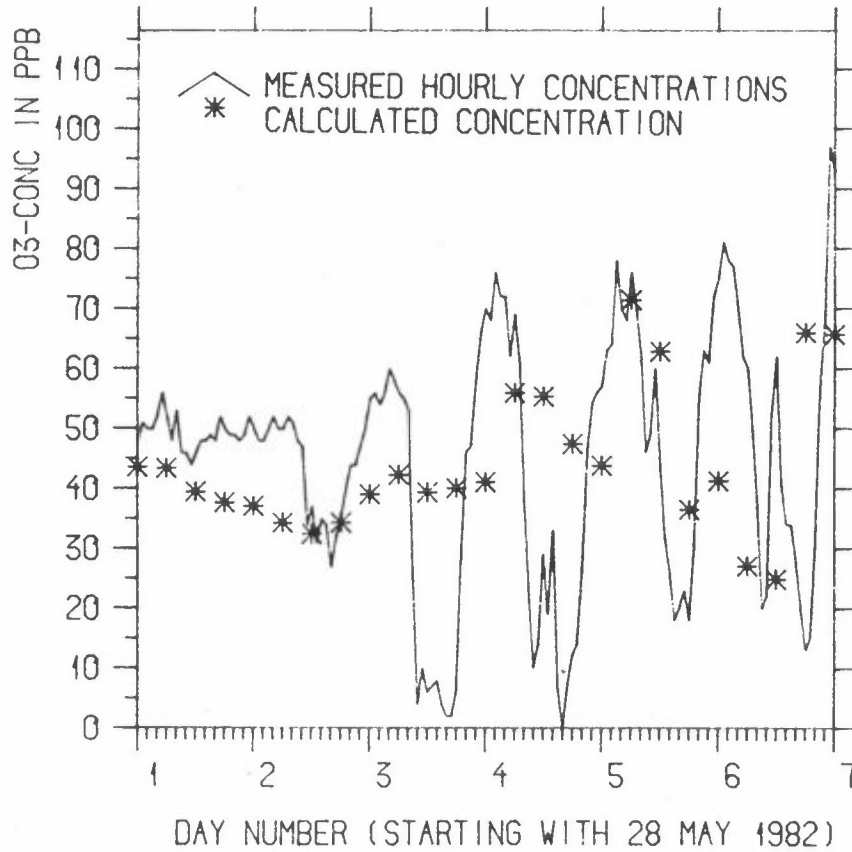
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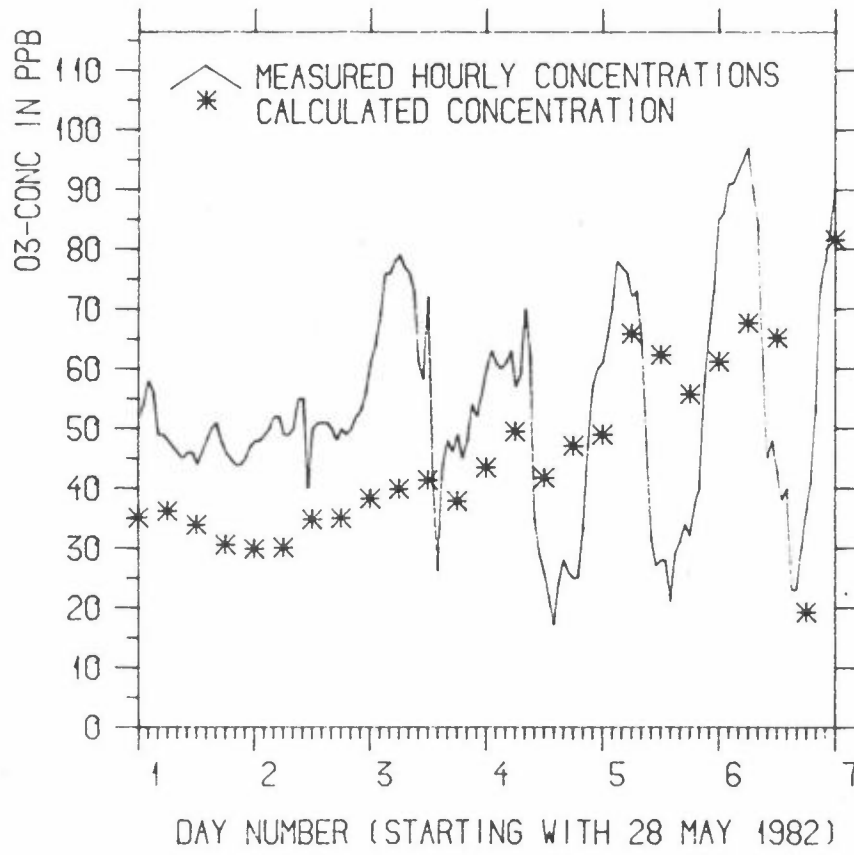
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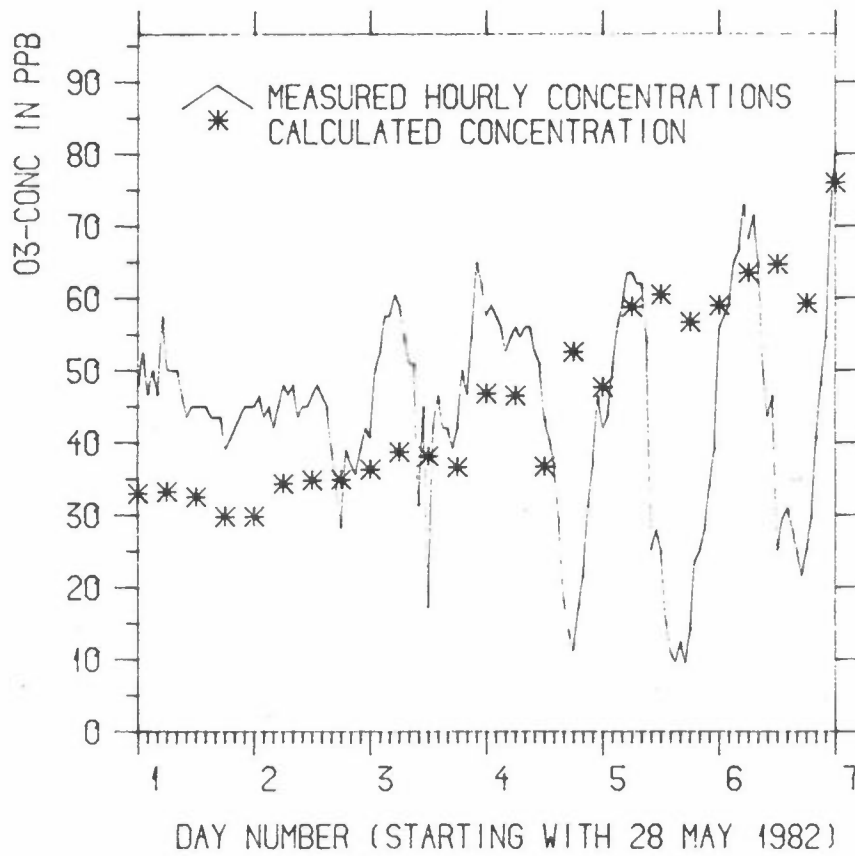
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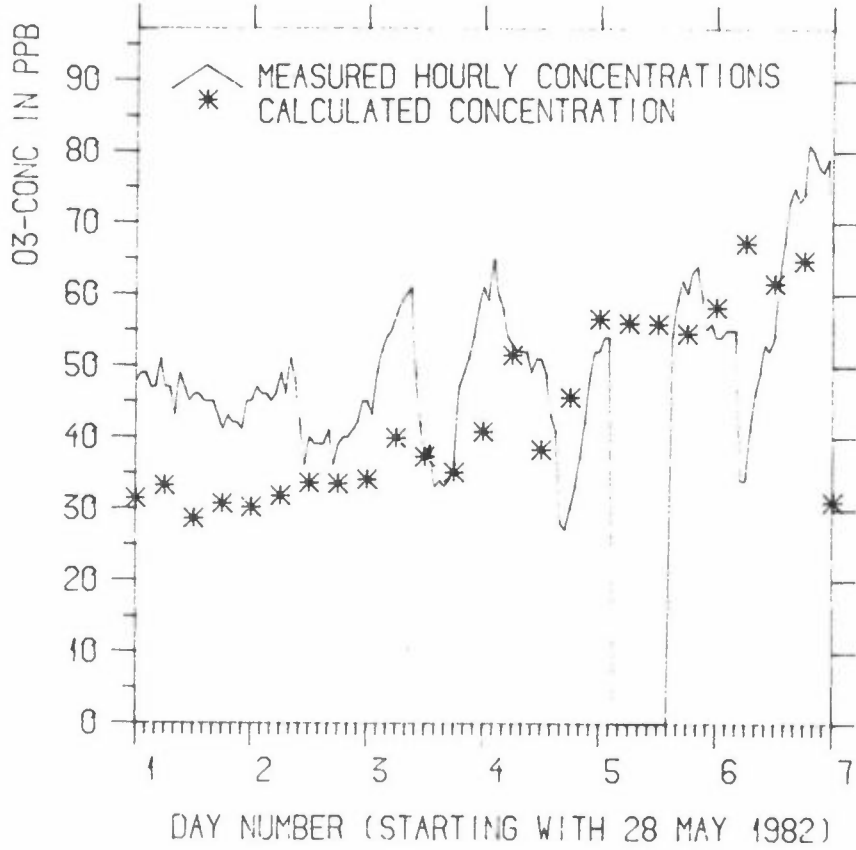
