1	Plant mercury pump controls seasonal variations in global atmospheric mercury
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#### 20 Abstract

21 Anthropogenic mercury emissions are transported through the atmosphere as gaseous 22 elemental mercury (Hg(0)) prior to deposition to earth surfaces. Two explanations for strong 23 Hg(0) seasonality in the Northern Hemisphere (NH) are variations in primary anthropogenic 24 Hg(0) emissions, thought to peak in winter due to higher energy consumption, and faster 25 atmospheric Hg(0) oxidation rates in summer. Global Hg models simulate equally pronounced, oxidation-driven Hg(0) seasonality in the Southern Hemisphere (SH) where 26 27 Hg(0) levels are observed to be constant year-round. Furthermore, power consumption 28 related Hg emissions in N-America are higher in summer than in winter, which is inconsistent 29 with low N-American summertime Hg(0) concentrations. Here, we assess the role of Hg(0) 30 uptake by plants as <u>an alternative</u> driving mechanism for global Hg(0) variability. We find that data from all terrestrial background sites in the NH show a co-variation of Hg(0) with CO<sub>2</sub>, 31 32 which is known to exhibit concentration minima in summer due to CO<sub>2</sub> assimilation by 33 vegetation. The amplitude of seasonal Hg(0) oscillations <u>increase with latitude in the</u> NH and are larger at inland terrestrial background sites compared to coastal sites. We find that 34 35 satellite-derived vegetation photosynthetic activity (NDVI) correlates with Hg(0) levels at 36 individual sites and across continents. These findings <u>suggest</u> vegetation <u>to be a global Hg(0)</u> 37 pump, <u>that</u> dominates seasonal variations of atmospheric Hg(0). We suggest that decreasing Hg(0) levels in the NH over the last 20 years were partly caused by increased Hg(0) uptake by 38 39 vegetation due to increased terrestrial net primary production.

#### 41 Atmospheric mercury cycling

Gaseous elemental mercury (Hg(0)) is the dominant form of natural and anthropogenic Hg emissions and is transported globally through the atmosphere<sup>1</sup>. Hg deposition to earth surfaces occurs either by direct Hg(0) dry deposition or after oxidation to reactive Hg<sup>II</sup>, which is scavenged from the atmosphere by wet and dry deposition.<sup>2</sup> In addition to primary anthropogenic emissions, Hg is also re-emitted as Hg(0) from <u>land and ocean</u> surfaces to the atmosphere after reduction of Hg(II) pools in terrestrial and aquatic ecosystems<sup>2</sup>.

48 The short-term balance between Hg emission, deposition and re-emission processes governs 49 diurnal and seasonal Hg(0) variations. For example, strong seasonal variations in atmospheric 50 Hg(0) concentration have been described for most background sites in the temperate 51 Northern Hemisphere (NH) with peaks in winter and minima in summer.<sup>3-6</sup> Two widely 52 accepted hypotheses for the observed Hg(0) seasonality exist. The first hypothesis assumes 53 that Hg(0) variations are directly related to primary anthropogenic Hg(0) emissions from coal 54 combustion which are thought to be higher in winter due to higher energy demands for 55 heating.<sup>3,5</sup> Global Hg models have not vet been able to test this hypothesis, as current 56 anthropogenic Hg emission inventories have no seasonal resolution and are kept constant 57 throughout the year in models.<sup>7-9</sup> The second hypothesis suggests that atmospheric Hg(0) 58 oscilations are due to variations in atmospheric Hg(0) oxidation rate and subsequent Hg(II) 59 deposition <u>which are both</u> highest in <u>summer in global Hg models</u>.<sup>7-10</sup> <u>M</u>odels incorporating 60 atmospheric Hg(0) oxidation <u>simulate equally-pronounced (amplitudes of approx. 18%)</u> seasonal Hg(0) variations in both Hemispheres. Hg(0) measurements at temperate sites in the 61 62 Southern Hemisphere (SH) however are constant with no seasonal oscillations.<sup>6,11</sup> Methane 63 concentrations in the SH, for which atmospheric oxidation by OH radicals represents the most 64 important sink, do show strong seasonal variations.<sup>12</sup> The absence of Hg(0) seasonality in the 65 SH questions the relative importance of atmospheric Hg(0) oxidation.