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LANGESUND



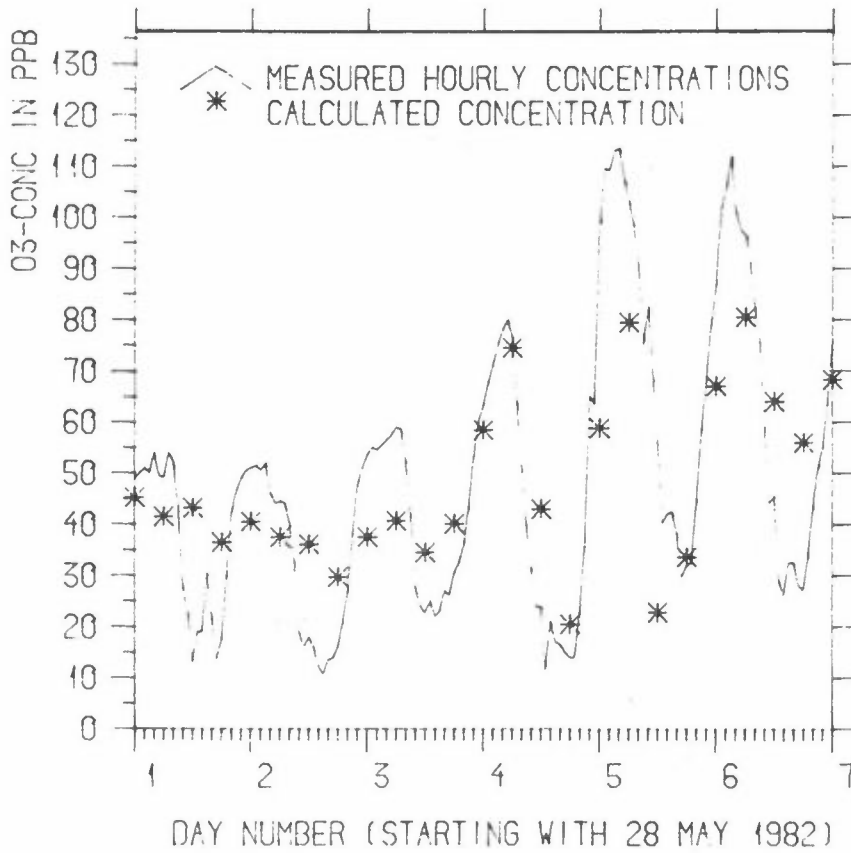
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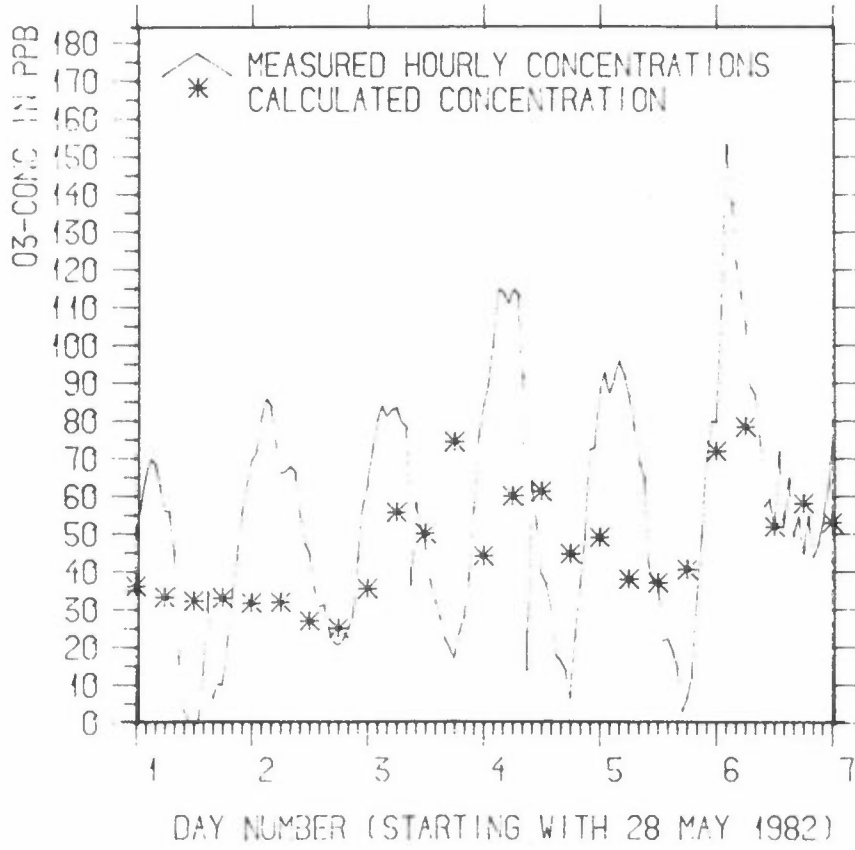
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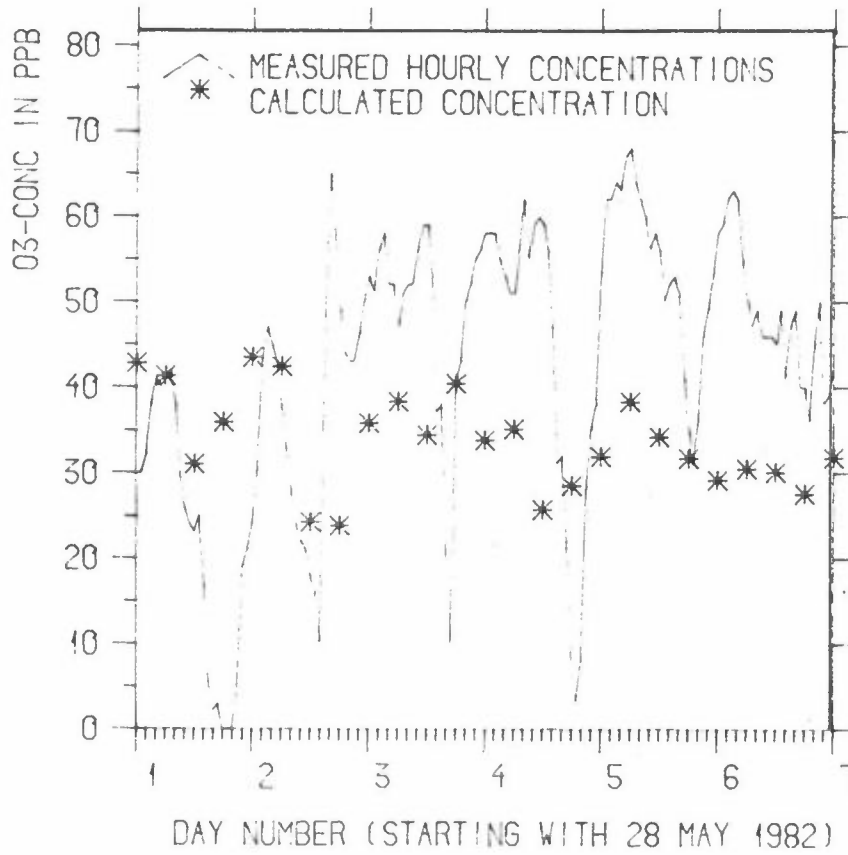
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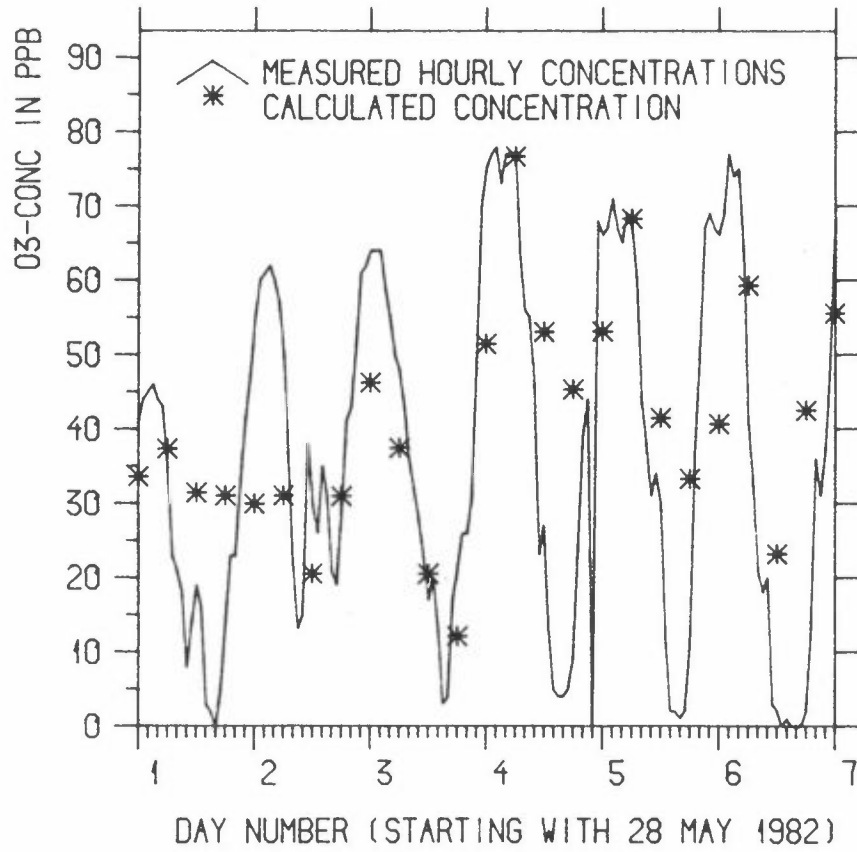