Foliar uptake of atmospheric Hg(0) has long been recognized as a potentially inportant
pathway for atmospheric Hg deposition.<sup>13,14</sup> The magnitude of terrestrial-atmosphere net
Hg(0) exchange (foliar Hg(0) uptake - Hg(0) re-emission) however has been debated and most
studies suggested a net emission of Hg(0).<sup>15-18</sup> A recent review of Hg(0) flux measurements
over terrestrial surfaces concludes that there is no consensus whether terrestrial ecosystems
are a net sink or source for Hg(0) (-513 to 1651 Mg a<sup>-1</sup>, 25 to 75 percentile)<sup>19</sup>. Most Hg(0) flux
studies however focused on soil re-emission, and thus neglected foliar Hg(0) exchange.<sup>19</sup>

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73 There are several lines of evidence for net Hg(0) deposition to terrestrial ecosystems 74 that all invoke direct plant uptake of atmospheric Hg(0). First, foliage tissue Hg concentrations increase continously over the growing season, implying that Hg(0) is taken up by plants 75 76 through gas exchange and sequestered in leaf tissue.<sup>14,20</sup> Structurally incorporated Hg in leaf 77 tissue derived from atmospheric Hg(0) therefore likely has a lower re-emission potential due 78 to photo-reduction,<sup>20,21</sup> compared to Hg(II) deposited onto external leaf surfaces.<sup>22</sup> Second, 79 global litterfall Hg deposition to soils - representing Hg(0) net uptake by foliage - is estimated 80 to be 1000-1200 Mg yr<sup>-1</sup> <sup>17,23</sup> representing half of primary anthropogenic emissions.<sup>24</sup> Third, 81 vegetation and soil Hg stable isotope signatures, which discern foliar Hg(0) uptake from Hg(II) 82 wet deposition, put firm constraints on plant Hg(0) uptake dominating (50-80%) Hg 83 deposition at terrestrial sites. <sup>21,25-27</sup> Here, we investigate the role of Hg(0) uptake by plants in modulating global 84 85 atmospheric Hg(0) seasonality. We compare the atmospheric dynamics of Hg(0) with those of 86  $CO_2$  - a tracer for gas exchange by plants - for which seasonal oscillations in the Northern 87 Hemisphere (NH) are attributed to the balance of CO<sub>2</sub> assimilation and respiration during the 88 vear.<sup>28-30</sup>

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#### 90 Correlation of Hg(0) and CO<sub>2</sub> seasonality

91 We <u>investigated</u> multi-year time series of Hg(0) and CO<sub>2</sub> at seven <u>NH and SH background sites</u> 92  $(mean Hg(0) < 2 ng/m^3)$  and Hg(0) at an additional 43 global sites. Multi-year Hg(0) and CO<sub>2</sub> 93 observations of the five NH background sites, normalized for linear long-term trends, show 94 seasonal variations with highest concentrations in winter and minima in summer and fall (Fig. 95 **1**, **S1**, **S3** and **S6**). For all NH sites, we found a significant positive correlation between Hg(0) 96 and CO<sub>2</sub> (Pearson's product-moment correlation of monthly means, normalized for long-term 97 trends, Birkenes, Norway (r=0.37, p<0.005), Schauinsland, Germany (r=0.50, p<0.001), Mace Head, Ireland (r=0.40, p<0.001), Pallas, Finland (r=0.54, p<0.001) and Egbert, Canada (r=0.53, 98 99 p<0.001)). We investigated two coastal SH sites, Cape Point (South Africa) and Amsterdam 100 Island (France), with the latter surrounded by 3000 km of ocean on all sides.\_Cape Point 101 showed a weaker (r=0.30, p<0.01) correlation and no clear seasonal variation and Amsterdam 102 Island\_did not show a significant correlation between the two trace gases. Both SH sites show 103 small day-to-day variability, small diurnal variation and small seasonal amplitudes in  $CO_2$ , (<3 104 ppm, <1%) and Hg(0) (<0.1 ng/m<sup>3</sup>, <10%) (Fig. 1). This suggests low anthropogenic

contributions (i.e., no pollution plumes) as well as atmosphere-ocean exchange of Hg(0) that
is relatively constant throughout the day and season<u>s</u>.

We find that the interhemispheric gradient in Hg(0) concentrations (ΔHg<sub>45°N-45°S</sub>) is
 largest in February (0.66 ng/m<sup>3</sup>) and smallest in September (0.43 ng/m<sup>3</sup>) with a yearly
 average of 0.54 ng/m<sup>3</sup> (Fig 2A). Two thirds of the primary anthropogenic emissions are to the
 NH,<sup>31</sup> which is the main explanation for the higher NH Hg(0) concentrations compared to the
 SH.<sup>6</sup>.

112 The <u>seasonal</u> amplitudes of Hg(0) measurements at <u>temperate <u>NH</u> background sites</u> 113 <u>are</u> much larger, averaging 0.39 ng/m<sup>3</sup> (equivalent to 25% of the average annual Hg(0) concentration of 1.5-1.7 ng/m<sup>3</sup>),<sup>6</sup> and were 6 times larger than the corresponding  $CO_2$ 114 115 amplitudes (4%). Excluding <u>Polar</u> sites (see below), we find a positive correlation (r=0.88, 116 p<0.01, Pearson's product-momentum correlation) between seasonal Hg(0) (Sep-Feb) and 117 CO<sub>2</sub> (Aug-Feb) amplitudes (Fig. 2C). Amplitudes of seasonal CO<sub>2</sub> oscillations are largest at high 118 NH latitudes and decrease towards the south with minimal oscillations in the SH, which has 119 been related to varying CO<sub>2</sub> exchange during the vegetation period in the NH, and fewer land 120 masses in the SH.<sup>28-30</sup> We find a significant correlation between the seasonal Hg(0) amplitudes 121 of all background sites (n=38) and latitude (R<sup>2</sup> = 0.42, p<0.001, Figure 2B). However, Hg(0) 122 <u>amplitudes</u> are confounded by additional factors. For example, we explain a lower amplitude 123 of Egbert (Ca)\_by relatively high contributions of polluted air masses throughout the year 124 (Figure S12). Comparing all investigated sites, we find significantly lower Hg(0) amplitudes at 125 coastal sites compared to inland background sites, both for Europe (0.25 ng/m<sup>3</sup> vs. 0.39 126 ng/m<sup>3</sup>, n=13, p<0.05, 2-sided t-test) and North America (0.27 ng/m<sup>3</sup> vs. 0.41 ng/m<sup>3</sup>, n=18, 127 p<0.05, 2-sided t-test). A special case is a different seasonality of Hg(0) in the Arctic, which is 128 typically characterized by spring minima due to sea-salt induced atmospheric mercury 129 depletion events (AMDEs) and an unusual summertime maximum.<sup>32</sup> Observations from SH 130 terrestrial background sites are few, yet the first measurements in Bariloche, Patagonia (Ar) 131 revealed seasonal Hg(0) oscillations with minima in SH summer and autumn.<sup>33</sup>

132 If seasonal variations in atmospheric Hg(0) were mainly due to oxidation rates and 133 ocean re-emission,<sup>7-9</sup> one would expect higher Hg(0) amplitudes at coastal sites and in the SH, 134 which is contrary to the patterns described above. On an ecosystem scale, studies have 135 reported enhanced Hg deposition fluxes during vegetation growth periods<sup>21,34</sup> and lower local 136 atmospheric Hg(0) levels over vegetated surfaces.<sup>35,26,36</sup> The fact that <u>the</u> highest seasonal 137 amplitudes <u>are</u> observed at inland terrestrial background sites — along with similar patterns

observed for CO<sub>2</sub> — argues for a dominant role <u>of vegetation Hg(0) uptake in modulating</u>
global atmospheric Hg(0) concentrations. <u>The NH is not only more populated but also has a</u>
<u>much larger land mass (68%) compared to the SH (32%), resulting in higher plant Hg(0)</u>
<u>uptake which may explain the weaker interhemispheric gradient during the NH growing</u>
<u>season.</u>