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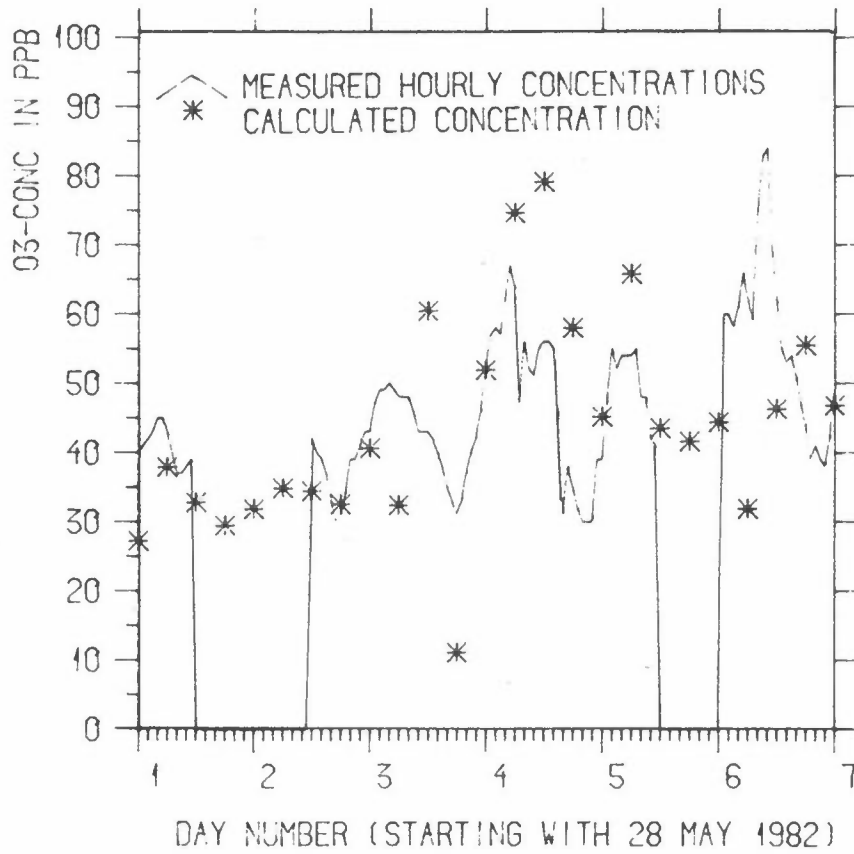
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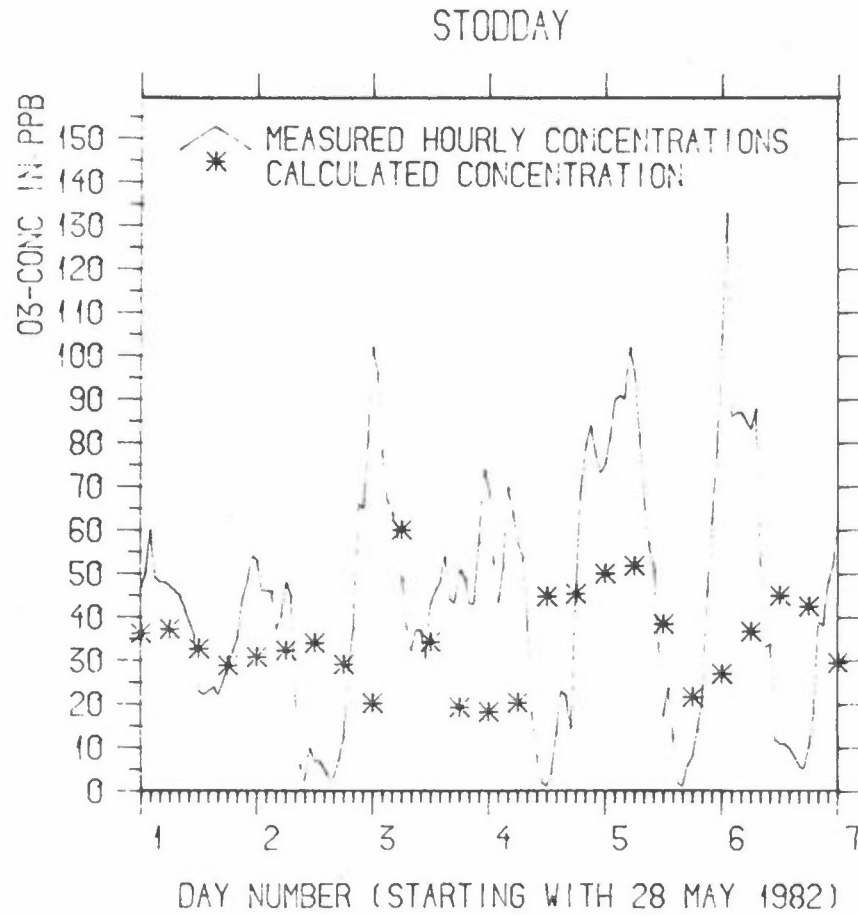


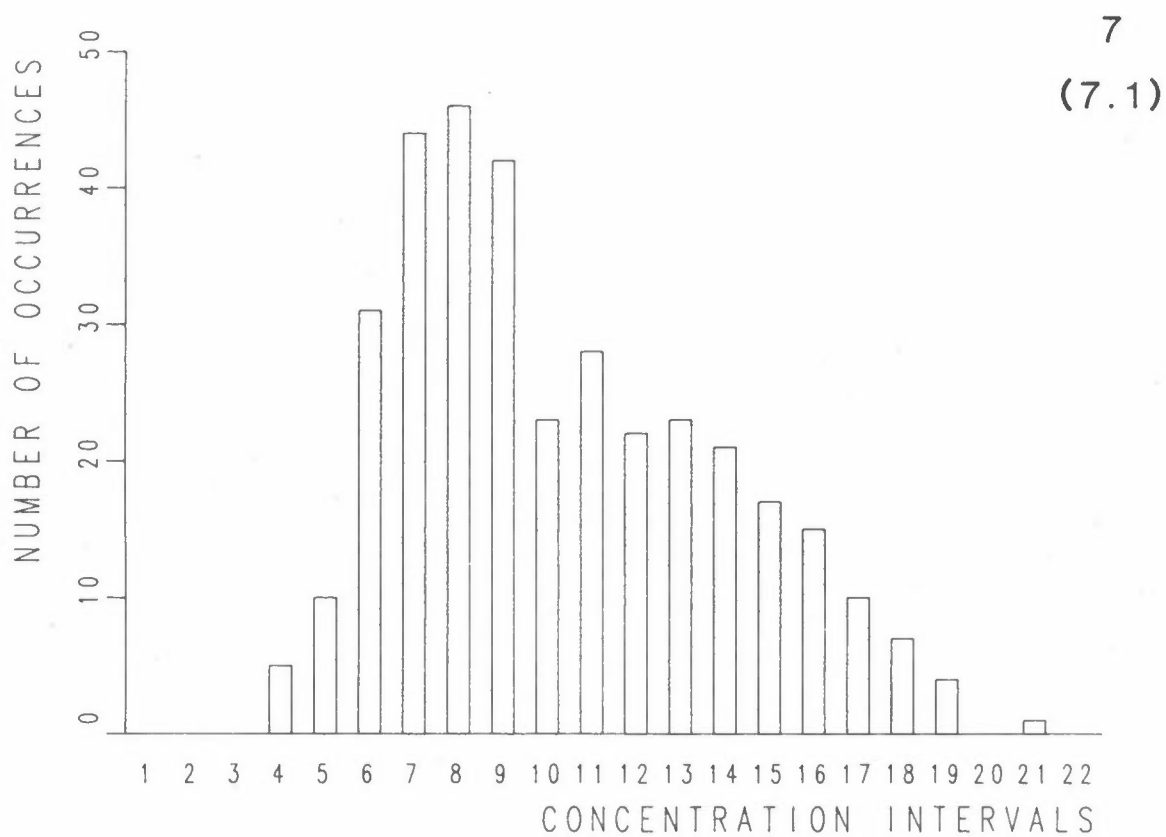
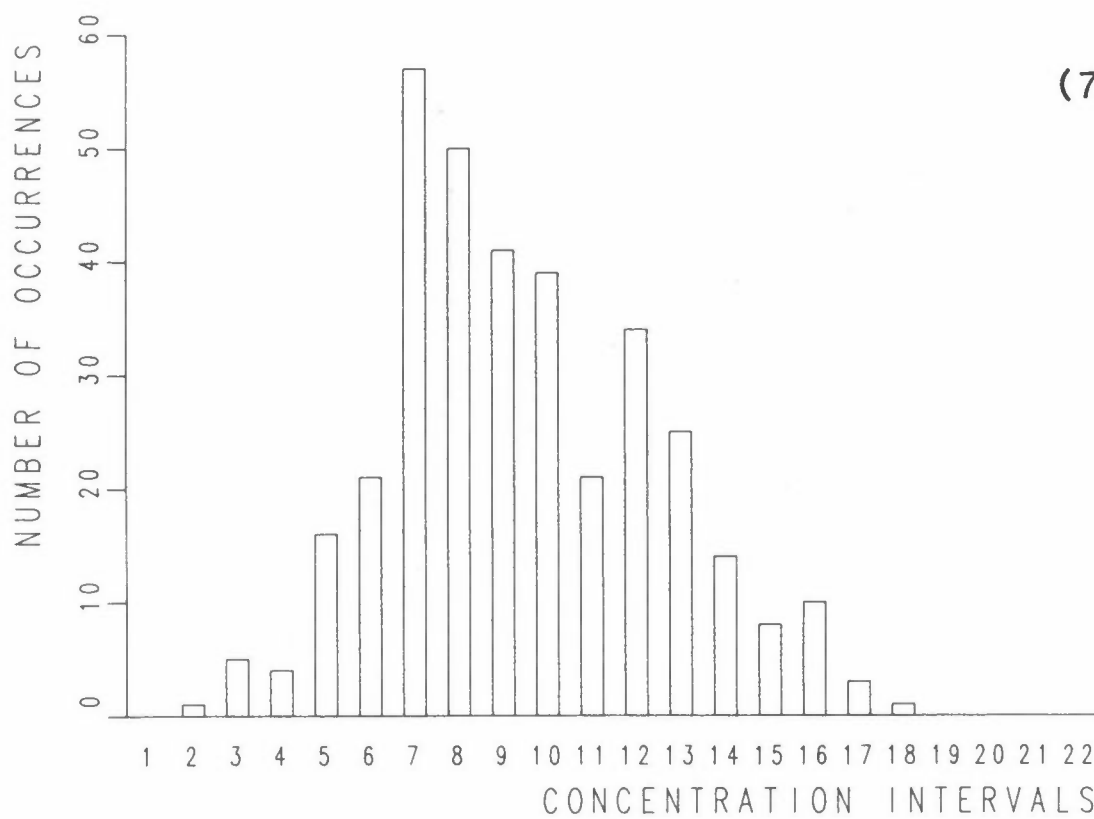
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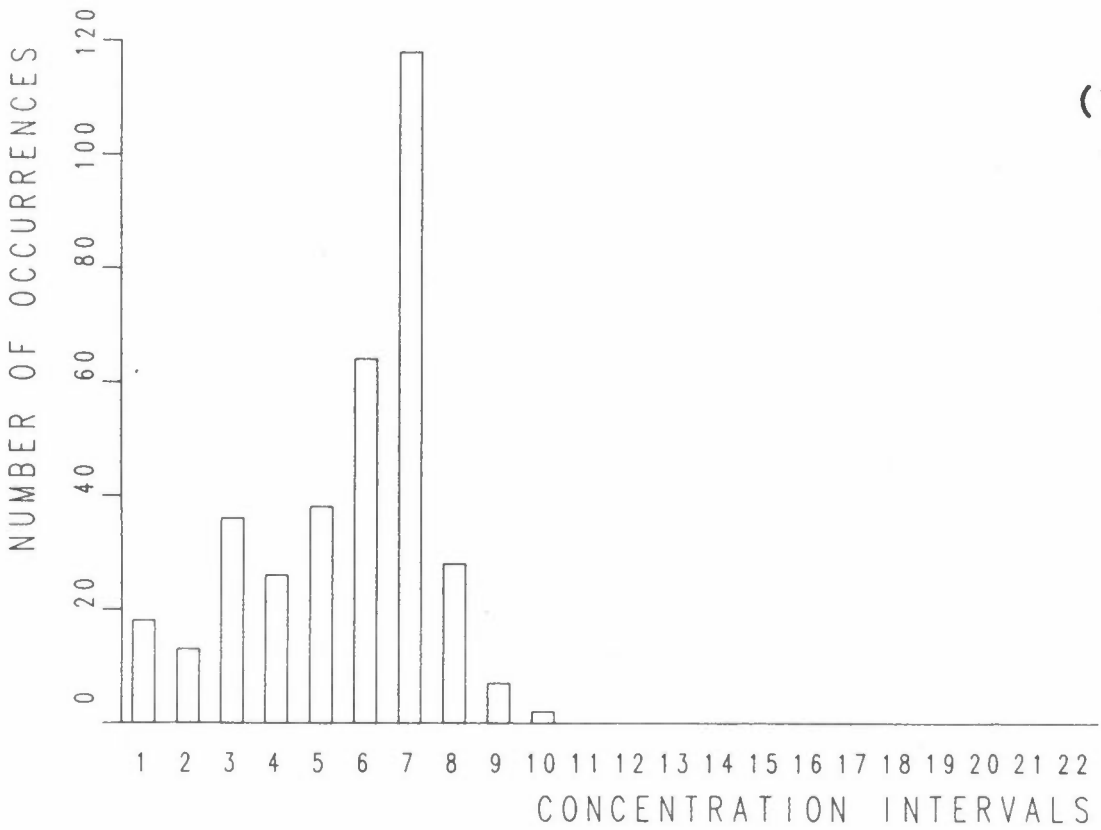