143 Most NH sites reveal a remarkable <u>1-month offset in Hg(0)</u> summertime <u>minima (Sep)</u> 144 compared to CO<sub>2</sub> (Aug) (Fig. S3, S6). Opposite oscillation in diurnal Hg(0) and CO<sub>2</sub> 145 concentrations (Fig. S4) and later seasonal minima of Hg(0) relative to  $CO_2$  suggests that Hg(0)146 uptake by plants is <u>continuing</u> during periods of net CO<sub>2</sub> respiration during nighttime and in fall, when the ecosystem exchange of  $CO_2$  turns from net assimilation to net respiration.<sup>37</sup> 147 148 Hg(0) thus follows a similar behavior as carbonyl sulfide (COS), for which seasonal patterns 149 are coupled to CO<sub>2</sub>, whereas diurnal variations are decoupled.<sup>38,39</sup> This suggests that Hg(0) 150 plant uptake is controlled by plant gas exchange as represented by CO<sub>2</sub> here. Foliar Hg(0) re-151 emission however is decoupled from CO<sub>2</sub> respiration and is likely limited. This is reflected in 152 increasing Hg concentrations in foliage over the growing season.  $\frac{13,20,35}{10}$  The observed Hg(0) 153 concentration minimum in September may be amplified by a shallower planetary boundary 154 layer (PBL) in autumn leading to a reduced mixing with background air.<sup>30</sup>

# 156 Variations in Hg(0) vs. vegetation Index

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157 We compared <u>monthly</u> Hg(0) concentrations with the satellite-based Normalized Differenced Vegetation Index (NDVI), representing the vegetation photosynthetic activity and thus a proxy 158 159 for plant gas exchange, at the location of the <u>44 NH</u> monitoring sites (0.1° resolution). We find 160 a significant correlation between monthly Hg(0) and and NDVI for for 22 of the 44 investigated 61 sites. For the NH background sites where Hg(0) and CO<sub>2</sub> was measured, NDVI significantly 162 correlated with both trace gases (Fig. 3, Table S1), with the exception of Hg(0) at Schauinsland, De. For example, at Pallas NDVI was able to explain 69 % and 82% of seasonal Hg(0) and CO<sub>2</sub> 163 164 variation, respectively. The correlation was strongest at background sites in Canada (R<sup>2</sup> > 0.75 165 for e.g. Burnt Island, or Kejimkijuk, Figure S8, Table S2), Scandinavia (R<sup>2</sup> > 0.6 for e.g. Andoya, 166 Bredkäle, or Vavihill, Fig. S9, Table S3) or for some Mountain sites in China ( $R^2 \approx 0.5$  for e.g. Mt. 167 Damai, Mt. Leigong, Mt. Chanbai, Fig. S10, Table S4). For urban sites in North America (e.g. Salt 168 Lake City, Houston or Detroit, Fig. S7, Table S2) or Asia (Beijing, Nanjing, Fig. S10, Table S4) 169 we did not find a significant correlation between Hg(0) and NDVI (red points in Fig. S8), 170 indicating that local anthropogenic emissions dominate Hg(0) levels at urban sites. The

171 relatively strong correlation between Hg(0) concentrations and <u>NDVI</u> at regional background 172 sites suggests that vegetation uptake is responsible for Hg(0) depletion in summer months, when vegetation photosynthetic activity and thus gas exchange of plants is strongest (high 173 174 NDVI). Note that NDVI is able to describe a higher degree of the Hg(0) variance compared to 175 CO<sub>2</sub> concentrations which we relate to the fact that CO<sub>2</sub> respiration from terrestrial 176 ecosystems is decoupled from Hg(0) re-emission. We therefore suggest that NDVI satellite 177 data might be a robust parameter to estimate terrestrial Hg(0) uptake fluxes on a spatial and 178 temporal scale. Field studies comparing Hg(0) fluxes with vegetation cover are needed in 179 order to validate this approach. A recent study found a strong correlation of soil Hg pools with NDVI across western North America,<sup>40</sup> suggesting a long-term control of the plant Hg(0) pump 180 on soil Hg pools, which is strongly supported by Hg stable isotope data. <sup>21,25-27</sup> 181

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# 183 <u>Seasonality of Hg(0) emissions approximated by coal consumption statistics</u>

184 <u>Coal combustion represents a major primary Hg(0) emission source to the atmosphere.<sup>31</sup> We</u> estimated the seasonal variation of anthropogenic Hg emissions (Table S4) by using monthly 185 186 coal consumption data reported for the United States, Europe (EU27) and China. The 187 estimates show a peak in Hg emission during the cold winter months attributed to increased 188 energy consumption for heating (Fig. 4, Table S4). In North America, there is a second peak in coal consumption and thus Hg(0) emission (Fig. 4, Table S4) during hot summer months that 189 190 exceeds the winter peak and is attributed to high energy consumption for air-conditioning of 191 buildings.<sup>41</sup> Hg(0) concentrations measured at three urban monitoring sites in North America 192 (Salt Lake City, Houston and Detroit)<sup>42,43</sup> closely follow our seasonal emission estimates 193 including a noticeable summertime Hg(0) increase (Fig. 4). Regional background sites in 194 North America however show a consistent summertime minimum (Fig. 4), which cannot be 195 explained by anthropogenic Hg emissions, implying that other processes drive the lower Hg(0) levels in summer. Inverse modeling, that optimized terrestrial-atmosphere Hg(0) 196 197 exchange fluxes to fit observational Hg(0) data, suggested enhanced summertime net Hg(0) 198 deposition over eastern North America<sup>44</sup> supporting the strong role of vegetation as a pump 199 for Hg(0). In Europe and Asia, the anthropogenic Hg(0) emission estimates based on coal 200 consumption data are highest in winter (Table S4) with no summertime increase and a seasonality similar to Hg(0) observations (Fig. 4). Higher Hg(0) concentrations during 201 summer months at urban sites in Asia have been attributed to enhanced re-volatilization of 202 203 deposited Hg.<sup>45</sup> For the Waldhof site in central Europe, regional models have found a strong

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204 correlation between emission and particulate mercury but not for Hg(0).<sup>46</sup> We therefore
 205 conclude that variations in primary anthropogenic emissions alone cannot explain the
 206 observed seasonal Hg(0) variation in the NH.

#### 208 Implications for global Hg cycling

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209 Hg(0) concentrations measured in the PBL at terrestrial background sites reflect both 210 deposition and emission processes. Therefore observed Hg(0) oscillations can be considered 211 as variations in net exchange (vegetation uptake <u>– soil/vegetation</u> re-emission). The strong 212 depletion of atmospheric Hg(0) observed at terrestrial background sites in summer, despite 213 highest solar radiation and therefore potential photo-reductive re-emission, suggests that 214 terrestrial ecosystems are net sinks for Hg(0). We therefore conclude that vegetation uptake of Hg(0) is large and dominates over other terrestrial emission and deposition processes at a 215 216 regional and global scale. Considering an atmospheric Hg(0) pool of  $\frac{4800}{4800}$  Mg, $\frac{47}{27}$  the 20% 217 amplitude of seasonal Hg(0) oscillations indicate that terrestrial ecosystems are drawing 218 approximately 1000 Mg a<sup>-1</sup> of Hg(0) from the atmosphere <u>via</u> the vegetation pump, which is 219 in agreement with foliage/litterfall estimates<sup>17,23</sup>. This suggests that approximately half of the 220 annual primary anthropogenic emissions are assimilated by terrestrial vegetation, where it is 221 efficiently retained against re-emission to the atmosphere, though susceptible to transfer via 222 soils to continental and coastal aquatic ecosystems.<sup>21</sup>