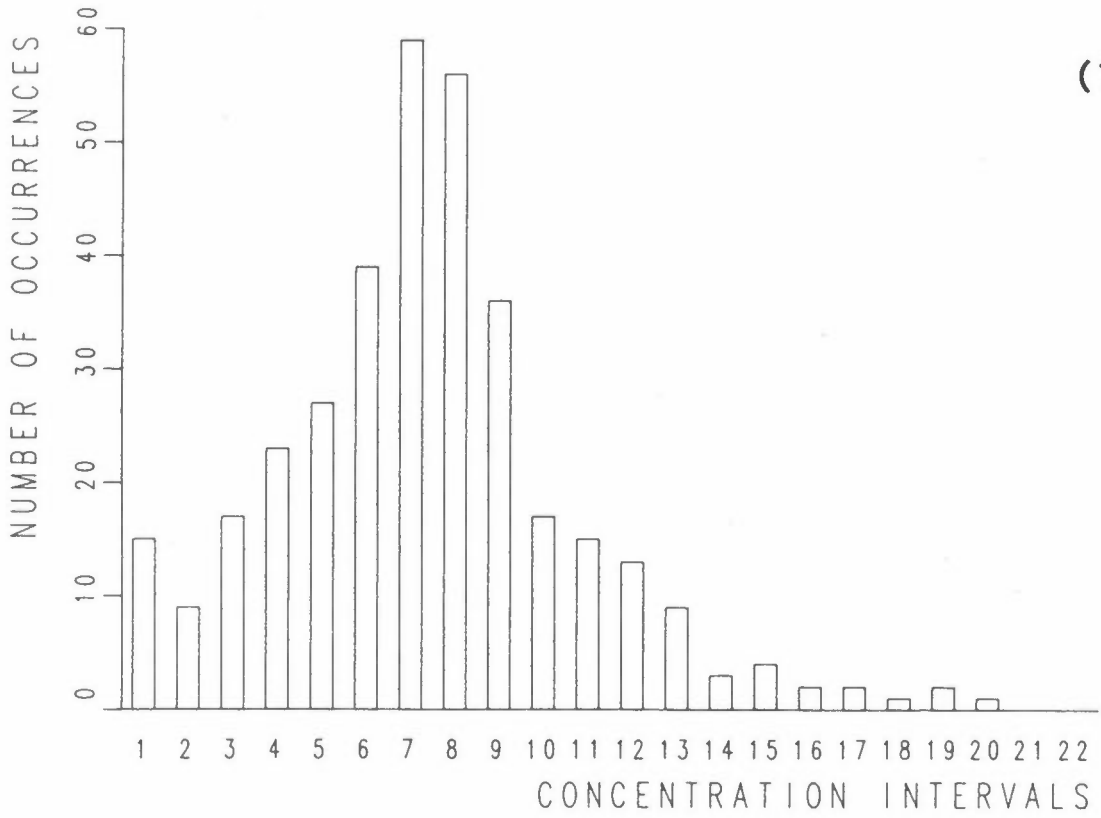
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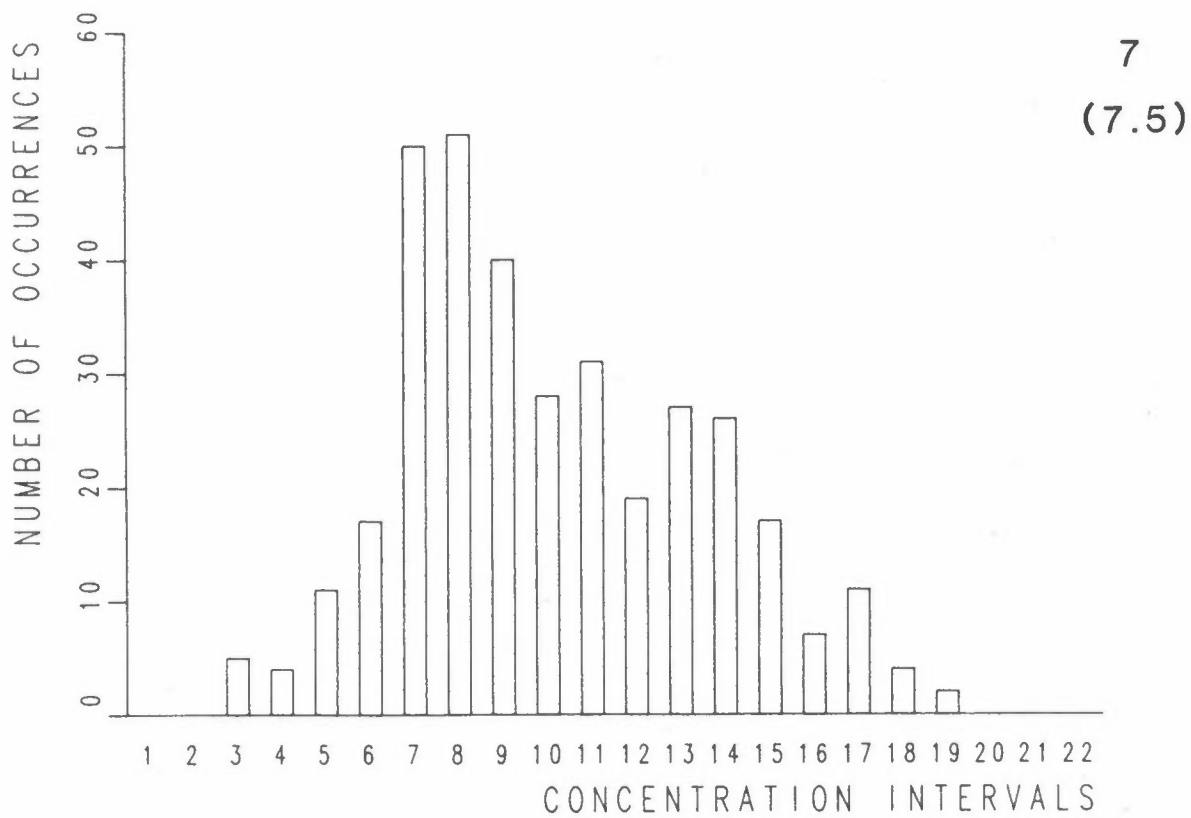
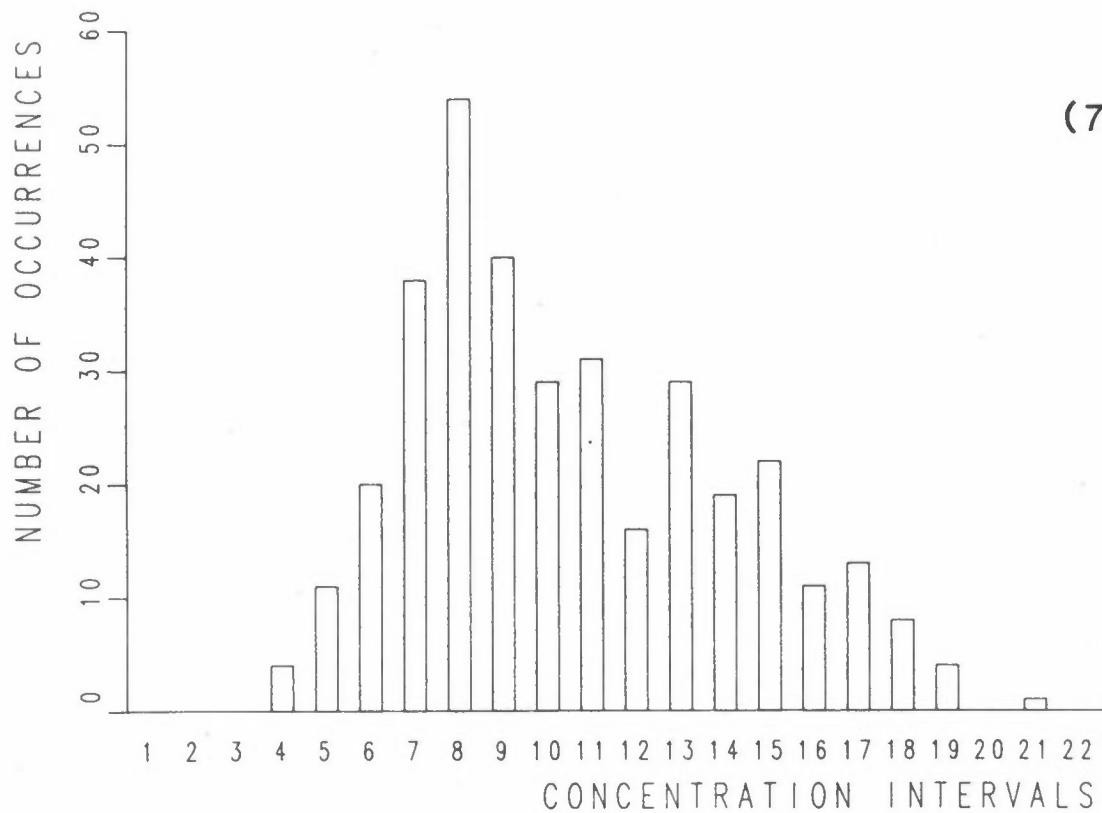


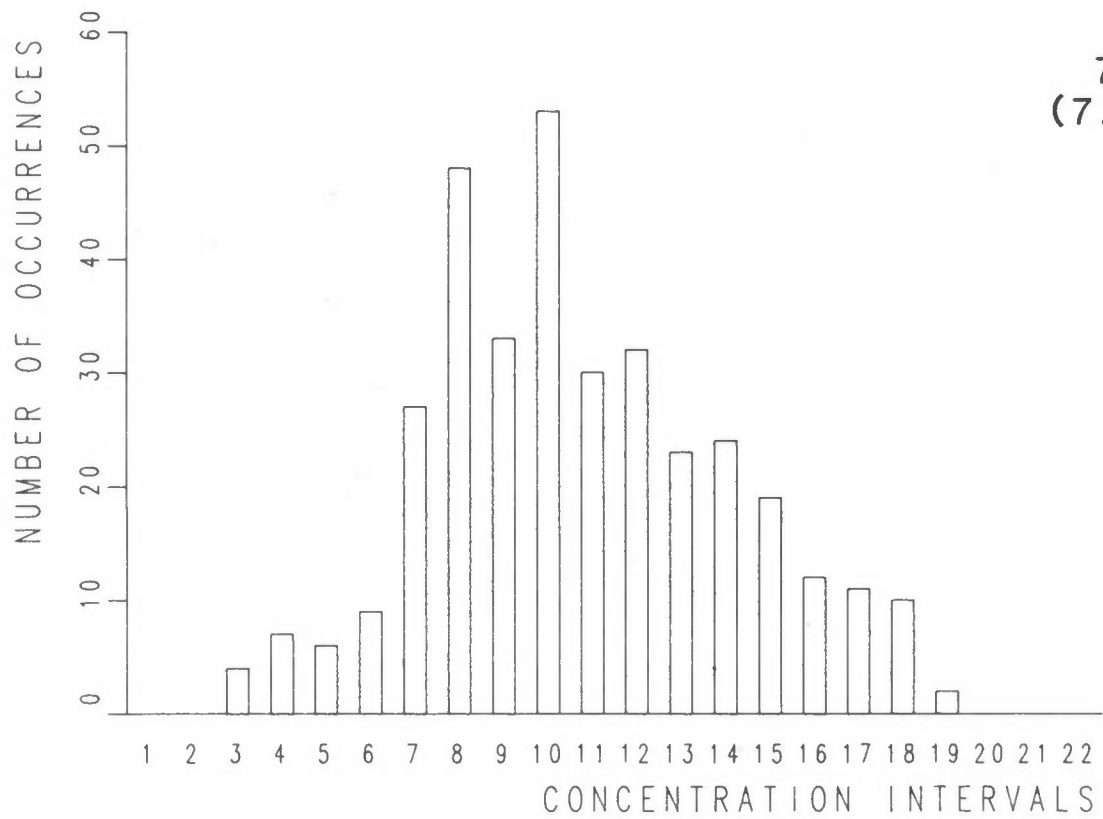
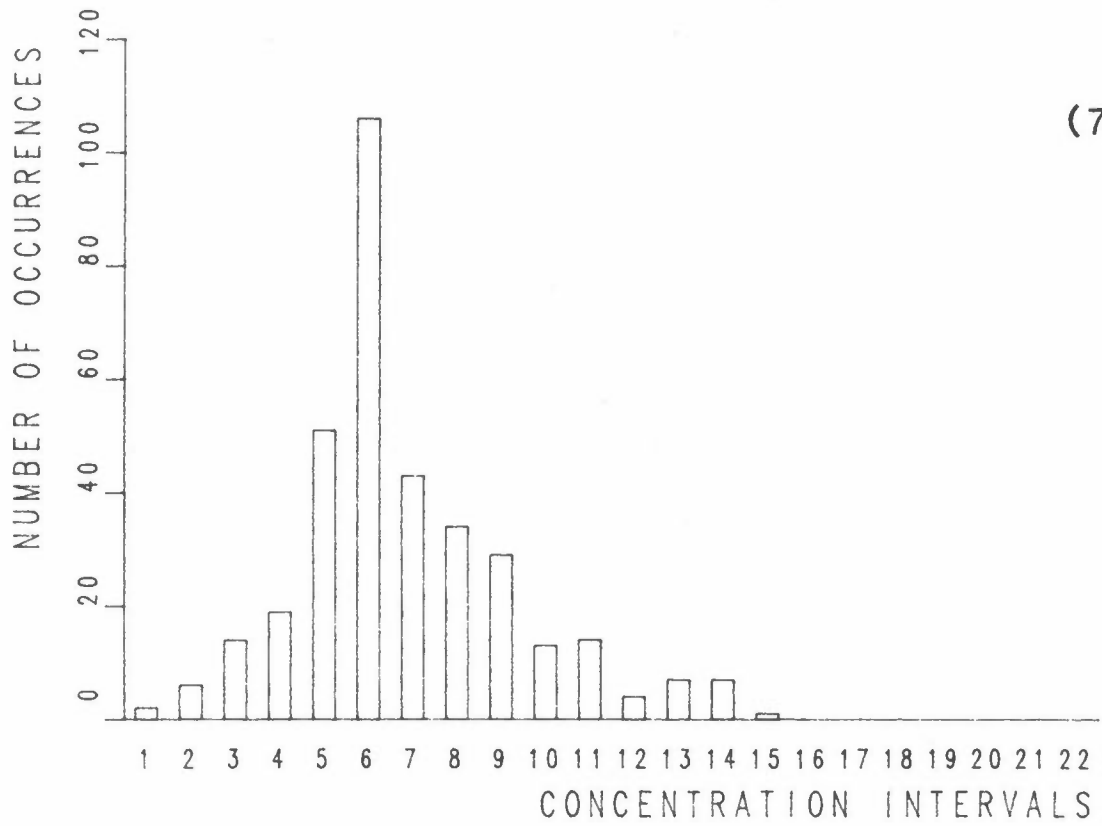
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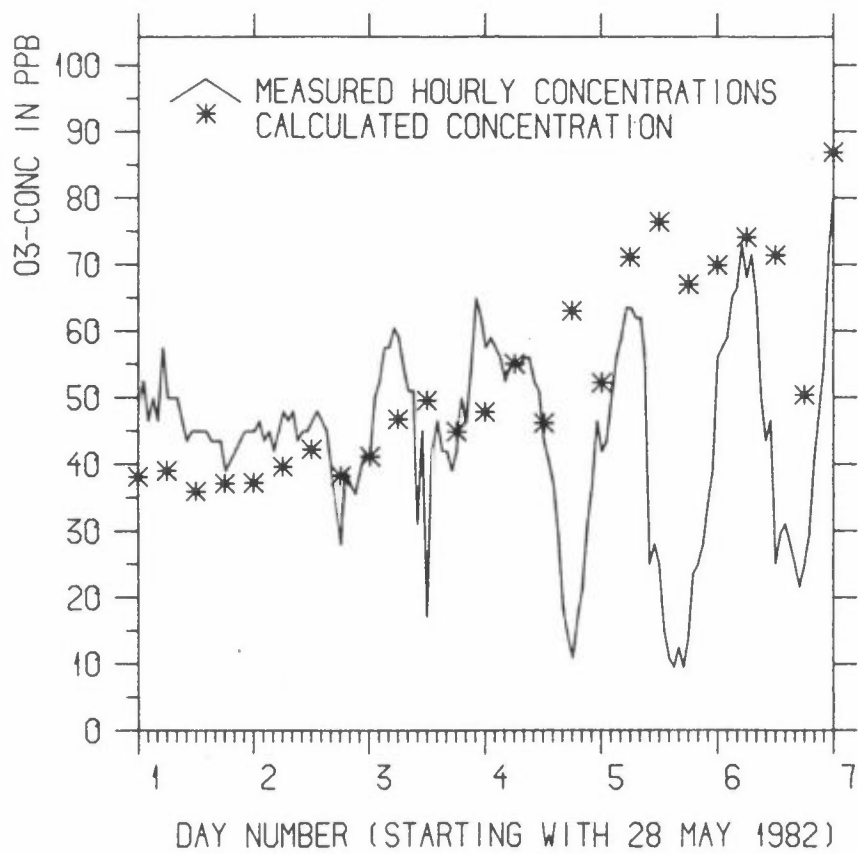