223 We suggest that the vegetation pump controls to a large extent diurnal and seasonal 224 cycling of atmospheric Hg(0) in the terrestrial PBL, which has large implications for global Hg 225 cycling and interpreting and forecasting long-term trends. The absence of Hg(0) seasonality 226 observed in the SH seems inconsistent with our current understanding of fast Hg(0) oxidation 227 in the atmosphere<sup>7.9.47</sup> A dominant role of plants in NH Hg(0) seasonality <u>may</u> imply that 228 atmospheric Hg(0) oxidation is less important than currently assumed. Average Hg(0) levels 229 in the NH have decreased by 20-40 % between 1990 and 2010 and were recently attributed 230 to a cumulative 30% ( $\approx 600$  Mg a<sup>-1</sup>) decrease in global primary anthropogenic Hg(0) 231 emissions.<sup>24</sup> Over the same time period, net primary production (NPP) is thought to have increased in the NH by 0.13-0.15 Pg C a<sup>-1</sup>.48.49 Assuming median foliar Hg concentrations of 24 232 233 ng  $g^{-1}$ ,  $\frac{17}{17}$  we estimate that Hg(0) deposition by NH vegetation uptake today has increased by 234  $\simeq 140$  Mg a<sup>-1</sup> relative to 1990. We therefore suggest that a significant fraction of the observed 235 decrease in NH atmospheric Hg(0) concentrations resulted from increased vegetation uptake of Hg(0). Recent estimates of gross primary production (GPP), based on atmospheric COS 236

237 cycling, suggested even larger increases in GPP by 31% during the twentieth century<sup>50</sup>, that 238 may have strongly counterbalanced estimated increases in Hg(0) emissions. Regional 239 reductions in vegetation cover by deforestation<sup>51</sup> or droughts<sup>49</sup> on the other hand would slow 240 down or cease the vegetation Hg(0) pump, potentially resulting in higher atmospheric Hg(0)241 concentrations. We therefore emphasize the need to incorporate seasonal and spatial 242 variability in vegetation uptake of Hg(0) into global Hg models to better assess its impact on 243 regional and global Hg cycling. Trends in vegetation cover should be incorporated in models 244 reconstructing past and predicting future Hg(0) levels. The effects of Hg(0) uptake by vegetation related to climate change and land-use change should be considered when 245 246 discussing mitigation strategies to reduce human Hg exposure. Finally, the paradigm shift we 247 propose regarding the dominant global atmospheric Hg deposition pathway demands revised Hg deposition monitoring strategies by environmental agencies. 248 249

#### 250 Data, Materials & Methods

251 Atmospheric CO<sub>2</sub> and Hg(0) measurements. The focus of the present analysis is on 252 atmospheric monitoring sites that simultaneously measured Hg(0) and CO<sub>2</sub> concentrations. 253 As atmospheric oxidized mercury species account for a minor fraction of total atmospheric Hg 254 (Hgtot) in the PBL, no difference was made between Hg(0) and Hgtot and all data are reported 255 as Hg(0). High resolution QA/QC controlled Hg(0) and CO<sub>2</sub> data were obtained from global air 256 monitoring databases or directly from the responsible site scientists. The European data are 257 harmonised and quality assured following the EMEP program and monitoring strategy, see 258 <u>Tørseth, et al. 52</u> Hg(0) concentrations at all sites except Pallas were measured at high 259 frequency (5-15 min) using Tekran Continous Mercury Vapor Analyzers 2537 and reported 260 as 1h averages. For Cape Point, SA and Amsterdam Island, Fr all Hg(0) measurements were 261 made in compliance with the GMOS quality assurance protocol (<u>www.gmos.eu</u>). Hg(0) 262 concentrations at Pallas were measured using a semi-automatic sampling on Au-traps (quartz 263 glass coated with gold) at an air sampling rate of 300 ml/min during 24h periods, two days 264 per week. The samples were analyzed on a Tekran 2500 CVAFS Mercury Detector at the IVL 265 laboratory in Gothenburg. CO<sub>2</sub> at Birkenes was measured with Cavity Ringdown spectroscopy 266 (CRDS), Picarro G1301 with 5 s resolution, average to 1 h if data coverage was more than 75%. 267 The data are calibrated against WMO- Global Atmospheric Watch standards at EMPA. CO<sub>2</sub> data 268 from Amsterdam Island, Fr was measured using CRDS and from Egbert, Ca and Cape Point, SA 269 using Non-Dispersive Infrared (NDIR) sensors. CO<sub>2</sub> data from Schauinsland, De were 270 measured by IR-Absorption. All data from Birkenes, No, Amsterdam Island, Fr, Egbert, Ca, 271 Cape Point, SA and Schauinsland, De were measured continuously and reported as hourly 272 averages. CO<sub>2</sub> data from Pallas, Fi and Mace Head, Ir were measured on an event base with 273 flasks and analysed using NDIR.53 Monthly average Hg(0) data were obtained from 4,42,43,45,54-<u>58,6,59-61</u> 274

Normal Differenced Vegetation Index, (NDVI) were obtained from from the NASA Earth
Observations (NEO) platform at 16 days and 0.1° resolution. NDVI data are derived from
Moderate Resolution Imaging Spectroradiometer (MODIS) images aboard the NASA's Terra
and Aqua satellites.

279 <u>Statistics</u>

The Co-variance between Hg(0) and CO<sub>2</sub> was assessed using Pearson's product moment correlation coefficient with the cor.test function of the stats (v3.1.3) package and long-term trends of time series as well as linear regression's of Hg(0) and CO2 with NDVI were assessed

- with the lm function of R.<sup>62</sup> Students t-tests were performed with Microsoft Excel assuming
  uneven variances.
- 285 <u>Seasonality of anthropogenic Hg emission.</u> The monthly Hg emission for the different regions
   286 (F(Hg)<sub>month</sub>) in Mg Hg/month was calculated according to Formula 1:
- 287  $F(Hg)_{month} = f_{coal} * f_{consumption} * FHg_{year} + (1 f_{coal}) * \frac{1}{12} * FHg_{year}$  (1)
- 288 where  $f_{coal}$  and F(Hg)<sub>vear</sub> corresponds to the fraction of Hg emission based on coal combustion 289 and the total yearly Hg emission for each investigated country, based on the AMAP/UNEP 290 2010 mercury emission estimates  $\frac{63}{5}$ .  $f_{consumption}$  corresponds to the monthly fraction of the yearly coal consumption. For the USA and Europe (27 EU countries) monthly coal 291 292 consumption data are reported by <u>governmental</u> agencies. Seasonality of Hg emission from 293 China was derived from the monthly power generation data between 2005 and 2010 reported 294 by Liu et al. 2015<sup>64</sup>. The monthly cement production in China was considered constant 295 throughout the year<sup>65</sup>.
- Normalization of monthly Hg concentration and emission data. In Figure 1 monthly Hg
   concentrations and emission fluxes are normalized to yearly averages for comparability. Hg
   concentrations are normalized as follows: C(Hg0)<sub>monthx</sub> /C(Hg0)<sub>year</sub>, were C(Hg0)<sub>monthx</sub>
   represents the multiyear monthly average of month x and C(Hg0)<sub>year</sub> the yearly average.
- 300 <u>Computation of Seasonal Amplitudes. The absolute seasonal amplitudes are calculated from</u>
- 301 <u>the multiyear monthly averages as follows:  $(\mu M_{max} \mu M_{min})$ , were  $\mu M_{max}$  and  $\mu M_{min}$  represent</u>
- the averages of the months with NH maximum (February for both Hg(0) and CO<sub>2</sub>) and
- 303 <u>minimum (August for CO<sub>2</sub> and September for Hg(0))</u> concentration (Table S6-S10),
- 304 <u>respectively. Relative amplitudes are calculated as follows:  $(\mu M_{max} \mu M_{min})/\mu Y$ , were  $\mu Y$ 305 <u>represents the yearly average.</u></u>
- 306