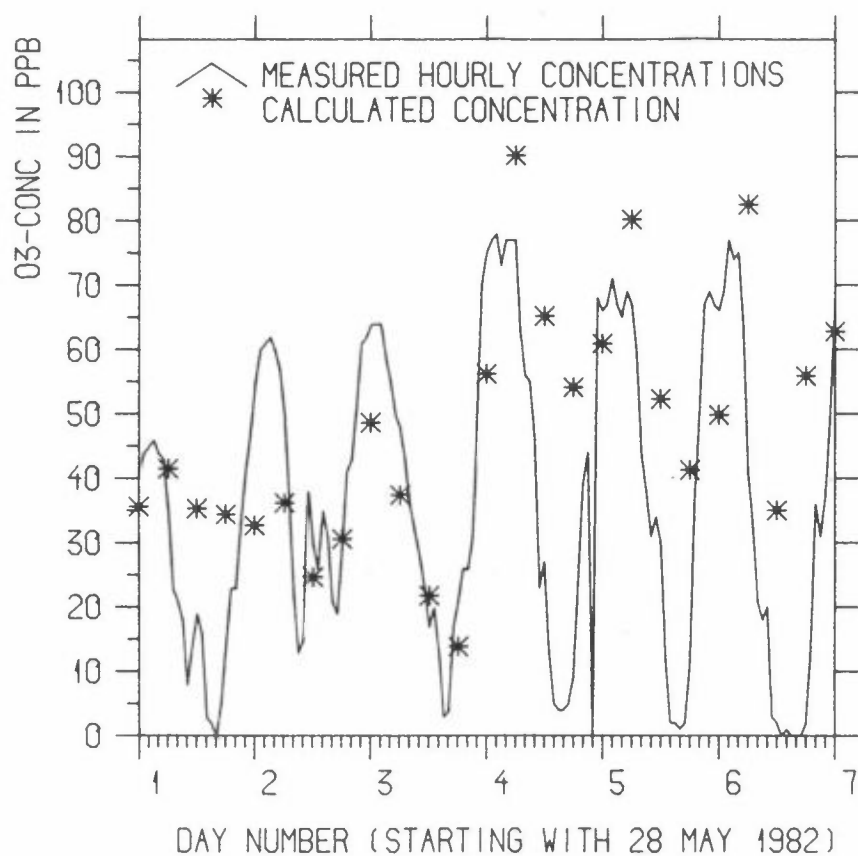


LANGESUND



8

BOTTFESFORD



8

NORSK INSTITUTT FOR LUFTFORSKNING (NILU)
 NORWEGIAN INSTITUTE FOR AIR RESEARCH
 POSTBOKS 64, N-2001 LILLESTRØM

RAPPORTTYPE OPPDRAGSRAPPORT	RAPPORTNR. OR 71/87	ISBN-82-7247-869-2	
DATO DECEMBER 1987	ANSV. SIGN. <i>J. Schjordepun</i>	ANT. SIDER 59	PRIS Kr 50.-
TITTEL Modelling of photochemical oxidants and precursors in Europe 28 May - 3 June, 1982.		PROSJEKTLEDER Øystein Hov	
		NILU PROSJEKT NR. O-8506 O-8624	
FORFATTER(E) Øystein Hov		TILGJENGELIGHET A	
		OPPDRAGSGIVERS REF.	
OPPDRAGSGIVER (NAVN OG ADRESSE) TNO, Apeldoorn, The Netherlands			
3 STIKKORD (à maks. 20 anslag) Photochemical oxidants Episodes Control strategies			
REFERAT (maks. 300 anslag, 7 linjer)			

TITLE
ABSTRACT (max. 300 characters, 7 lines) The Norwegian lagrangian long-range transport model with CBM-X chemistry has been applied to calculate the concentration of ozone and other photochemical oxidants at 14 measuring sites in Europe 28 May - 3 June 1982. Temperature, mixing height, cloud cover, dry removal and initial conditions are important for the calculated ozone concentrations. Reduction of VOC-emissions is calculated to reduce ozone. Reduction of NO _x or NO _x and VOC together may also reduce ozone.

* Kategorier: Åpen - kan bestilles fra NILU A
 Må bestilles gjennom oppdragsgiver B
 Kan ikke utleveres C