## 307 Data availability

- 308 Most European Hg(0) data used are associated to EMEP, and publically available through the
- 309 EMEP data base, EBAS: <u>http://ebas.nilu.no</u>. CO2 concentrations were obtained from the ESRL
- 310 NOAA Global Monitoring Division database
- 311 (ftp://aftp.cmdl.noaa.gov/data/trace\_gases/co2/flask/surface/co2\_mhd\_surface-
- 312 flask\_1\_ccgg\_event.txt,
- 313 ftp://aftp.cmdl.noaa.gov/data/trace\_gases/co2/flask/surface/co2\_pal\_surface-
- 314 flask\_1\_ccgg\_event.txt) and the World Data Center for Greenhouse Gases WDCGG
- 315 (http://ds.data.jma.go.jp/gmd/wdcgg/). Canadian Hg(0) concentrations were obtained from

- 316 the National Atmospheric Chemistry (NAtChem) Database and Analysis Facility of 317 Environment Canada (www.ec.gc.ca/natchem).
- 318 For the USA and Europe (27 EU countries) monthly coal consumption are reported by the U.S.
- 319 Energy Information Administration (http://www.eia.gov/totalenergy, downloaded 320 25.5.2016) and Eurostat (http://ec.europa.eu/eurostat/statistics-321 explained/index.php/Coal\_consumption\_statistics, downloaded 15.2.2016), respectively. NDVI data was obtained from the NASA Earth Observations (NEO) platform 322 323 (http://neo.sci.gsfc.nasa.gov)
- 324 325

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- 334

# 335 Author Contribution

MJ initiated the project, performed the data analysis and wrote the manuscript together with
JES and contributions <u>from</u> DO, JB and AD. JB, RE, CML, KA, IW, KK, DW, LM, CL, TM, MR, OM,
and AD provided data. All <u>Authors</u> read and commented on the manuscript.

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404 anomalies normalized for long-term linear trends. All data meeting QA/QC criteria of the

405 <u>respective site are presented.</u>

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- Fig. 2: <u>A: Interhemispherical Gradient of Hg(0) at background sites (<2 ng/m<sup>3</sup>). The lines</u>
   represent polynomial fits of the measurements for the yearly mean and the monthly means of
- 12 Elemente de la composition de la composition de version de version de la composition de la composit
- **February and September. B: Seasonal amplitude of Hg(0) (September mean February mean)**
- 413 as a function of latitude for coastal (squares) and terrestrial (circles) sites. The dashed line

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- 414 <u>represents the linear fit of the data and the shaded area the 95% confidence area. C: Seasonal</u>
- 415 <u>amplitudes of Hg(0) (Sep-Feb)</u> and CO<sub>2</sub> (Aug-Feb) variation measured at <u>7</u> atmospheric sites.
- **<u>416</u>** The error bars represent the interannual variability (1σ). The dashed line represents the linear</u>
- fit of the data and the shaded area the 95% confidence area.



Figure 2: Linear correlation between multi-year averages of monthly Hg(0)
 concentration and NDVI (left) and CO<sub>2</sub> concentration and NDVI (right) at five NH
 monitoring stations.



Fig. 4: Seasonal variation of Hg emissions in gray, atmospheric Hg(0) concentration at terrestrial background sites (Hg(0) Bg) in blue and at urban sites (Hg(0) urban) in red and Vegetation index (NDVI) in green. Monthly Hg emission and Hg(0) concentrations are reported

432	relative to yearly average. The bold lines represent averages and the shaded areas represent		
433	25th to 75th percentiles.		
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